
Revisiting the Coupled Behavior of the Subtropical Jet and Hadley Cell

MOLLY MENZEL

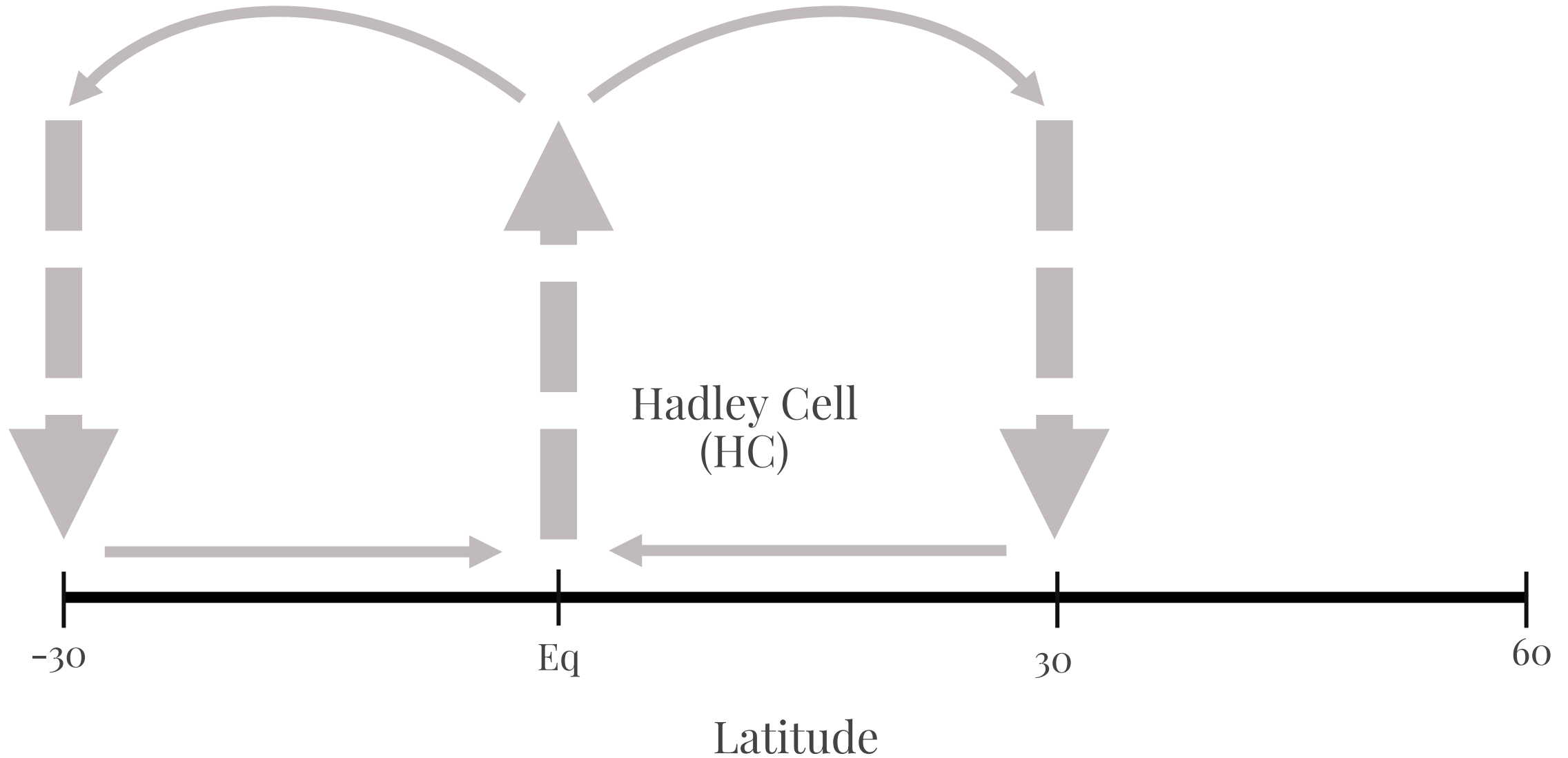
January 26, 2021

Advisor: Prof. Darryn Waugh
Johns Hopkins University

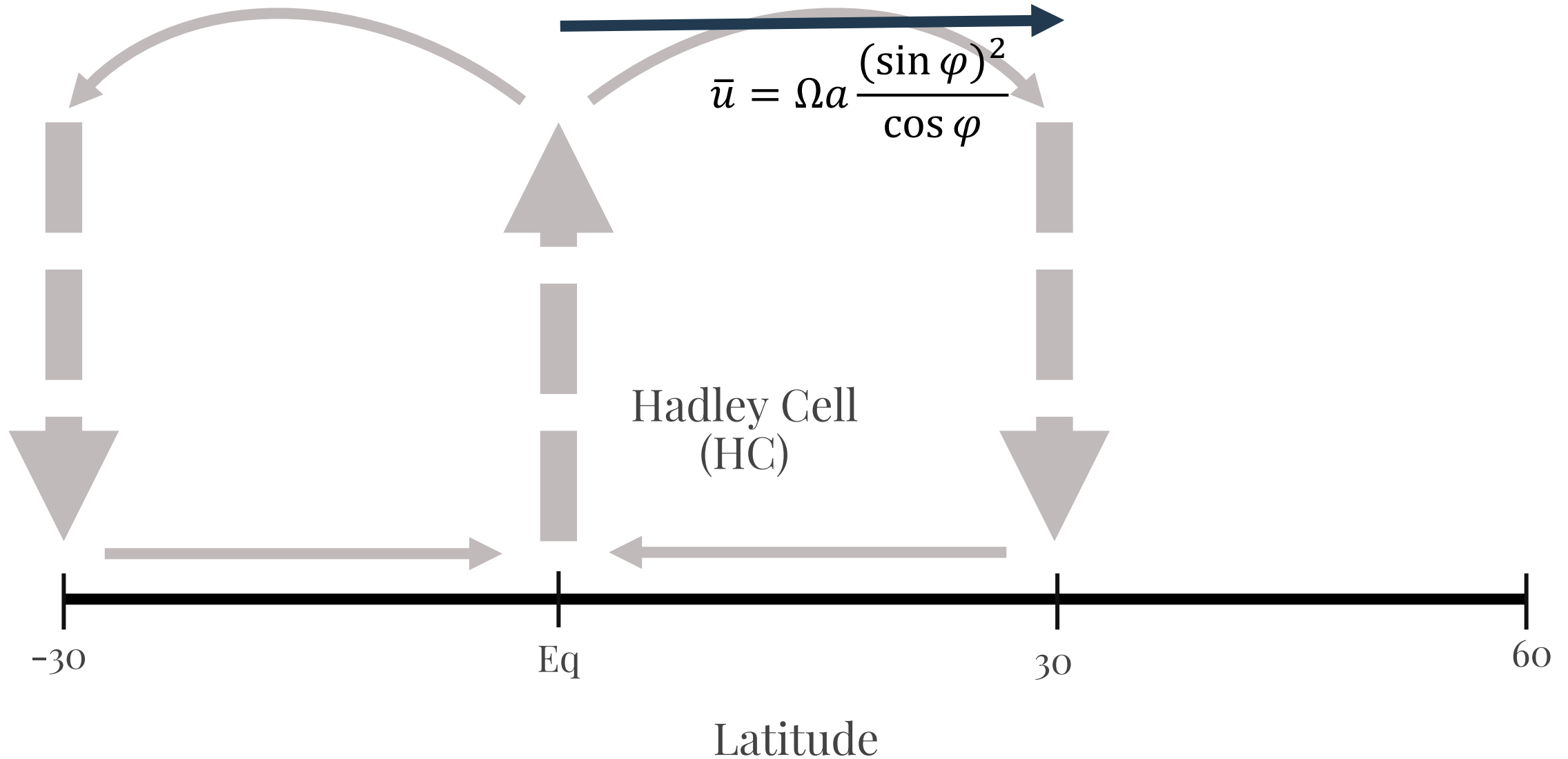
Introduction



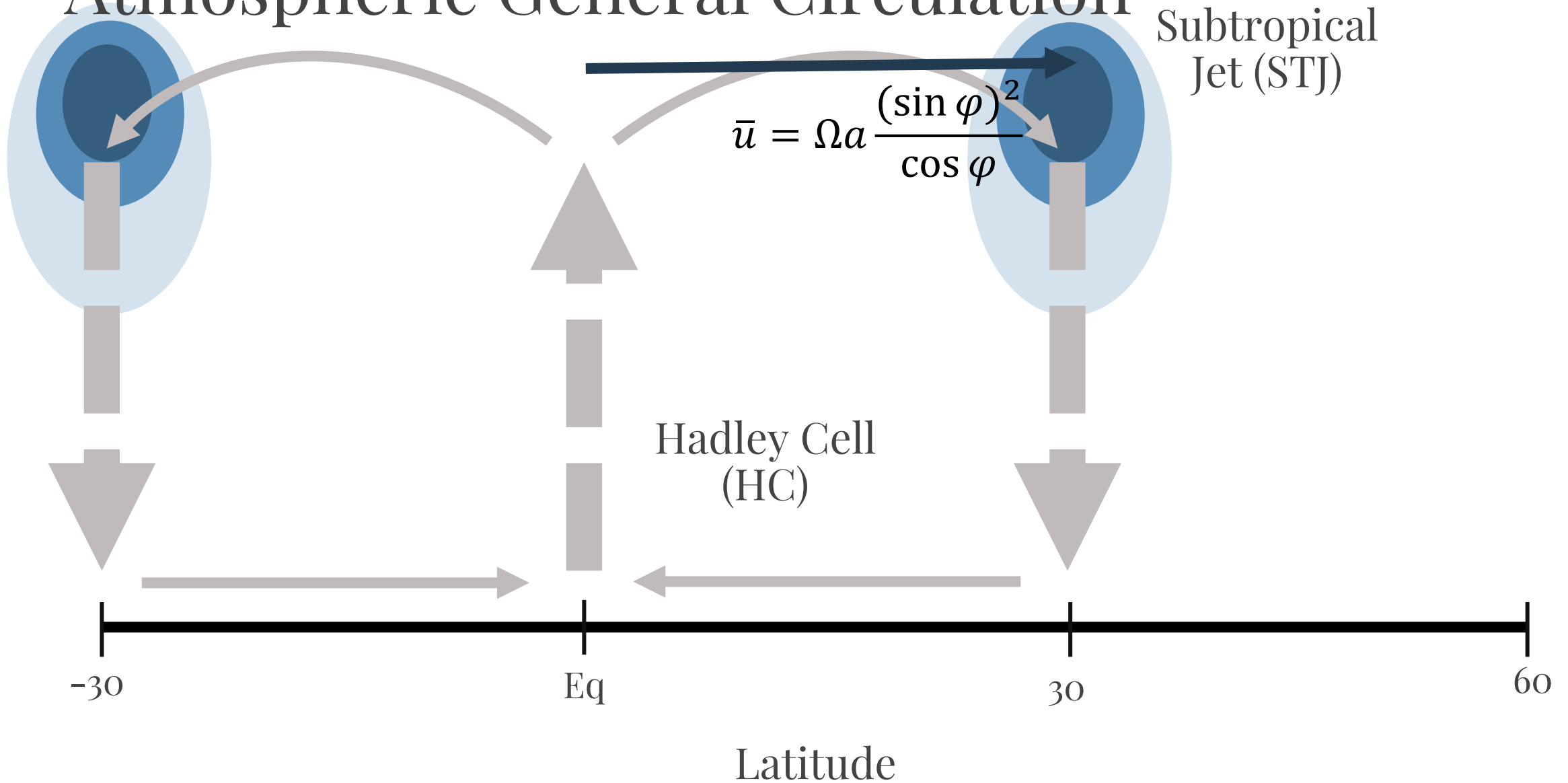
Atmospheric General Circulation



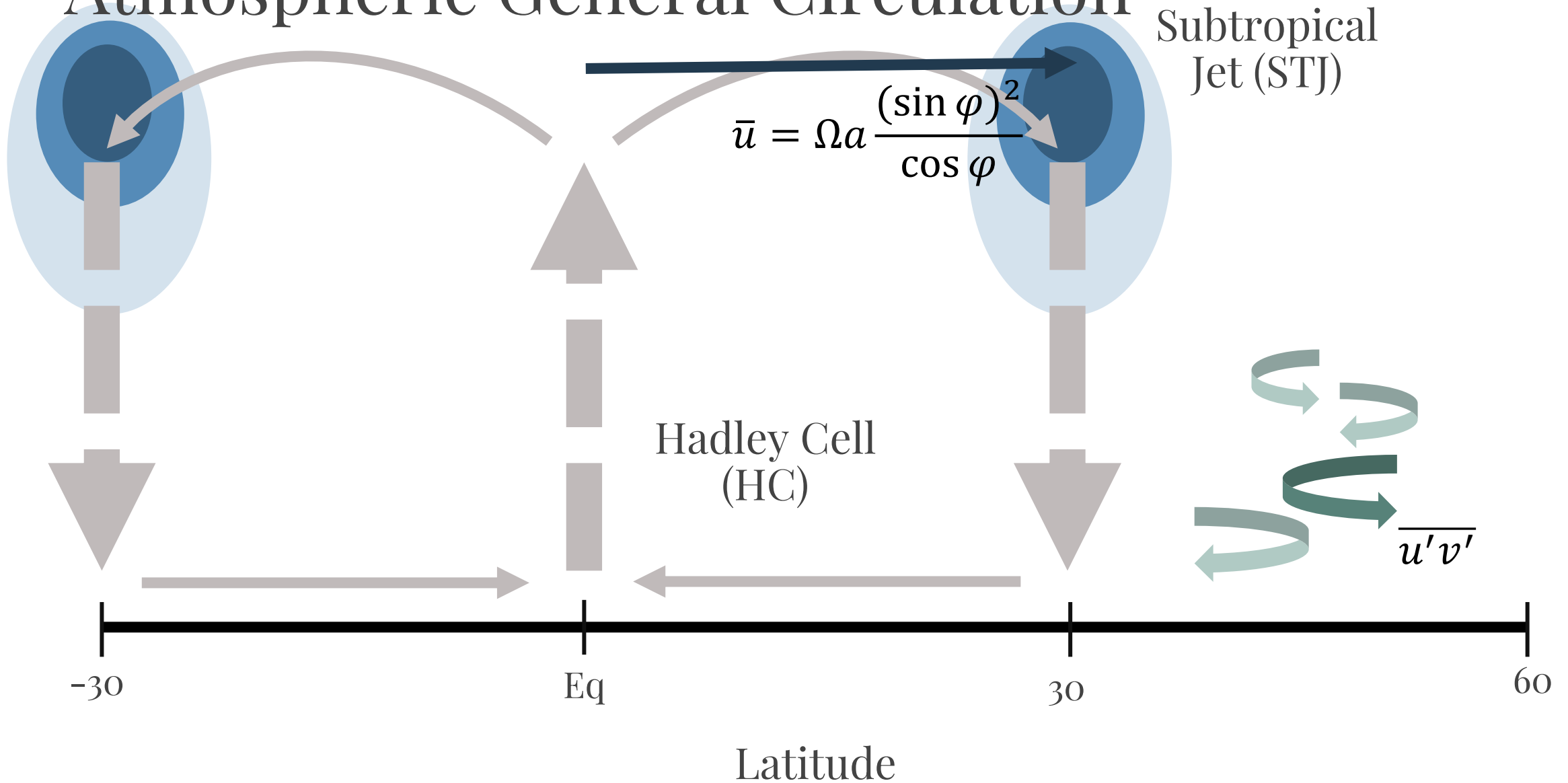
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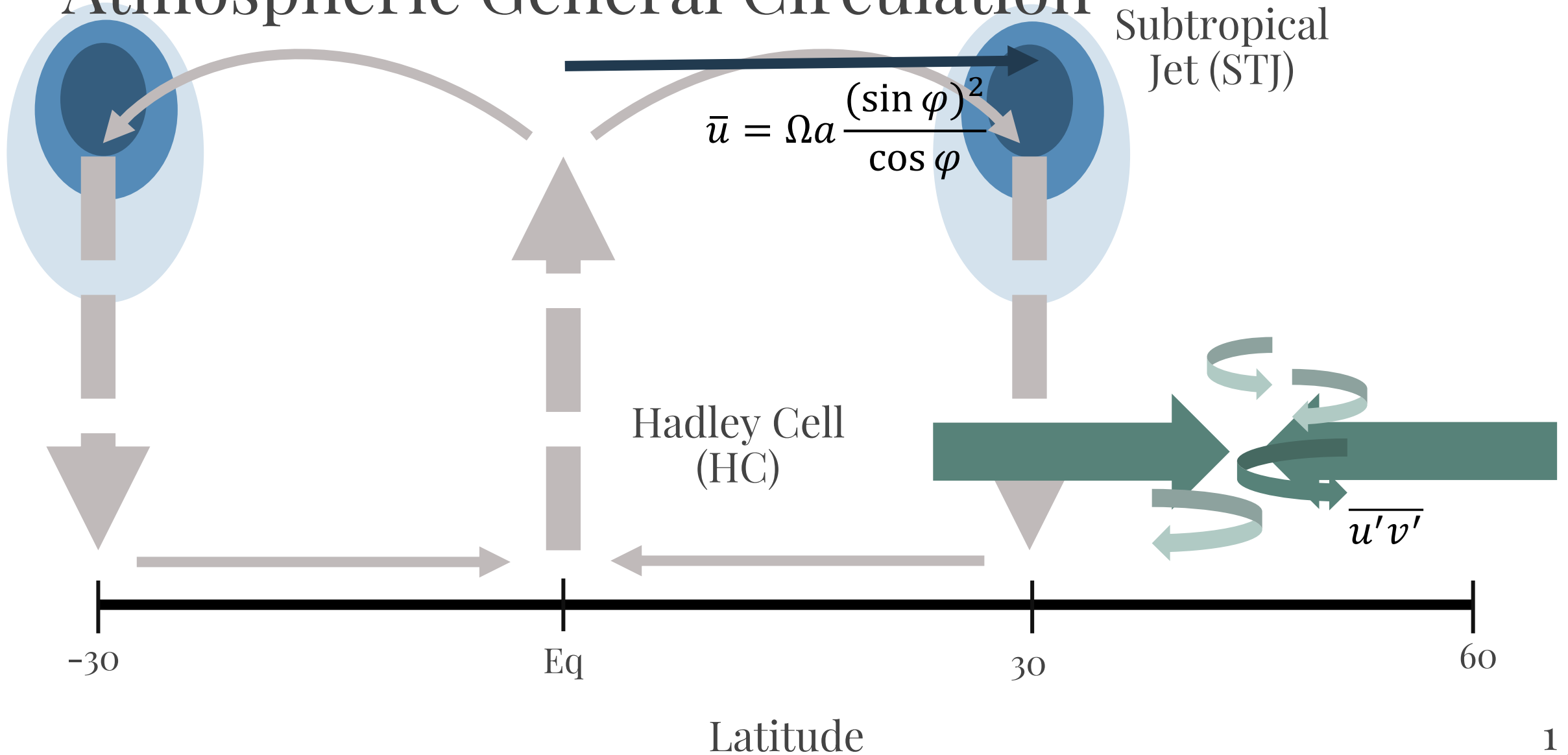
Atmospheric General Circulation



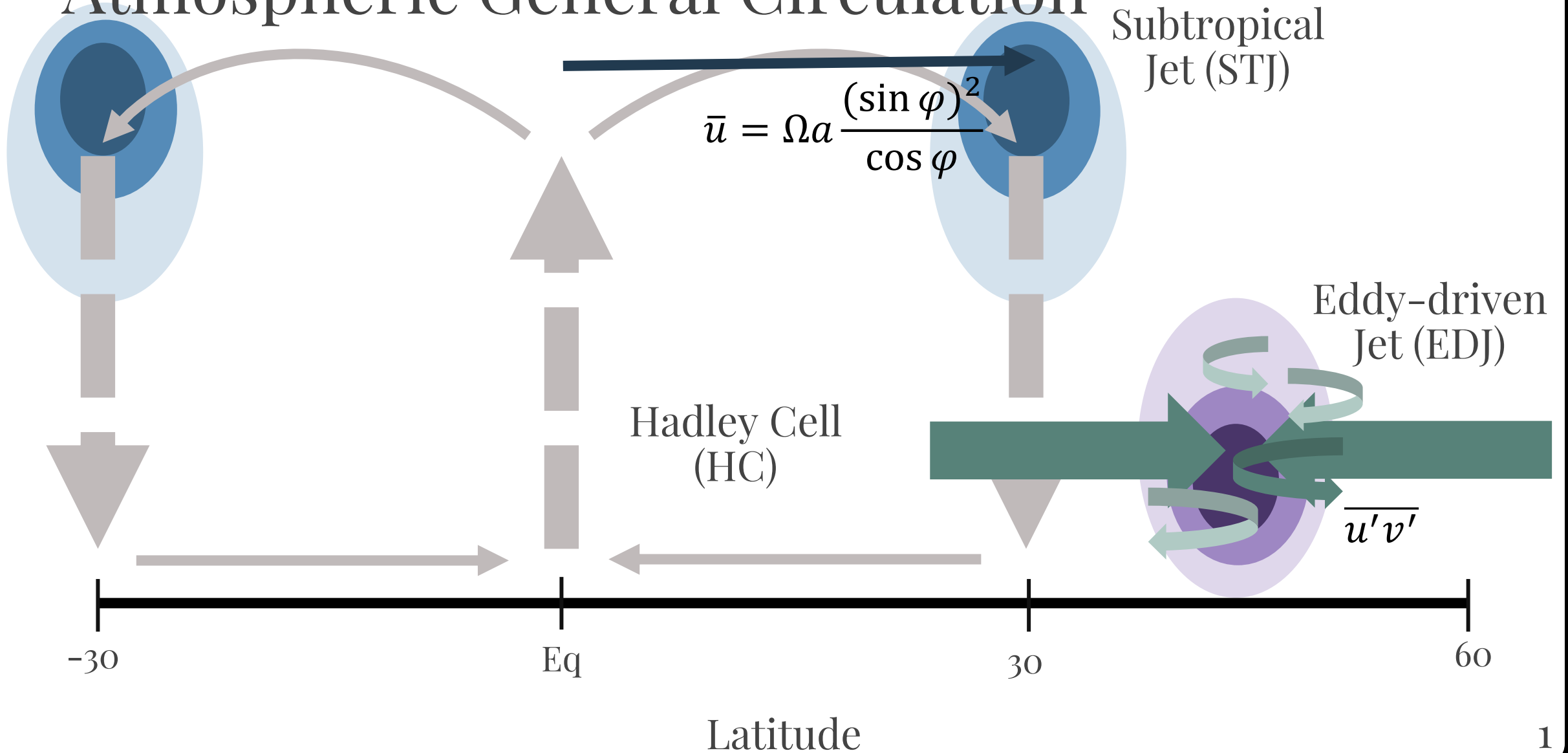
Atmospheric General Circulation



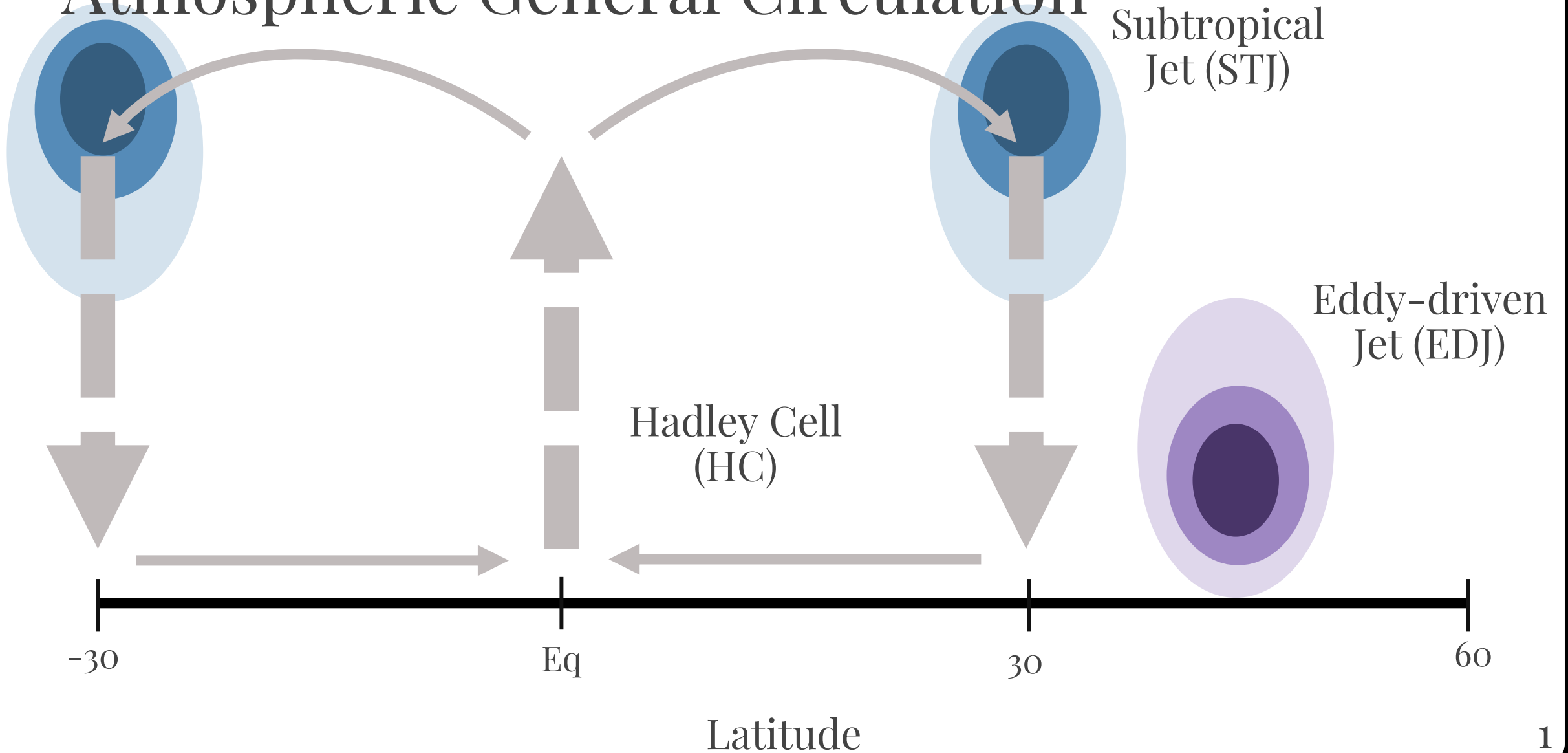
Atmospheric General Circulation



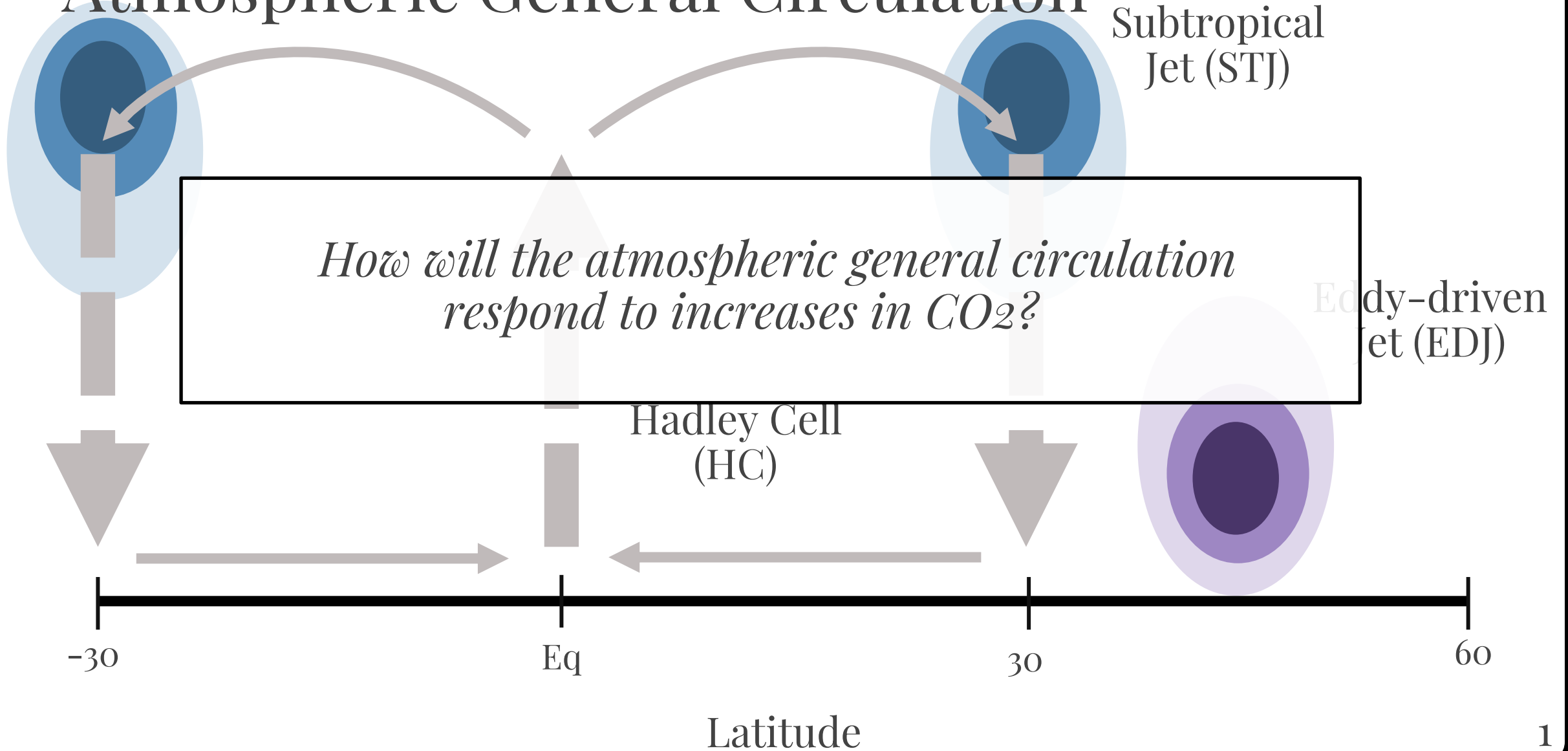
Atmospheric General Circulation



Atmospheric General Circulation

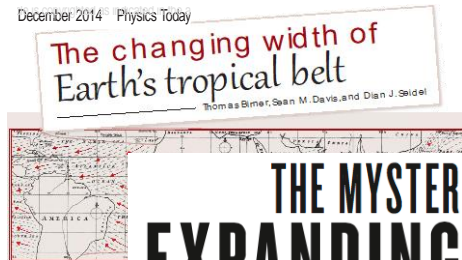


Atmospheric General Circulation

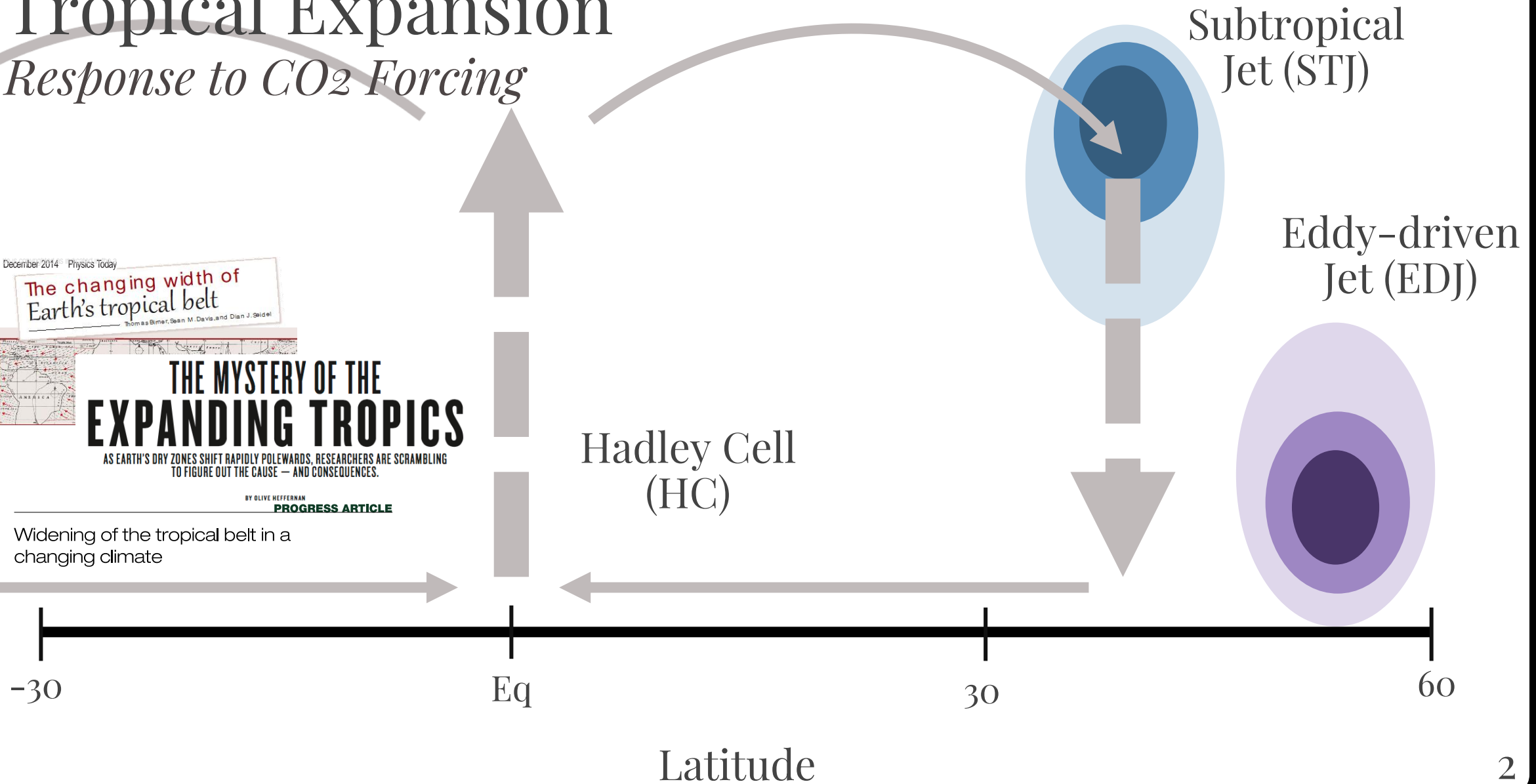


Tropical Expansion

Response to CO₂ Forcing

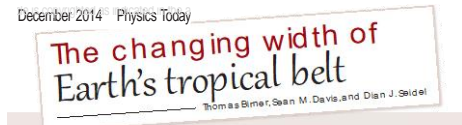


Widening of the tropical belt in a changing climate



Tropical Expansion

Response to CO₂ Forcing



THE MYSTERY OF THE EXPANDING TROPICS

AS EARTH'S DRY ZONES SHIFT RAPIDLY POLEWARDS, RESEARCHERS ARE SCRAMBLING TO FIGURE OUT THE CAUSE — AND CONSEQUENCES.

BY OLIVE HEFFERNAN
PROGRESS ARTICLE

Widening of the tropical belt in a changing climate

how to measure?

Hadley Cell (HC)

Subtropical Jet (STJ)

Eddy-driven Jet (EDJ)

-30

Eq

30

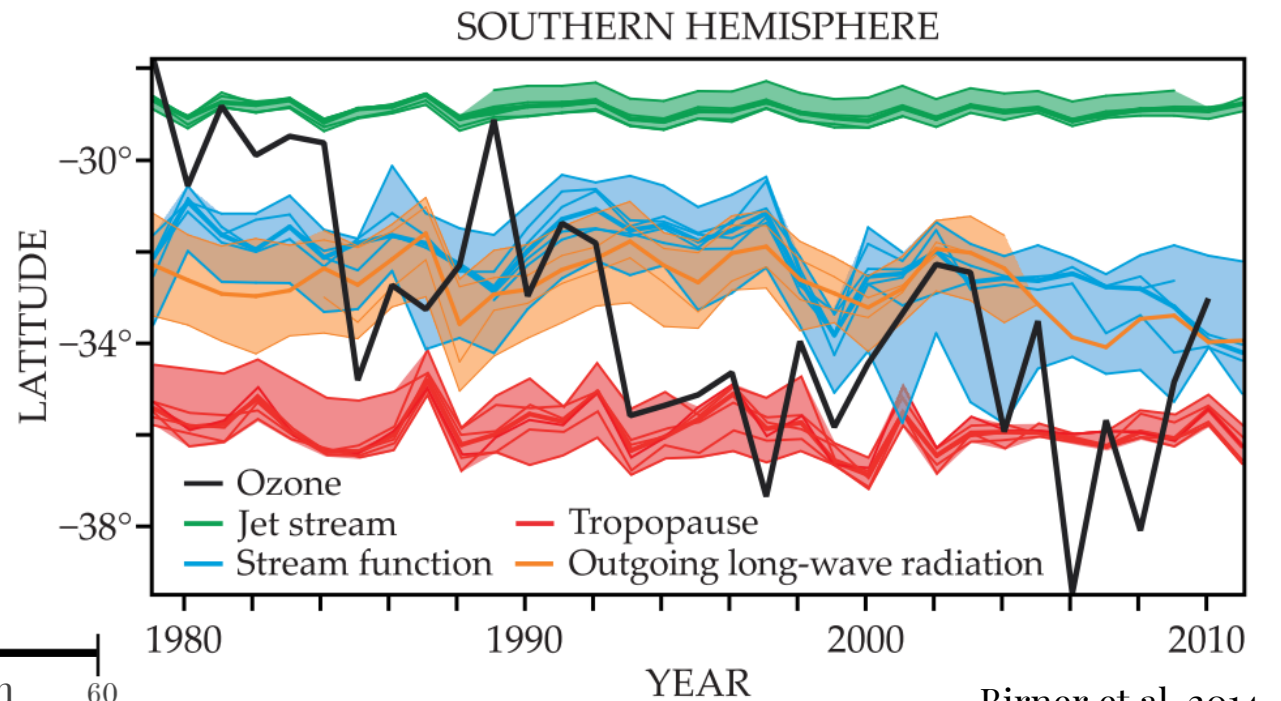
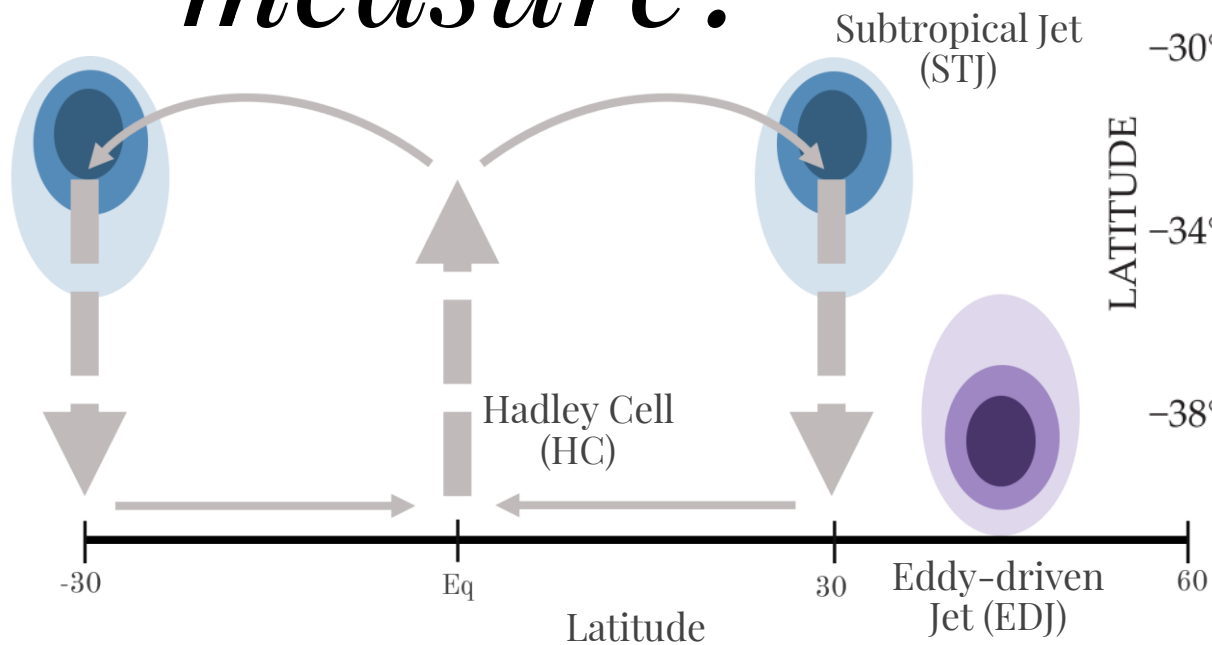
60

Latitude

Tropical Expansion

Response to CO₂ Forcing

*how to
measure?*



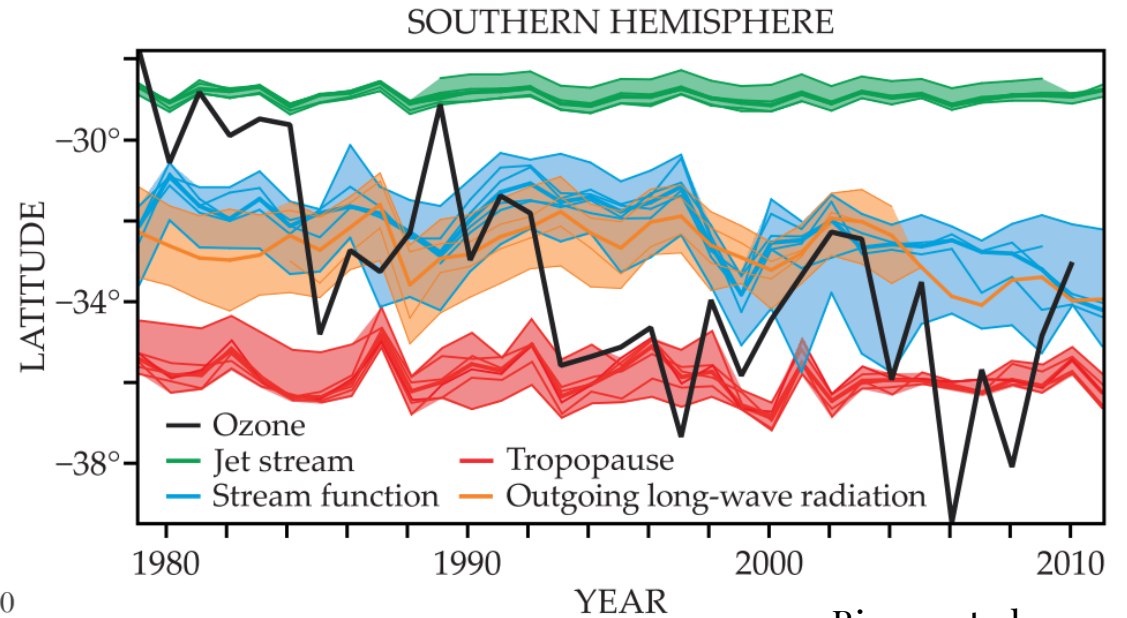
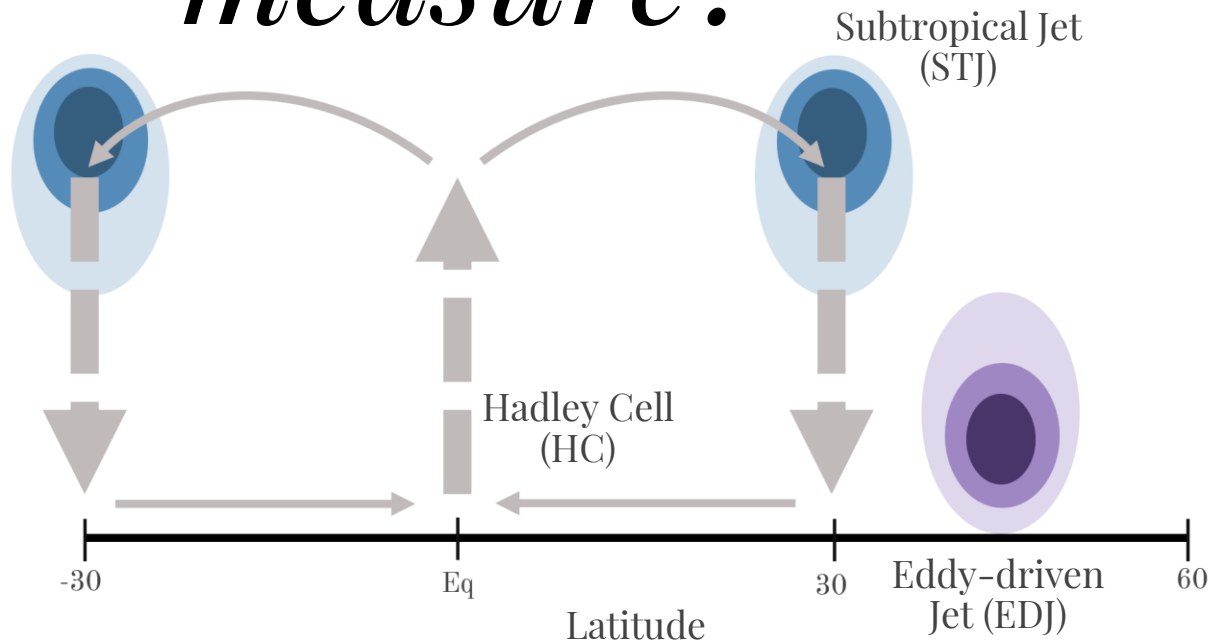
Birner et al. 2014

Tropical Expansion

Response to CO₂ Forcing

*how to
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Current understanding of atmospheric general circulation: the subtropical jet (STJ) should be coupled to the Hadley cells (HC)...



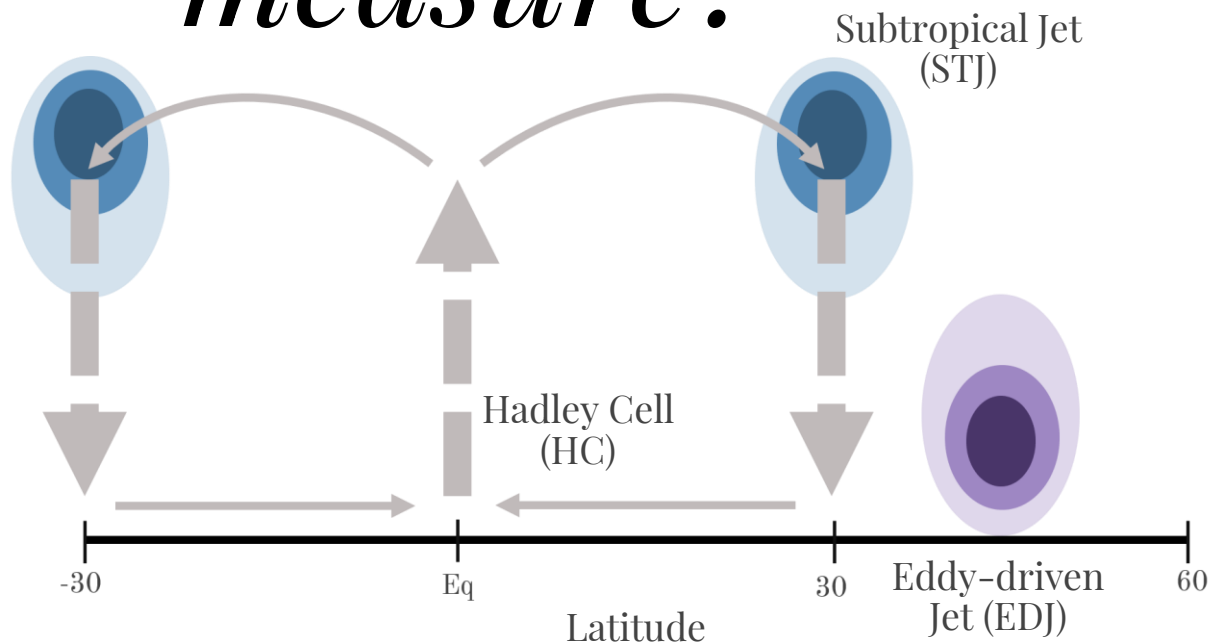
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Tropical Expansion

Response to CO₂ Forcing

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... reanalysis products and models do not support this!

Waugh et al. (2018)

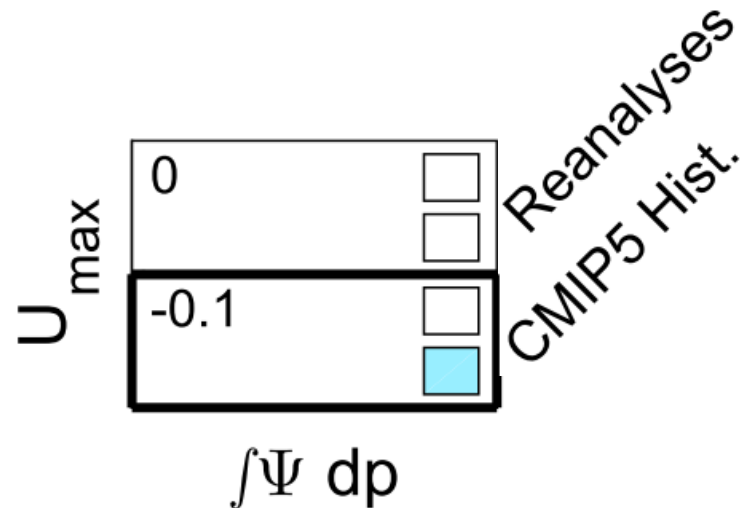
Davis & Birner (2017)

Solomon et al. (2017)

Tropical Expansion

Response to CO₂ Forcing

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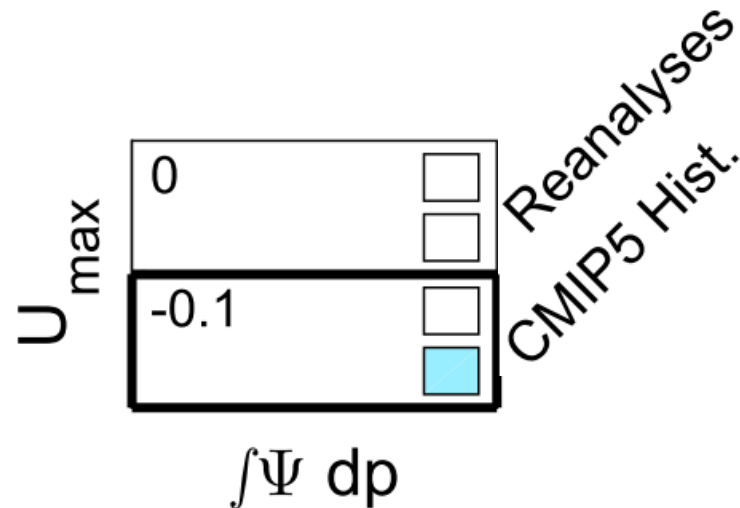
Tropical Expansion

Response to CO₂ Forcing

	(a) NH	(b) SH
DJF	0.02 (0.11)	0.05 (0.27)
MAM	0.33* (0.15)	0.10 (0.18)
JJA	0.28* (0.15)	0.13 (0.16)
SON	0.07 (0.11)	-0.03 (0.18)
Ann	0.28 (0.16)	0.06 (0.20)
	STJ	STJ

Waugh et al. 2018

Current understanding of atmospheric general circulation: the subtropical jet (STJ) should be coupled to the Hadley cells (HC)...



Davis & Birner 2017

... reanalysis products and models do not support this!

Waugh et al. (2018)

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Tropical Expansion

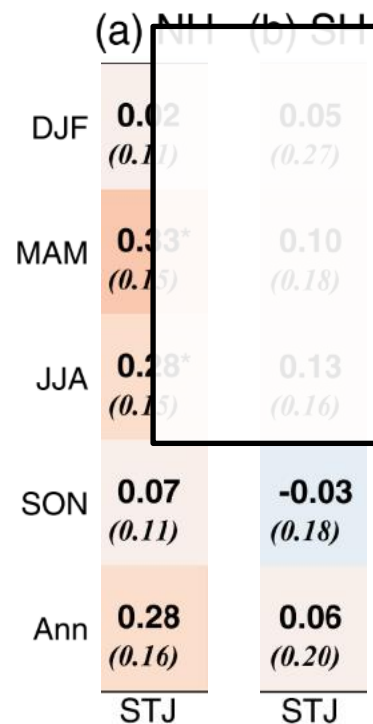
Response to CO₂ Forcing

Current understanding of atmospheric general circulation: the subtropical jet (STJ) should be coupled to the Hadley cells (HC)...

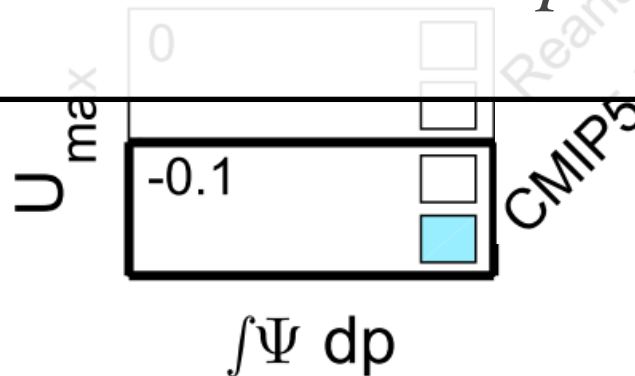
Knowledge Oversight!

Do we truly understand atmospheric circulation?

analysis products and models do not support this!



Waugh et al. 2018



Davis & Birner 2017

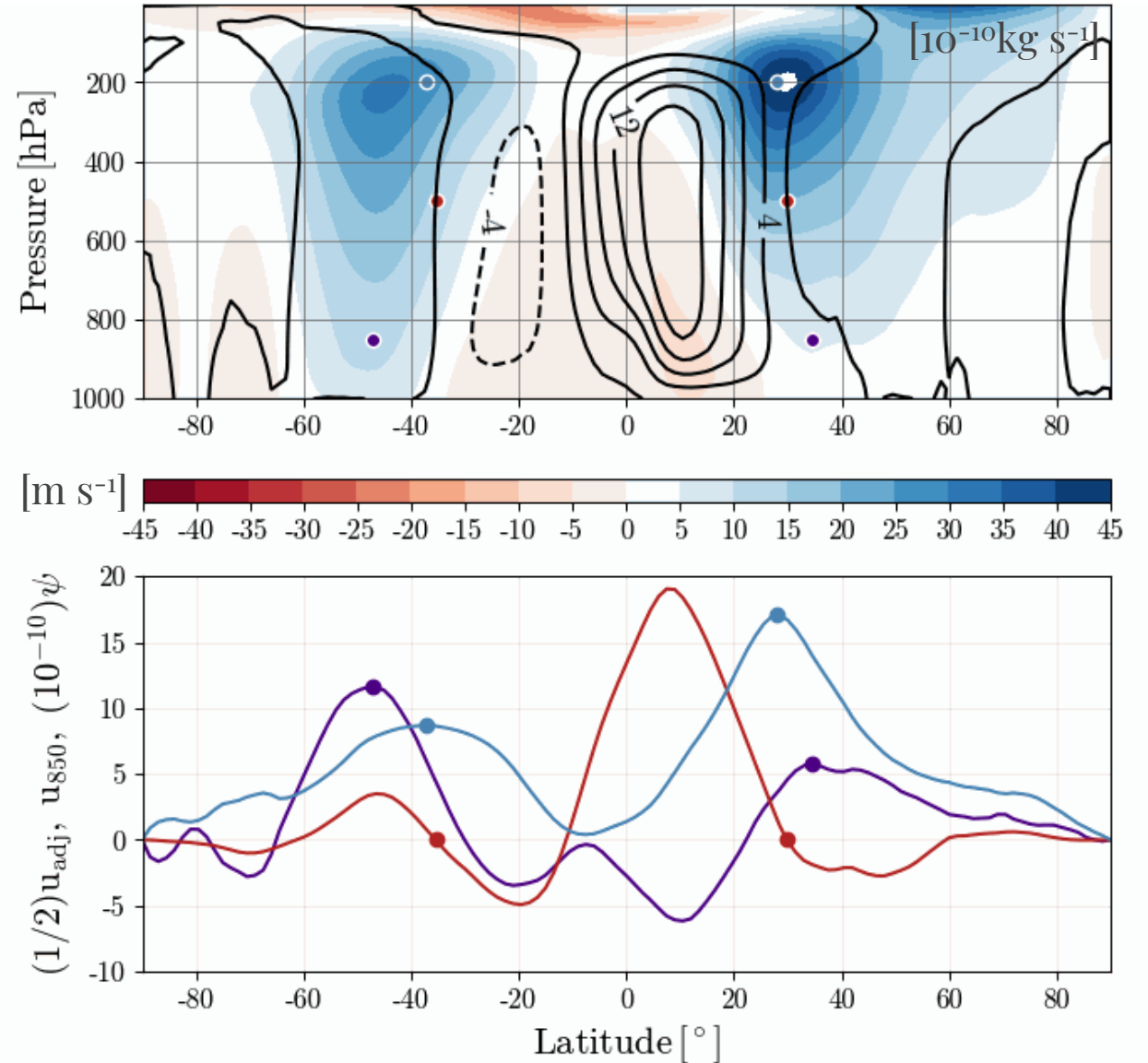
Waugh et al. (2018)
Davis & Birner (2017)
Solomon et al. (2017)

Circulation Metrics

Hadley Cell (HC)

● “PSI500”
 $\varphi_{HC} = \varphi(\psi_{500 \text{ hPa}} = 0)$

Zonal Wind, Streamfunction



Circulation Metrics

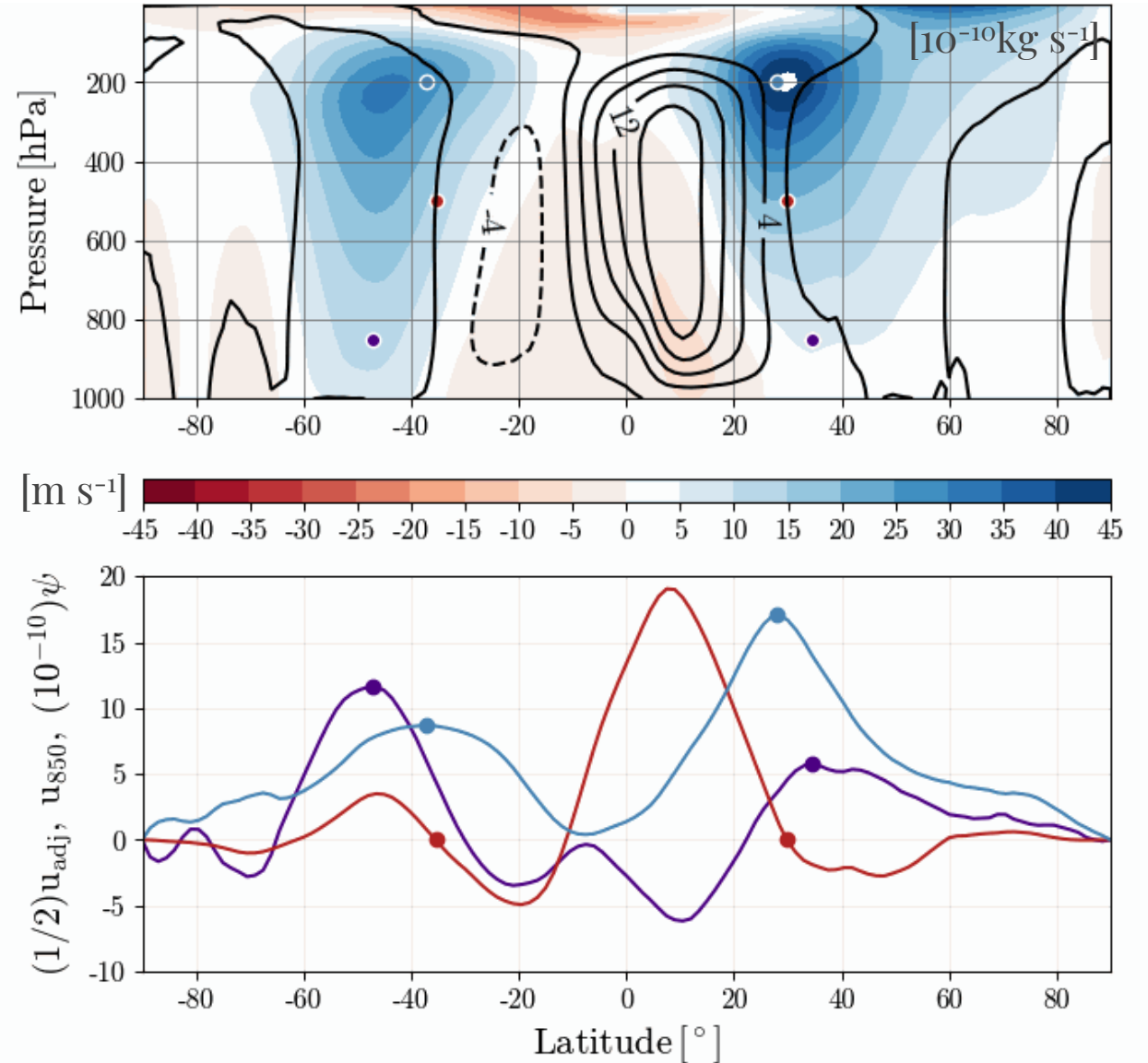
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Eddy-Driven Jet (EDJ)

● $\varphi_{EDJ} = \varphi(\max(u_{850 \text{ hPa}}))$

Zonal Wind, Streamfunction



Circulation Metrics

Hadley Cell (HC)

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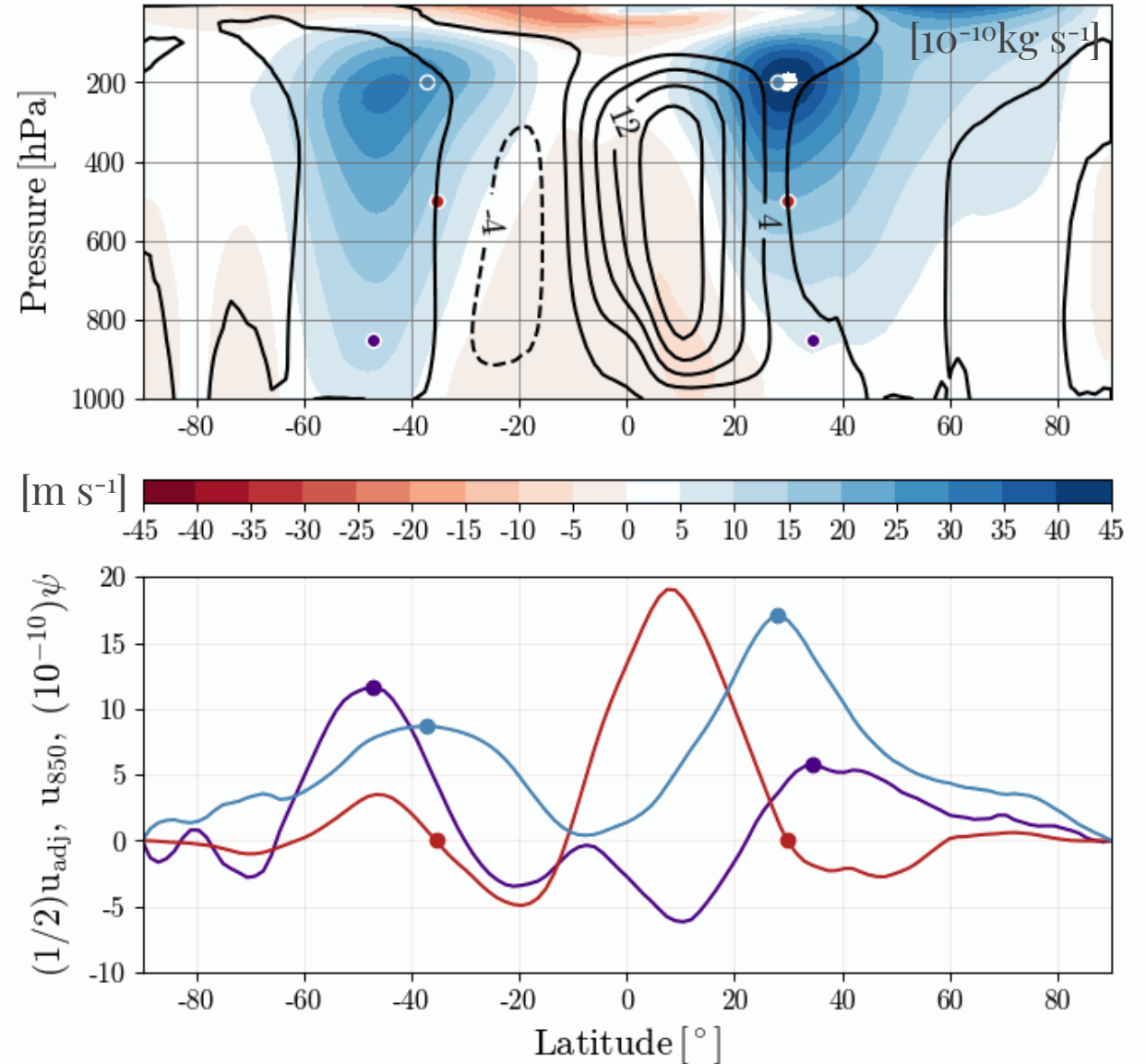
Eddy-Driven Jet (EDJ)

● $\varphi_{EDJ} = \varphi(\max(u_{850 \text{ hPa}}))$

Subtropical Jet (STJ)

● $\varphi_{STJ} = \varphi(\max(\Delta u))$
 $u_{STJ} = \Delta u(\varphi_{STJ})$
 $\Delta u = u_{100-400 \text{ hPa}} - u_{850 \text{ hPa}}$

Zonal Wind, Streamfunction



Circulation Metrics

Hadley Cell (HC)

● “PSI500”
 $\varphi_{HC} = \varphi(\psi_{500 \text{ hPa}} = 0)$

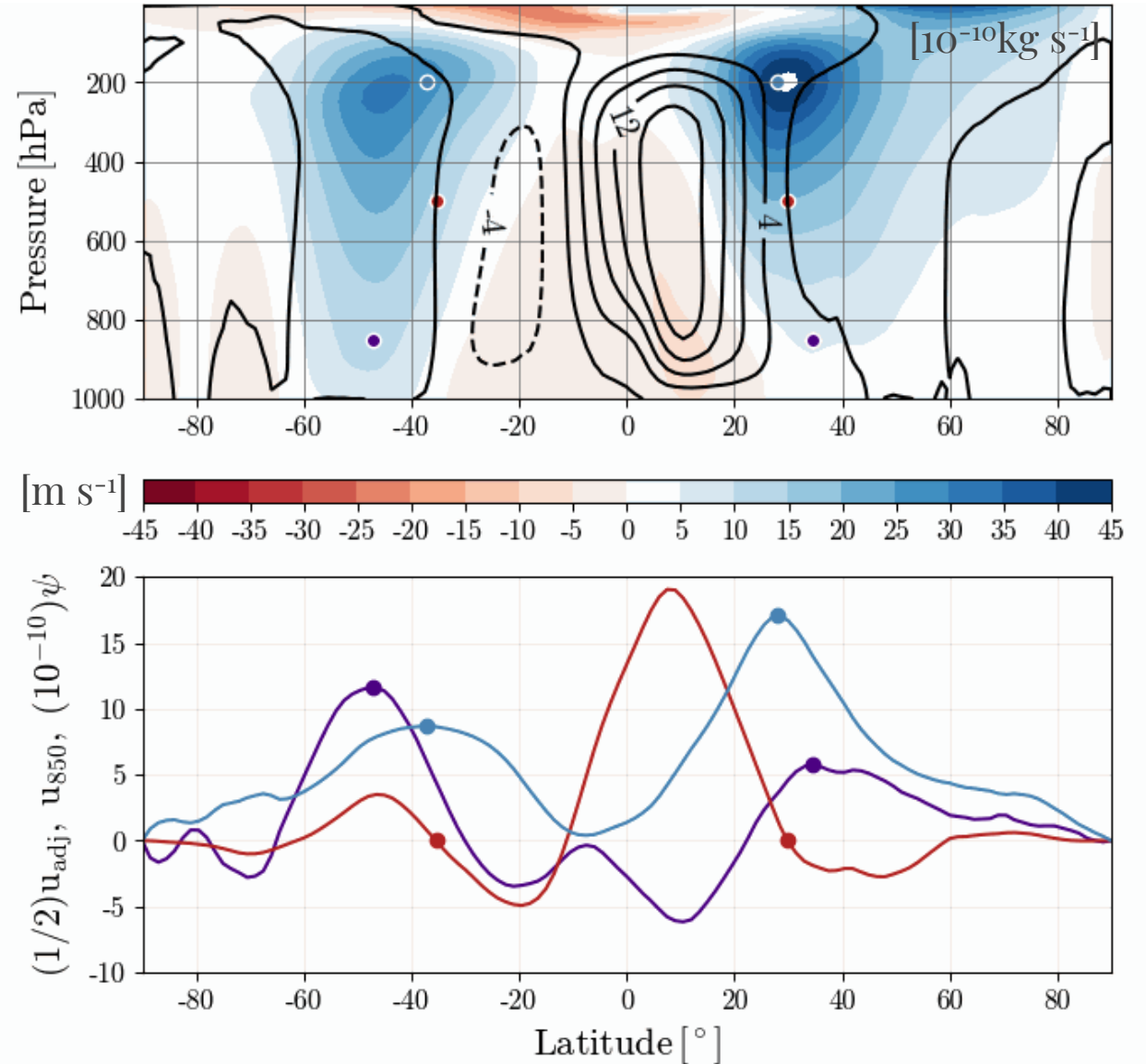
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Subtropical Jet (STJ)

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 $u_{STJ} = \Delta u(\varphi_{STJ})$
 $\Delta u = u_{100-400 \text{ hPa}} - u_{850 \text{ hPa}}$

Zonal Wind, Streamfunction



CMIP5

Coupled Model Intercomparison Project
(Phase 5)


Output from coupled simulations

ACCESS1-0	GISS-E2-R
bcc-csm1-1-m	HadGEM2-ES
bcc-csm1-1	Inmcm4
CanESM2	IPSL-CM5A-LR
CCSM4	IPSL-CM5B-LR
CNRM-CM5	MIROC5
CSIRO-Mk3-6-0	MIROC-ESM
FGOALS-s2	MPI-ESM-LR
GFDL-CM3	MPI-ESM-P
GFDL-ESM2G	MRI-CGCM3
GFDL-ESM2M	NorESM1-M
GISS-E2-H	

CMIP5

Coupled Model Intercomparison Project
(Phase 5)

Output from coupled simulations

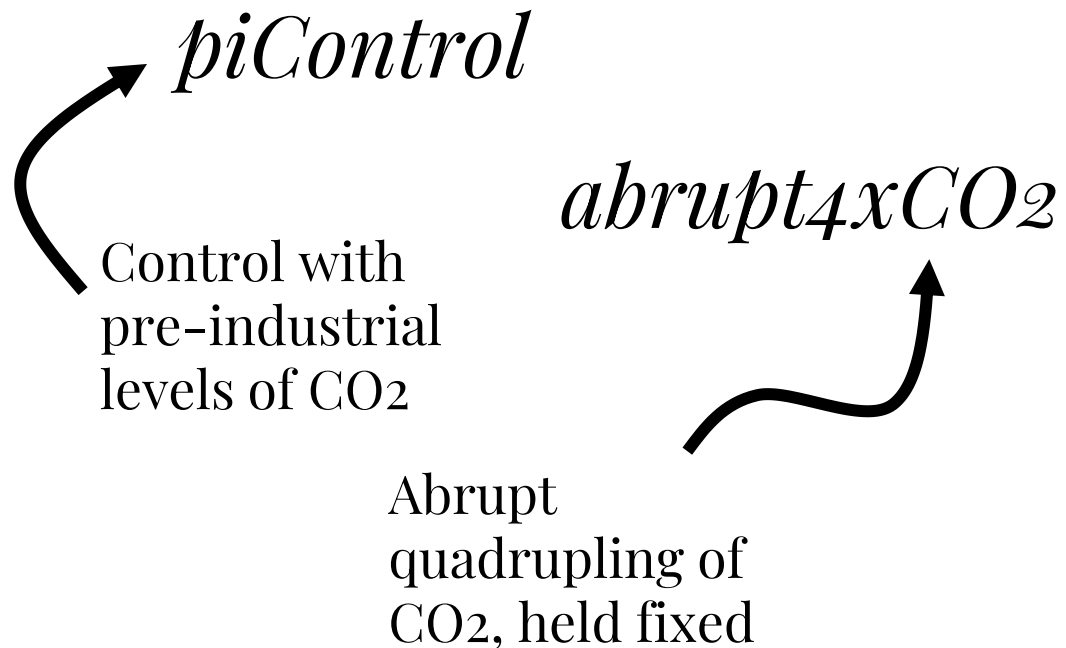
 *piControl*
Control with
pre-industrial
levels of CO₂

ACCESS1-0	GISS-E2-R
bcc-csm1-1-m	HadGEM2-ES
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CNRM-CM5	MIROC5
CSIRO-Mk3-6-0	MIROC-ESM
FGOALS-s2	MPI-ESM-LR
GFDL-CM3	MPI-ESM-P
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GISS-E2-H	

CMIP5

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GFDL-ESM2G	MRI-CGCM3
GFDL-ESM2M	NorESM1-M
GISS-E2-H	

CMIP5: Interannual

Southern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.07	-0.1	0.1	0.12	-0.03
ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)

Northern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.15	0.02	0.29*	0.2	-0.08
ϕ_{STJ}	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)

Menzel et al. 2019

CMIP5: Interannual

Southern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.07	-0.1	0.1	0.12	-0.03
ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)
ϕ_{HC}	0.52*	0.72*	0.46*	0.24	0.4*
ϕ_{EDJ}	(0.14)	(0.06)	(0.14)	(0.17)	(0.21)

Northern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.15	0.02	0.29*	0.2	-0.08
ϕ_{STJ}	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)
ϕ_{HC}	0.52*	0.47*	0.48*	0.4*	0.45*
ϕ_{EDJ}	(0.1)	(0.11)	(0.11)	(0.11)	(0.08)

Menzel et al. 2019

CMIP5: Interannual

Southern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.07	-0.1	0.1	0.12	-0.03
ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)
ϕ_{HC}	-0.19	-0.34	-0.14	-0.25*	-0.1
maxSTJ	(0.16)	(0.26)	(0.16)	(0.13)	(0.17)
ϕ_{HC}	0.52*	0.72*	0.46*	0.24	0.4*
ϕ_{EDJ}	(0.14)	(0.06)	(0.14)	(0.17)	(0.21)

Northern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.15	0.02	0.29*	0.2	-0.08
ϕ_{STJ}	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)
ϕ_{HC}	-0.39*	-0.3*	-0.52*	-0.29*	-0.15
maxSTJ	(0.14)	(0.13)	(0.13)	(0.18)	(0.15)
ϕ_{HC}	0.52*	0.47*	0.48*	0.4*	0.45*
ϕ_{EDJ}	(0.1)	(0.11)	(0.11)	(0.11)	(0.08)

Menzel et al. 2019

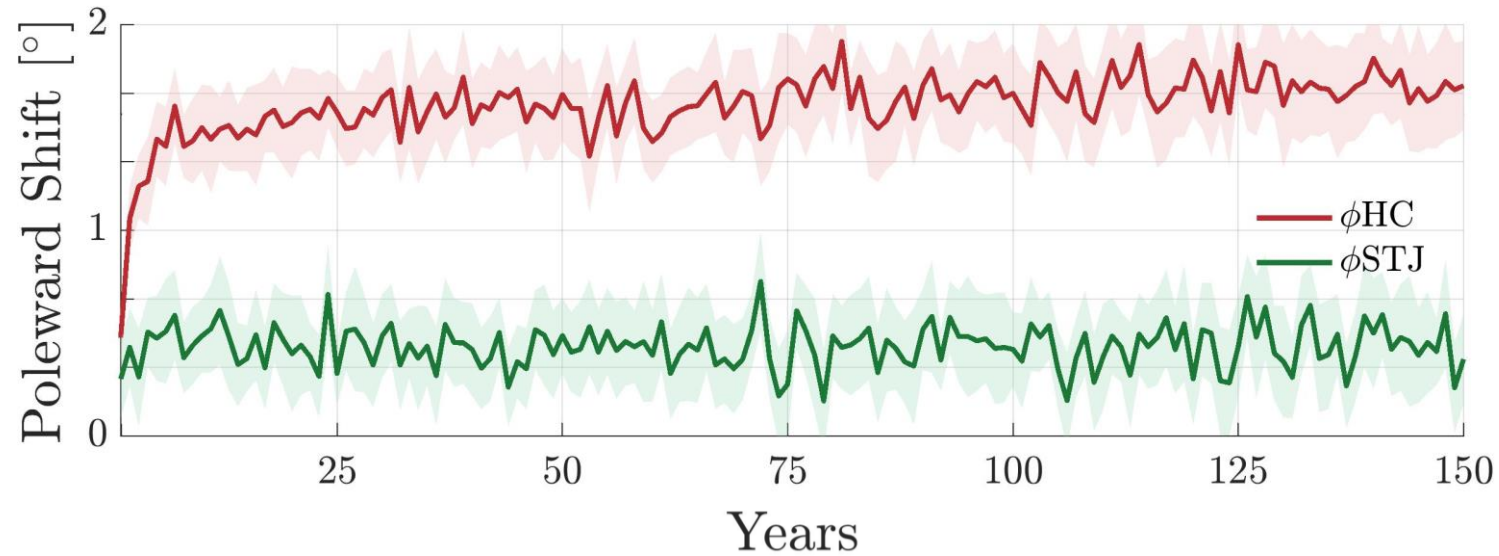
CMIP5: Interannual

Southern Hemisphere				Northern Hemisphere					
	ANN	DJF	MAM	Natural Variability		DJF	MAM	JJA	SON
ϕ_{HC}	0.07 (0.22)	-0.1 (0.2)	0.1 (0.22)			ϕ_{STJ}	0.02 (0.12)	0.29* (0.16)	0.2 (0.17)
maxSTJ	HC Location STJ Strength			—		-0.3* (0.13)	-0.52* (0.13)	-0.29* (0.18)	-0.15 (0.15)
ϕ_{EDJ}	0.52* (0.14)	0.72* (0.06)	0.46* (0.14)	More poleward Hadley cell, weaker subtropical jet		0.47* (0.11)	0.48* (0.11)	0.4* (0.11)	0.45* (0.08)

Menzel et al. 2019

Menzel et al. 2019

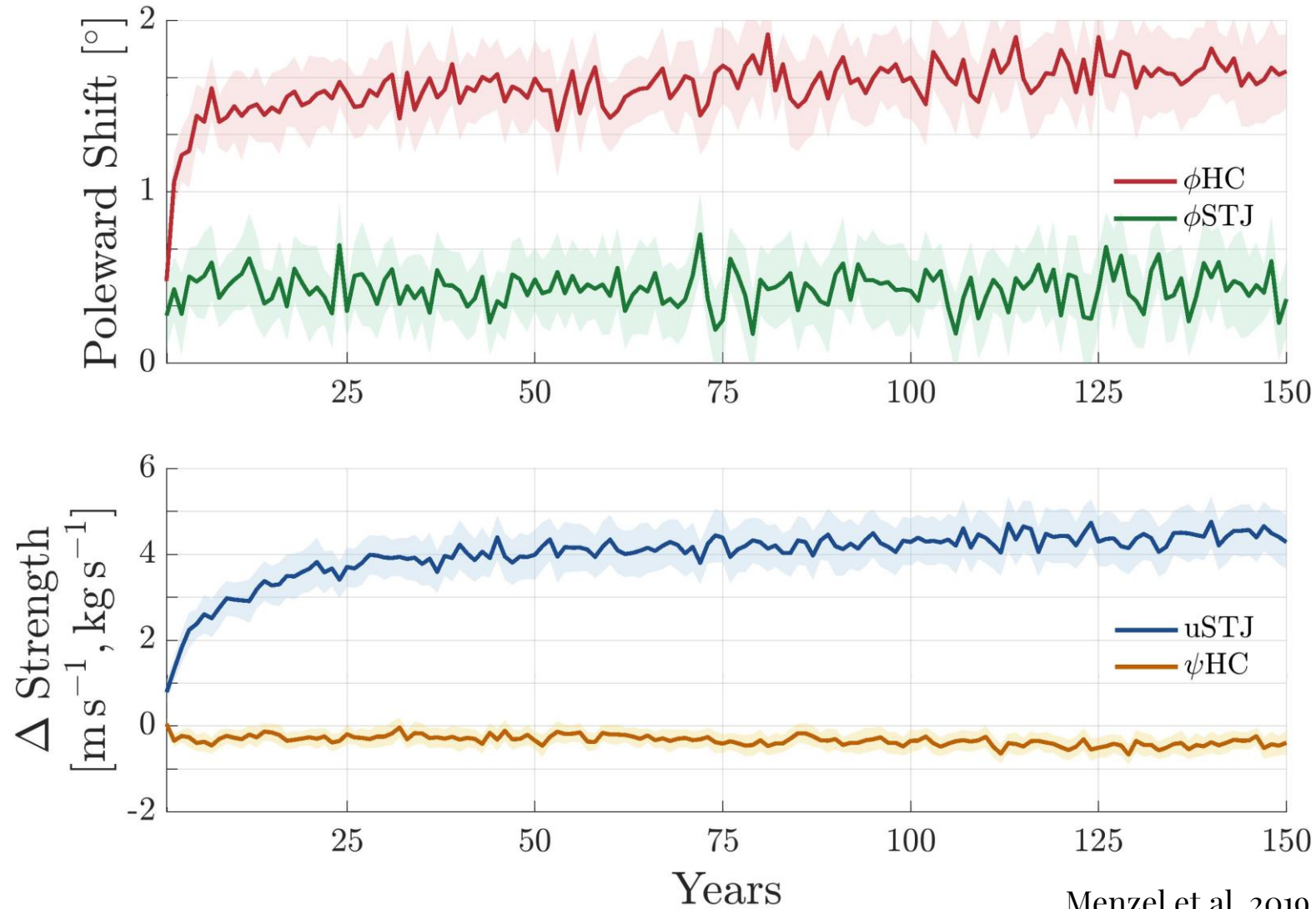
CMIP5: Transient Response



Time series of
metrics' response to
 $4\times\text{CO}_2$

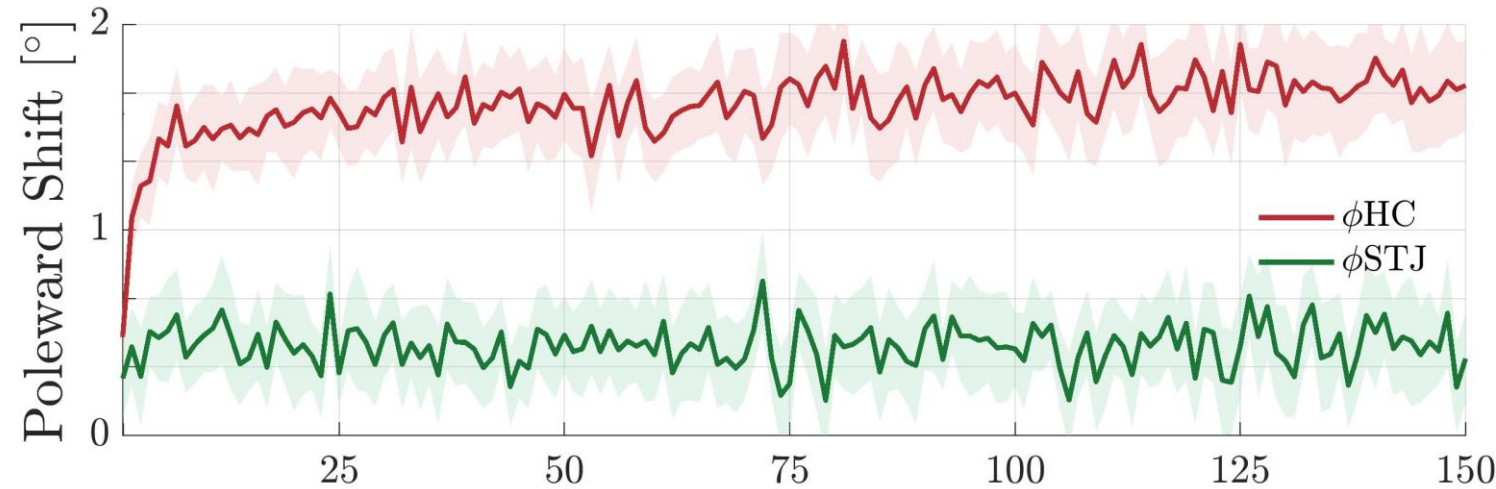
Menzel et al. 2019

CMIP5: Transient Response

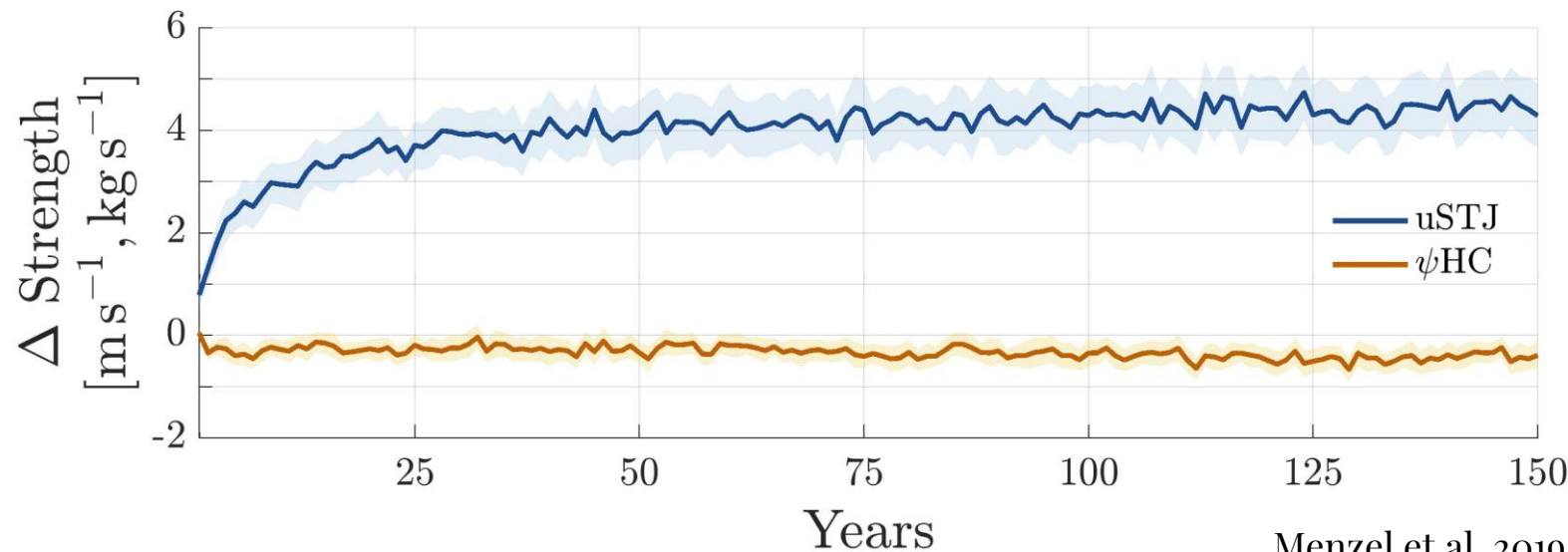


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CMIP5: Transient Response

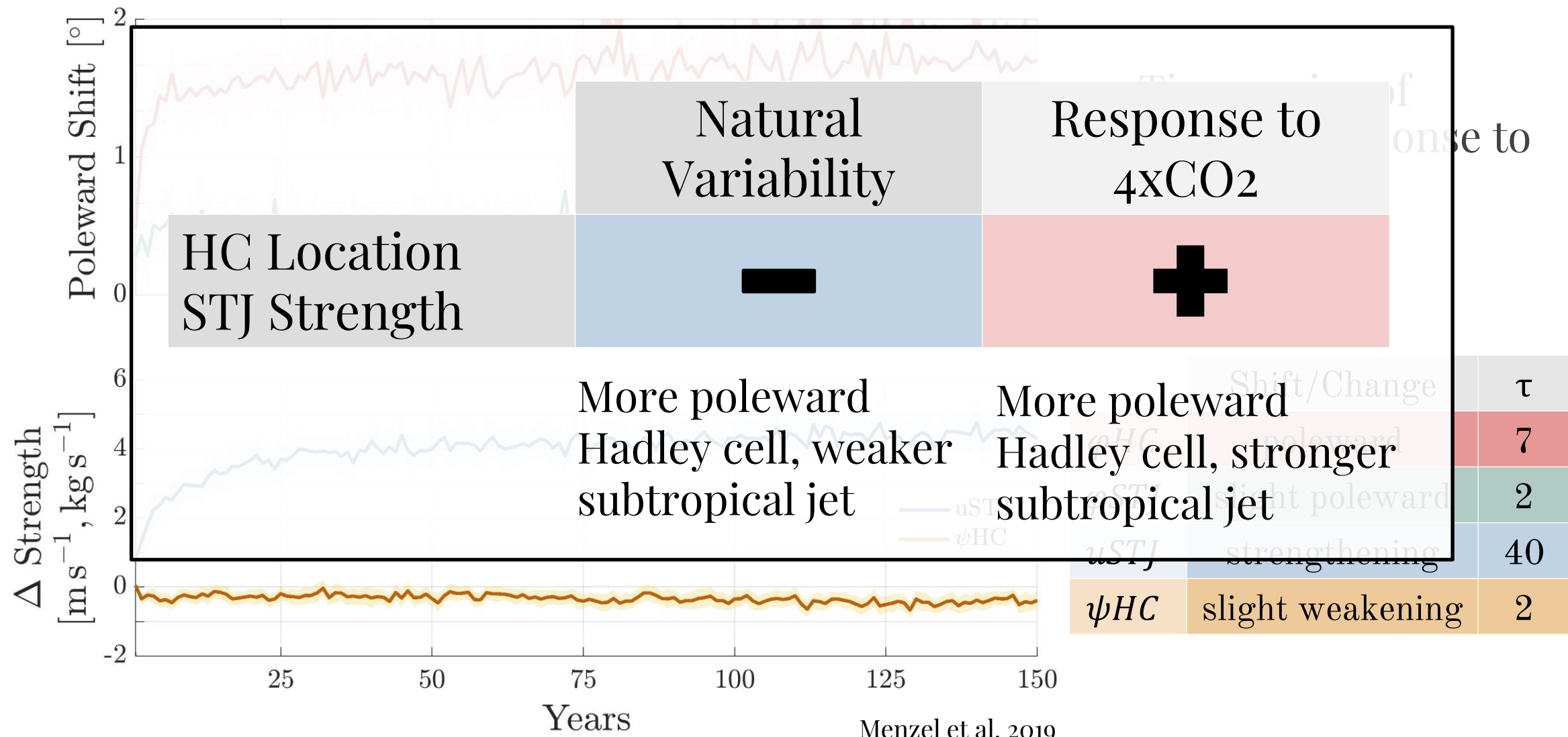


Time series of metrics' response to $4\times\text{CO}_2$



	Shift/Change	τ
ϕ_{HC}	poleward	7
ϕ_{STJ}	slight poleward	2
u_{STJ}	strengthening	40
ψ_{HC}	slight weakening	2

CMIP5: Transient Response



Metric Analysis

Conclusion:

The subtropical jet (STJ) is not coupled to the Hadley cell (HC), there must be physical processes responsible for their distinctive behavior.

Southern Hemisphere						Northern Hemisphere				
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ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)

Menzel et al. 2019

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Menzel et al. 2019

Lingering Question:

What are the physical processes responsible for the distinct behavior of the HC and STJ?

Physical Balances



Meridional Flow Balance

The diagram illustrates the meridional flow balance equation, $(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y}(\overline{u'v'})$. The terms are color-coded and labeled with arrows:

- Coriolis parameter**: Points to the f term (green box).
- relative vorticity**: Points to the $\vec{\zeta}$ term (red box).
- mean meridional flow**: Points to the \bar{v} term (blue box).
- eddy momentum flux**: Points to the $\frac{\partial}{\partial y}(\overline{u'v'})$ term (purple box).

The equation is presented as:

$$(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y}(\overline{u'v'})$$

Meridional Flow Balance

$$(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y} (\overline{u'v'})$$

The diagram shows the meridional flow balance equation with color-coded terms and arrows pointing to their physical meanings:

- f (green box) is labeled "Coriolis parameter" (green arrow).
- $\vec{\zeta}$ (red box) is labeled "relative vorticity" (red arrow).
- \bar{v} (blue box) is labeled "mean meridional flow" (blue arrow).
- $\frac{\partial}{\partial y} (\overline{u'v'})$ (purple box) is labeled "eddy momentum flux" (purple arrow).

If eddies are negligible...

$$(f + \vec{\zeta})\bar{v} = 0$$

Meridional flow is
angular-momentum
conserving!

Meridional Flow Balance

$$(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y} (\overline{u'v'})$$

The diagram shows the equation $(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y} (\overline{u'v'})$ with four colored boxes highlighting specific terms: a green box around f , a red box around $\vec{\zeta}$, a blue box around \bar{v} , and a purple box around the entire right-hand side $\frac{\partial}{\partial y} (\overline{u'v'})$. Four curved arrows point from text labels to these boxes: a green arrow from "Coriolis parameter" to f , a red arrow from "relative vorticity" to $\vec{\zeta}$, a blue arrow from "mean meridional flow" to \bar{v} , and a purple arrow from "eddy momentum flux" to the right-hand side.

If eddies are negligible...

$$(f + \vec{\zeta})\bar{v} = 0$$

Meridional flow is angular-momentum conserving!

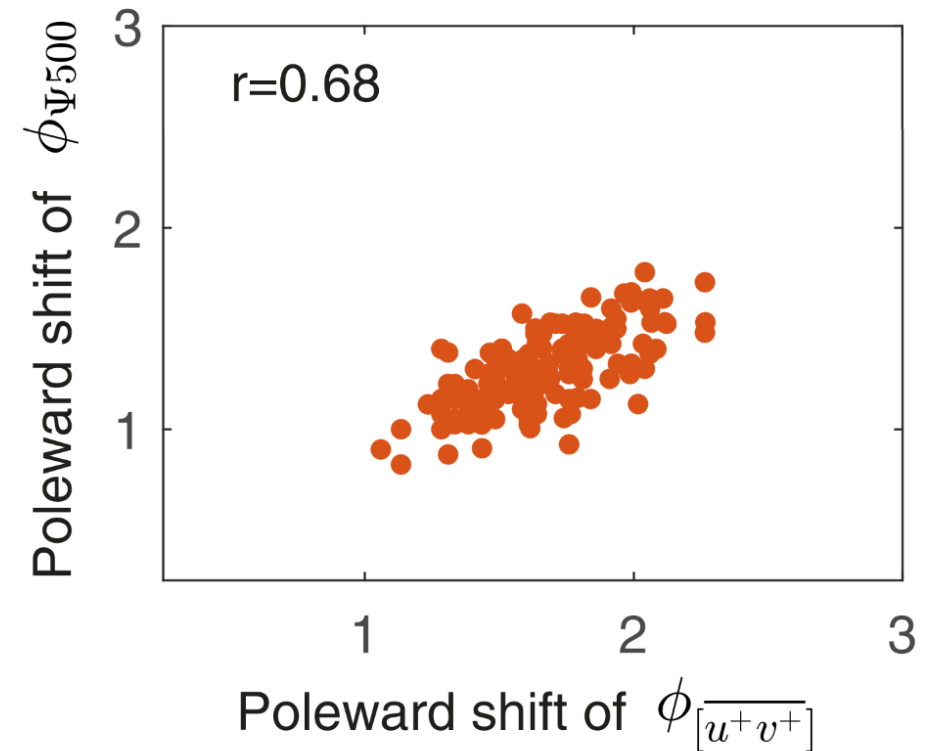
If not...

$$(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y} (\overline{u'v'})$$

Meridional flow is set by eddies!

Meridional Flow Balance

$$(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y} (\overline{u'v'})$$

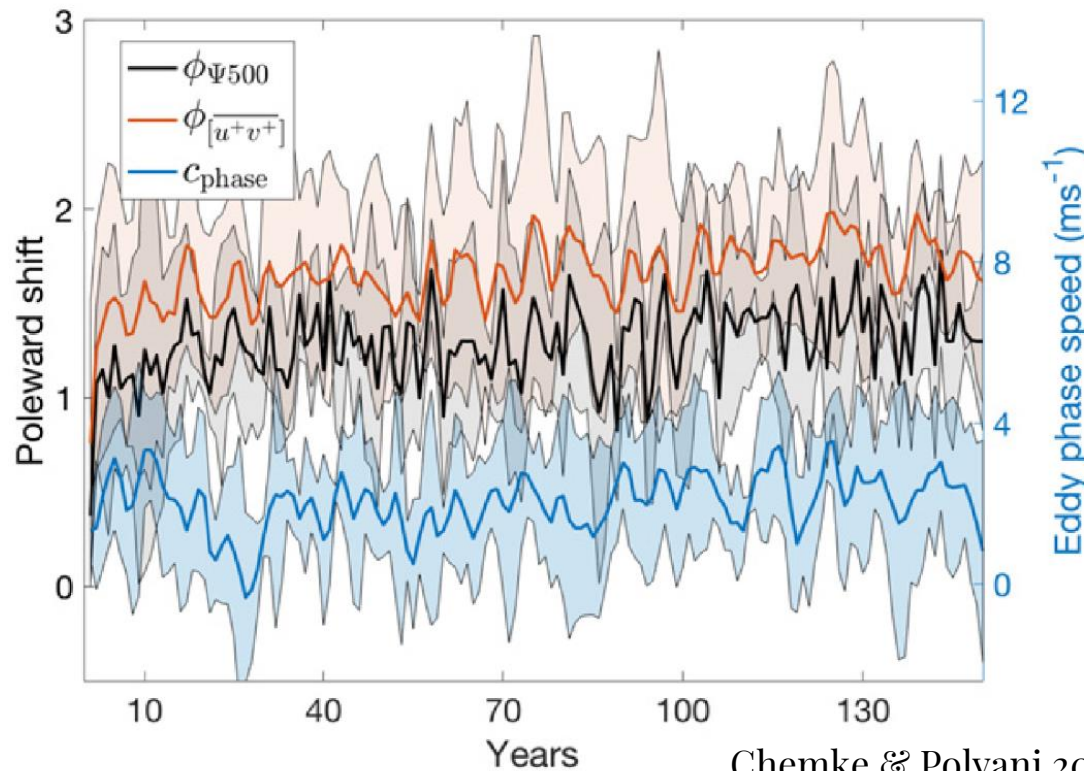


Chemke & Polvani 2019

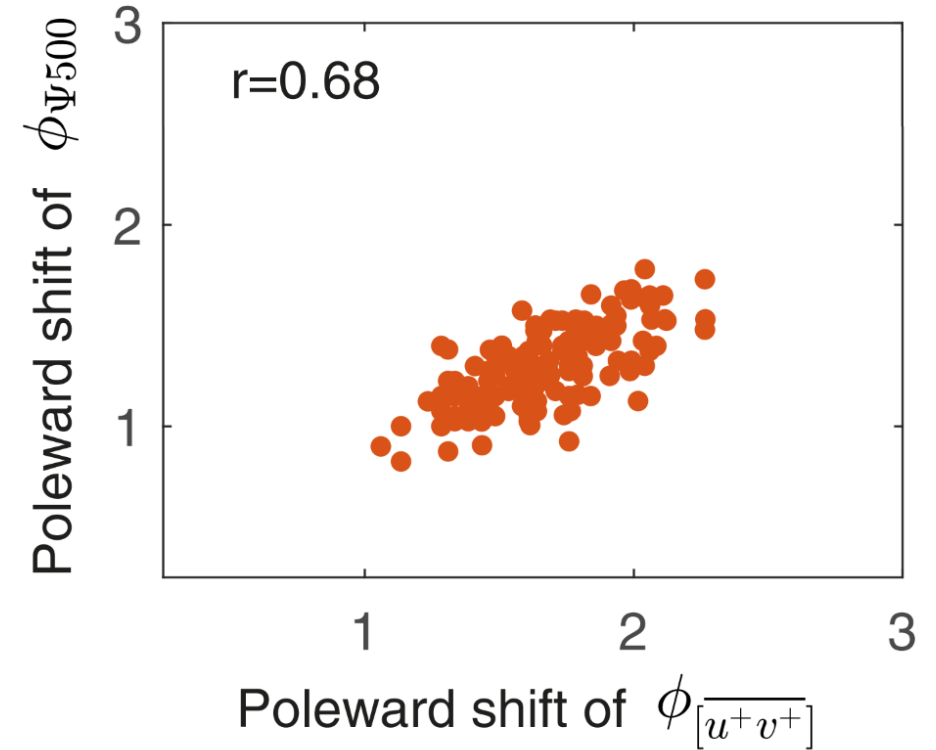
Meridional flow is
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Meridional Flow Balance

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Chemke & Polvani 2019



Chemke & Polvani 2019

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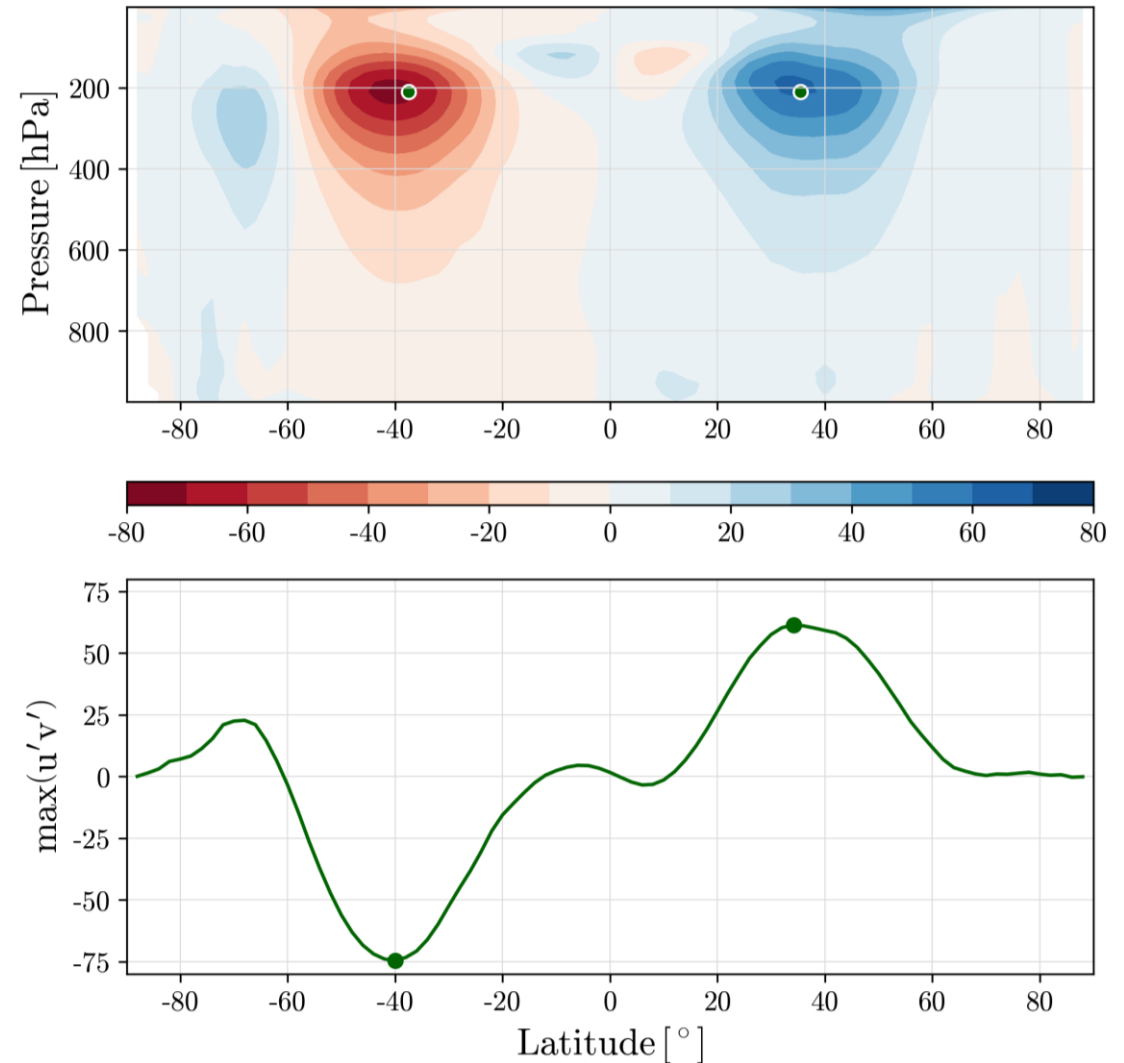
Meridional Flow Balance

$$(f + \vec{\zeta})\bar{v} = \frac{\partial}{\partial y}(\overline{u'v'})$$

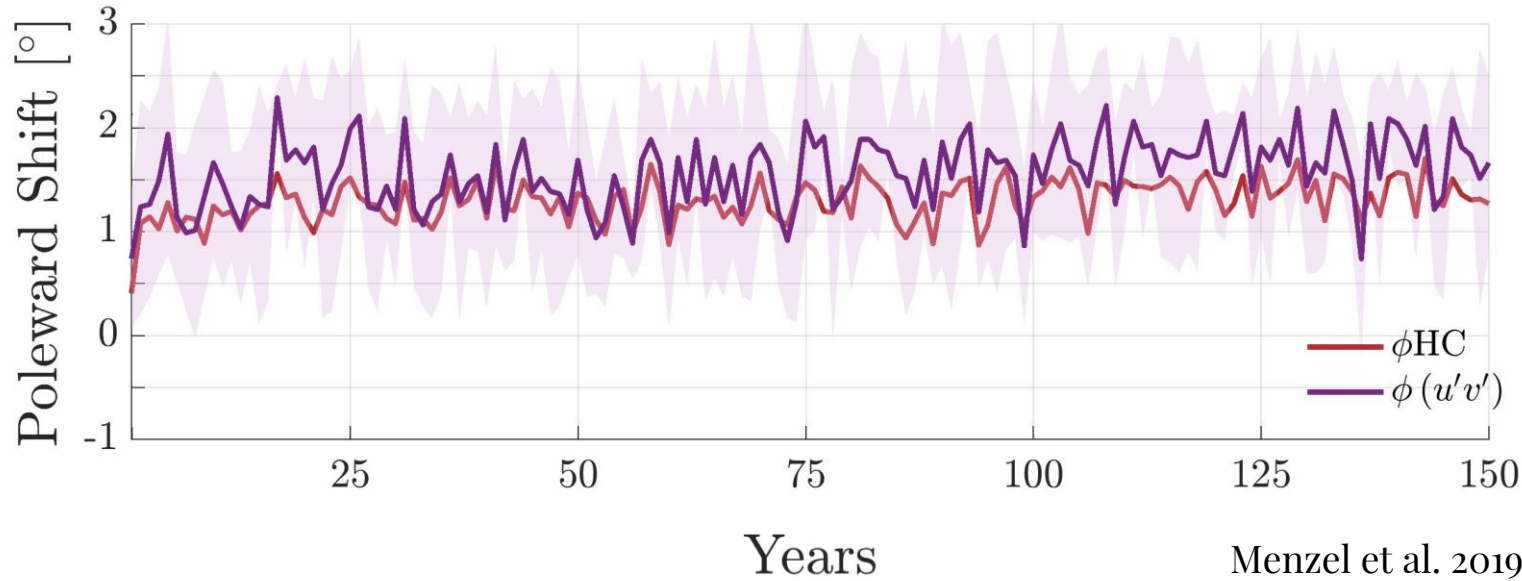
Maximum Eddy Momentum Flux

● $\varphi(\overline{u'v'}) = \varphi(\max(\overline{u'v'}))$

Eddy Momentum Flux



Meridional Flow Balance

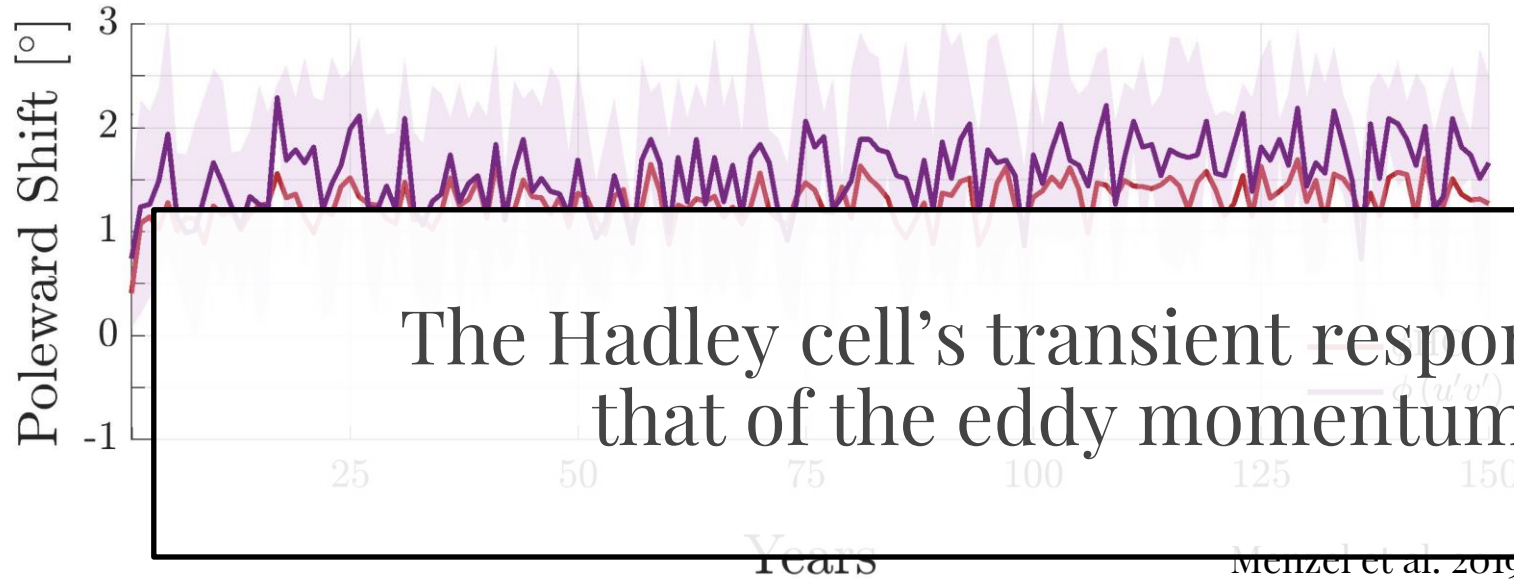


Menzel et al. 2019

	Shift/Change	τ
ϕ_{HC}	poleward	7
$\phi(u'v')$	poleward	5

HC edge:
– latitude of max
eddy momentum
flux $\phi(u'v')$

Meridional Flow Balance



	Shift/Change	τ
φ_{HC}	poleward	7
$\varphi(u'v')$	poleward	5

The Hadley cell's transient response follows that of the eddy momentum flux

HC edge:
– latitude of max
eddy momentum
flux $\varphi(u'v')$

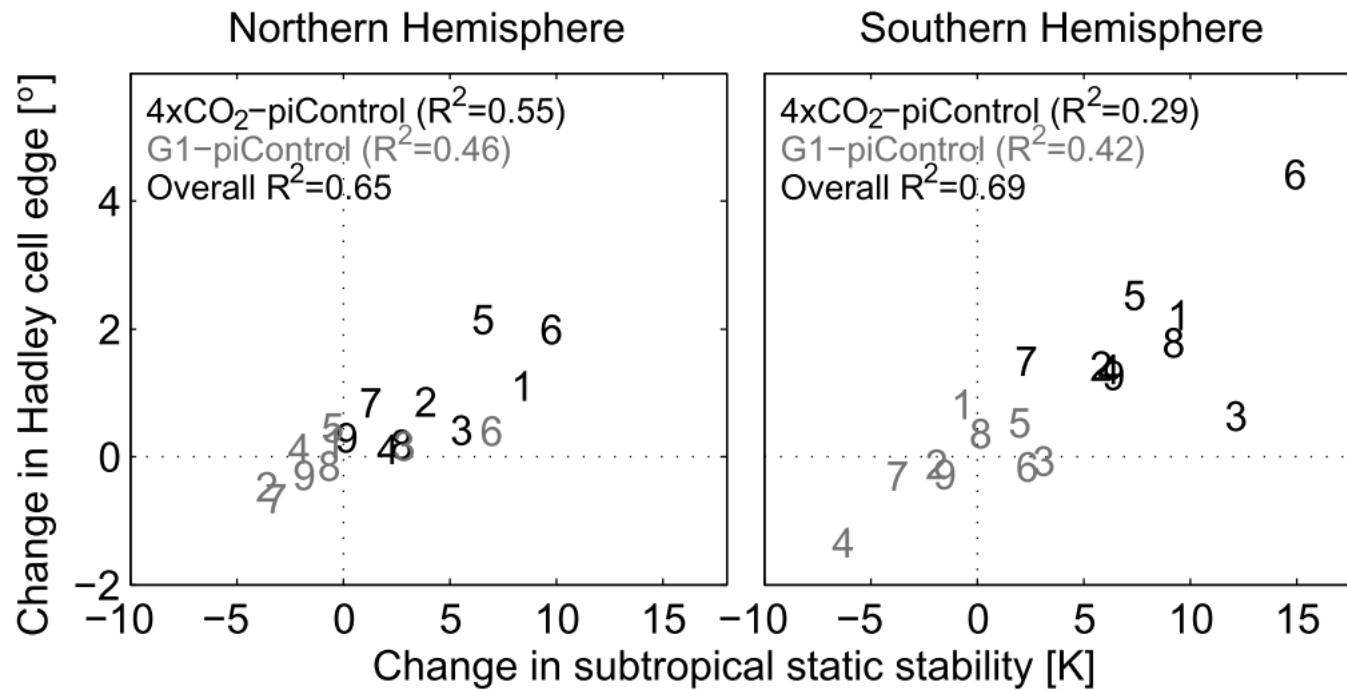
Menzel et al. 2019

Subtropical Static Stability

Connection between Hadley cell and subtropical static stability

Subtropical Static Stability

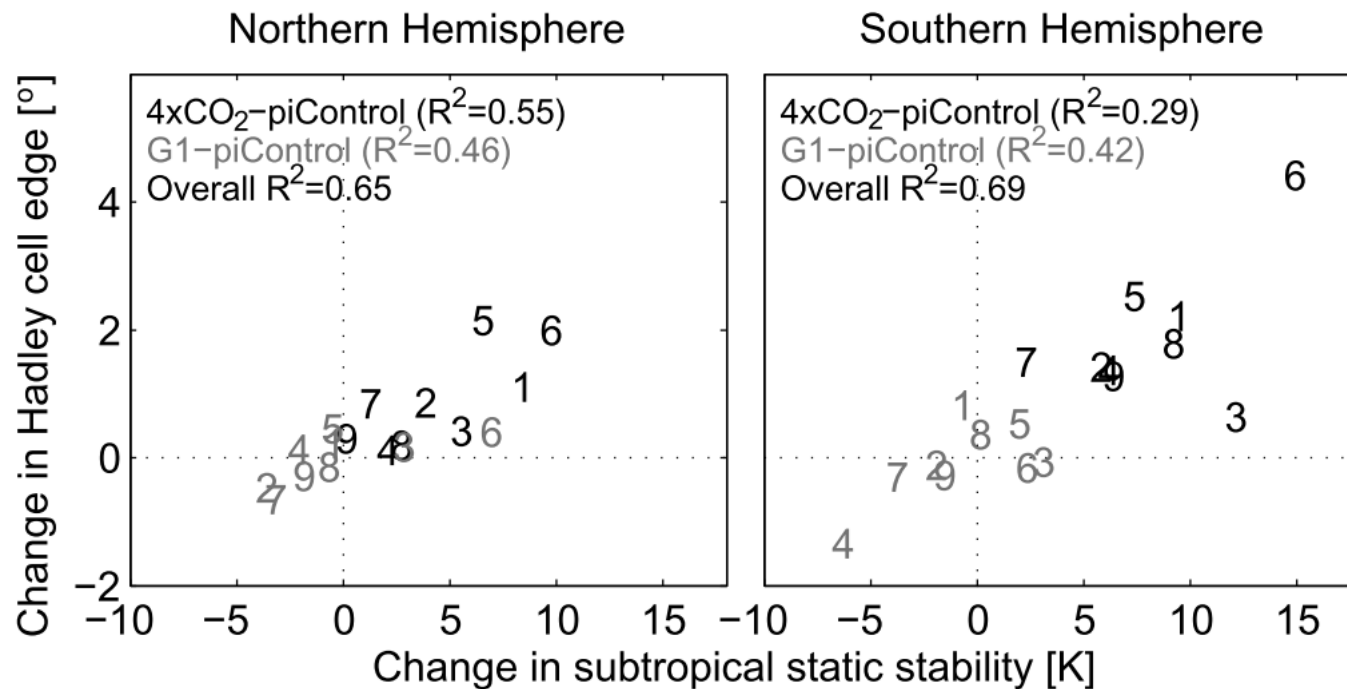
Connection between Hadley cell and subtropical static stability



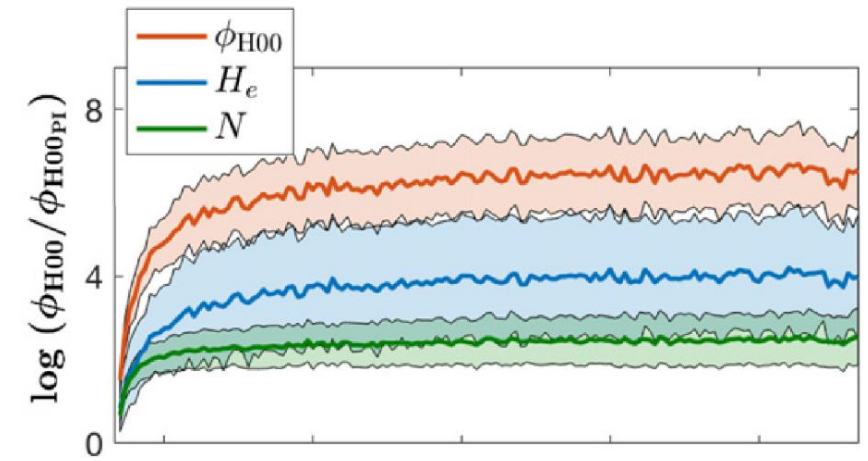
Davis et al. 2016

Subtropical Static Stability

Connection between Hadley cell and subtropical static stability

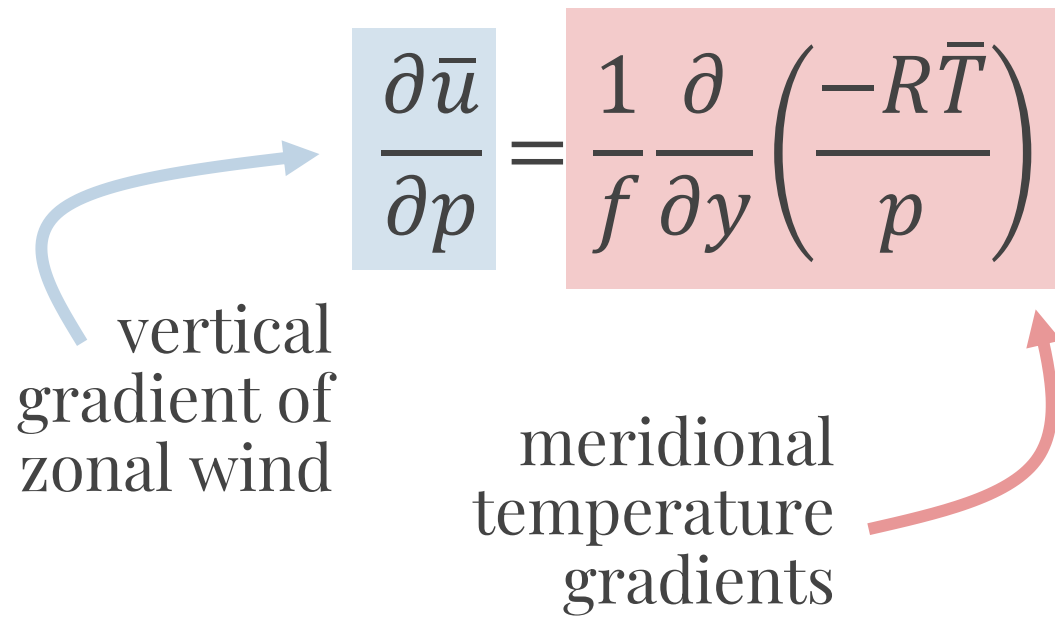


Davis et al. 2016



Chemke & Polvani 2019

Thermal Wind Balance



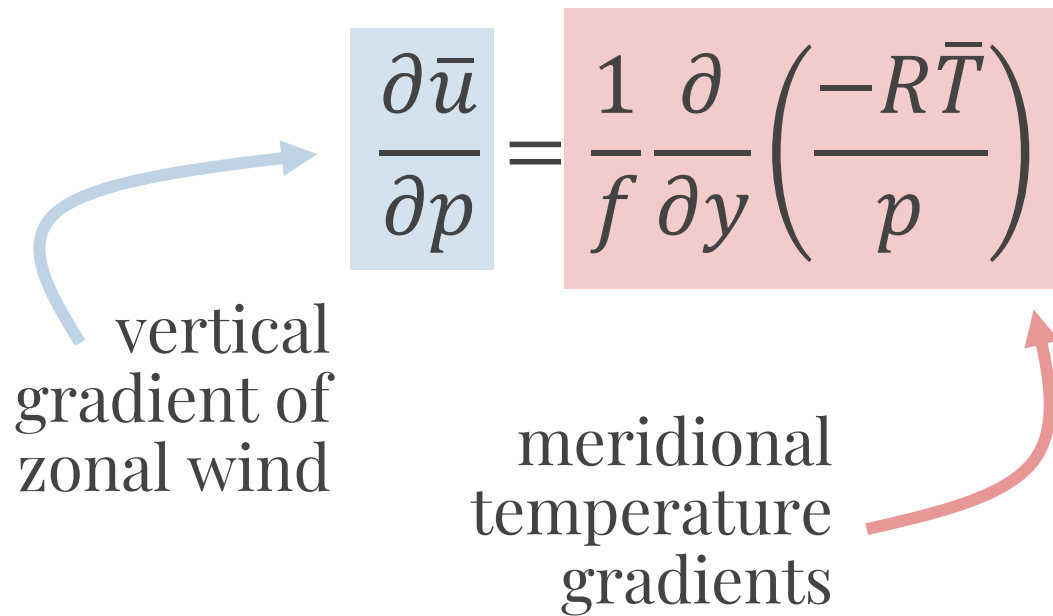
The diagram shows the Thermal Wind Balance equation: $\frac{\partial \bar{u}}{\partial p} = \frac{1}{f} \frac{\partial}{\partial y} \left(\frac{-R\bar{T}}{p} \right)$. A blue arrow points from the text 'vertical gradient of zonal wind' to the left side of the equation, $\frac{\partial \bar{u}}{\partial p}$. A red arrow points from the text 'meridional temperature gradients' to the right side of the equation, $\frac{1}{f} \frac{\partial}{\partial y} \left(\frac{-R\bar{T}}{p} \right)$.

$$\frac{\partial \bar{u}}{\partial p} = \frac{1}{f} \frac{\partial}{\partial y} \left(\frac{-R\bar{T}}{p} \right)$$

vertical
gradient of
zonal wind

meridional
temperature
gradients

Thermal Wind Balance



vertical
gradient of
zonal wind

$$\frac{\partial \bar{u}}{\partial p} = \frac{1}{f} \frac{\partial}{\partial y} \left(\frac{-R\bar{T}}{p} \right)$$

meridional
temperature
gradients

The diagram shows the thermal wind balance equation. A blue arrow points from the text 'vertical gradient of zonal wind' to the left side of the equation, $\frac{\partial \bar{u}}{\partial p}$. A red arrow points from the text 'meridional temperature gradients' to the right side of the equation, $\frac{1}{f} \frac{\partial}{\partial y} \left(\frac{-R\bar{T}}{p} \right)$.

$$\frac{\partial \bar{u}}{\partial p} \propto \frac{\partial T}{\partial y}$$

Upper level winds
proportional to
temperature gradients!

Thermal Wind Balance

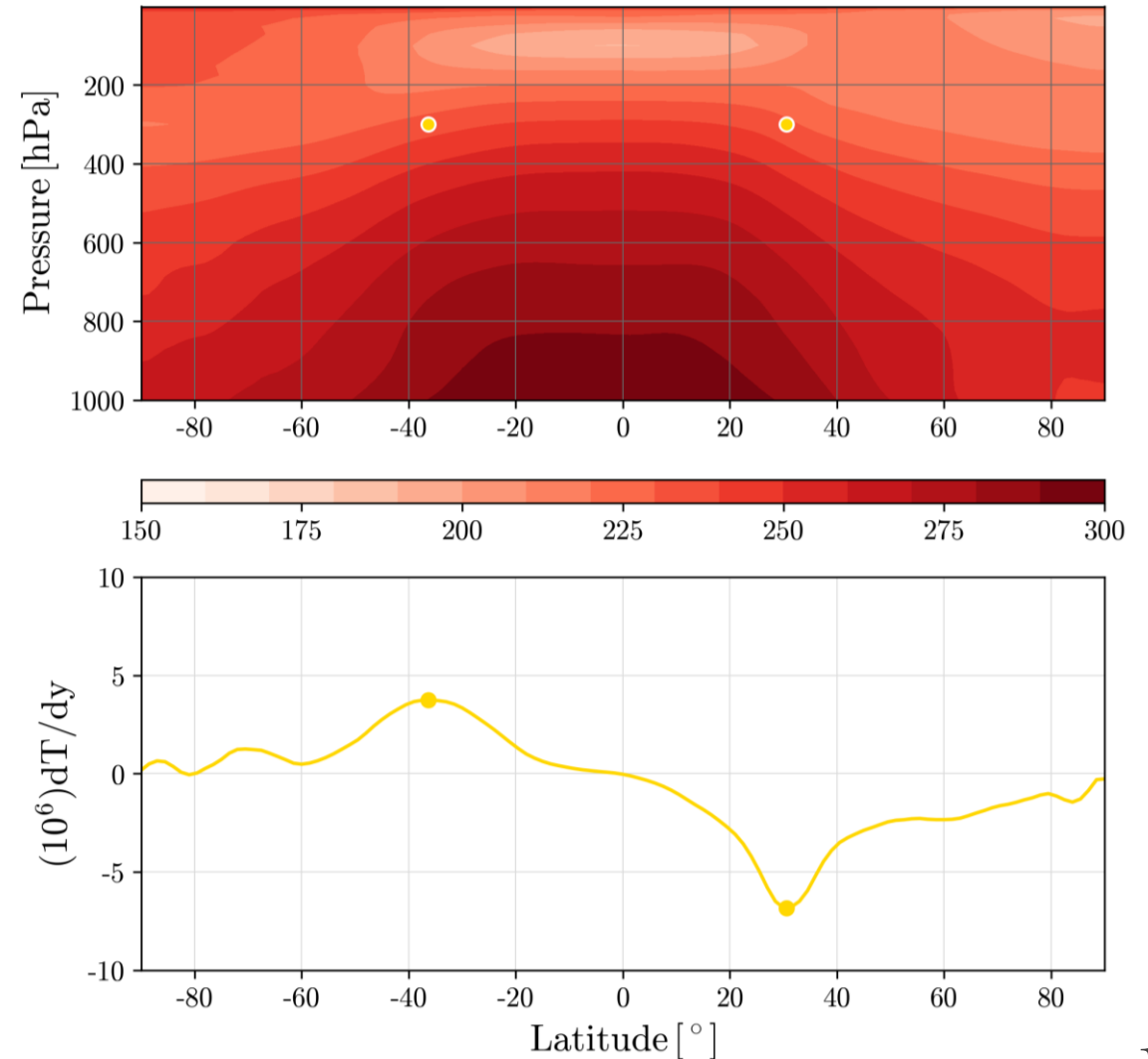
$$\frac{\partial \bar{u}}{\partial p} = \frac{1}{f} \frac{\partial}{\partial y} \left(\frac{-R\bar{T}}{p} \right)$$

Meridional Temperature Gradients

● $\varphi \frac{\partial T}{\partial y} = \varphi \left(\max \left(\frac{\partial T}{\partial y} \right)_{200-400 \text{ hPa}} \right)$

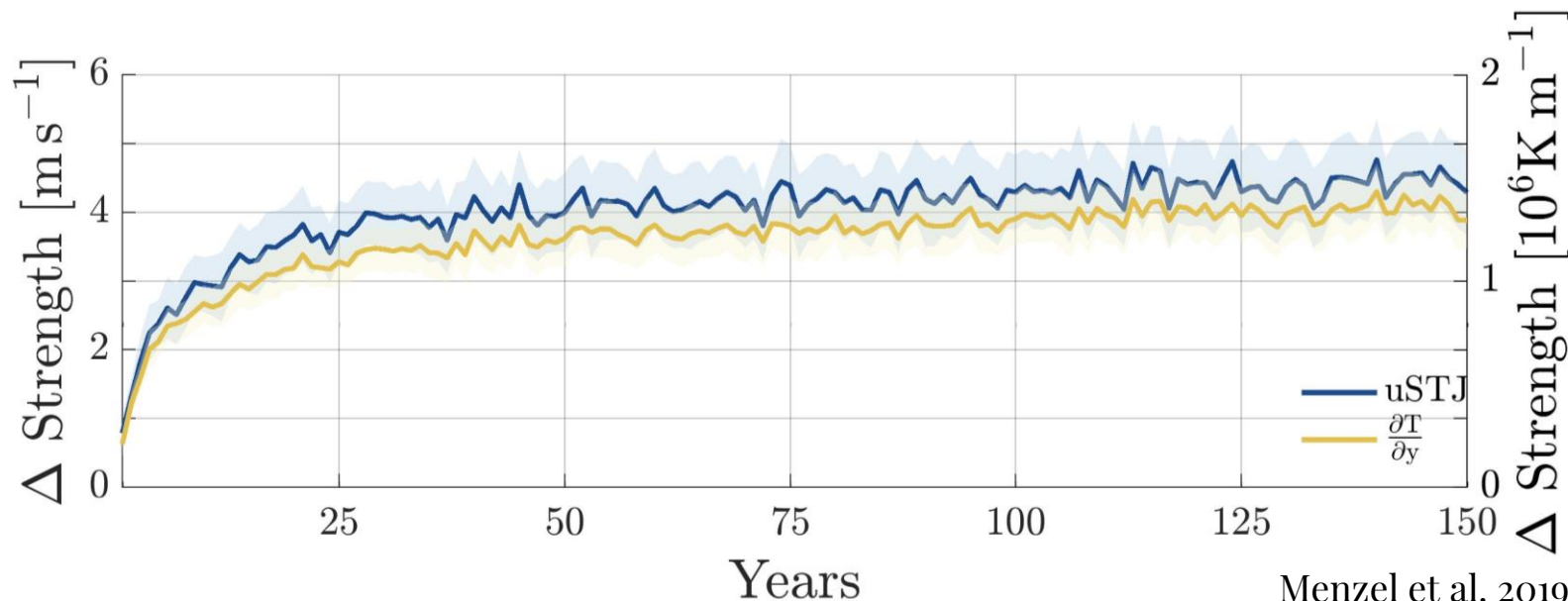
$$\max \frac{\partial T}{\partial y} = \max \left(\frac{\partial T}{\partial y} \right)_{200-400 \text{ hPa}}$$

Meridional Temperature Gradient



Thermal Wind Balance

	Shift/Change	τ
u_{STJ}	strengthening	40
$\partial T / \partial y$	strengthening	40

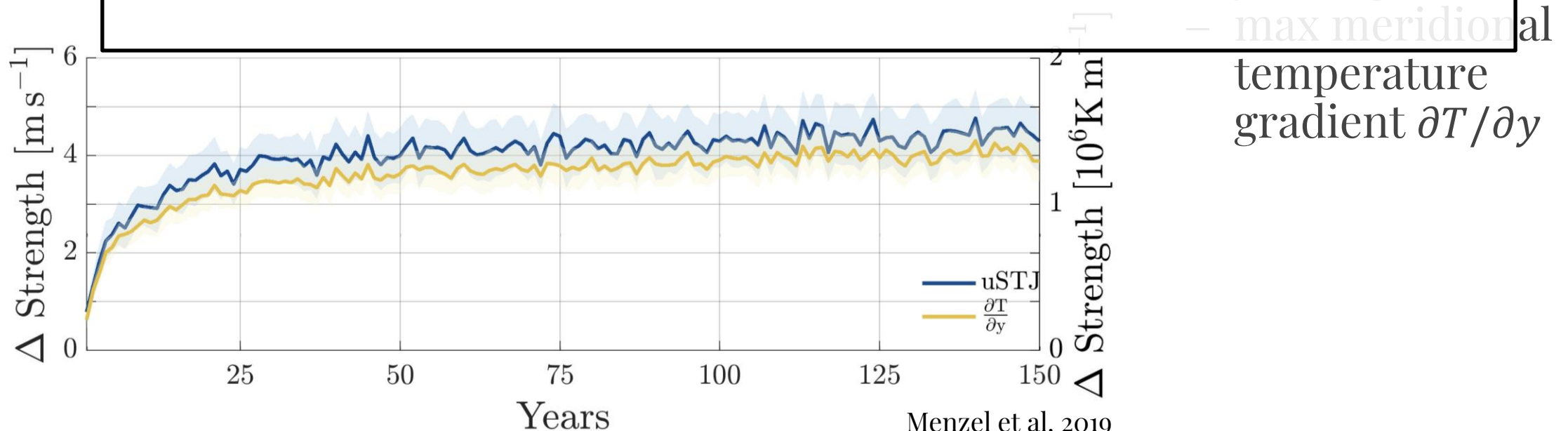


Menzel et al. 2019

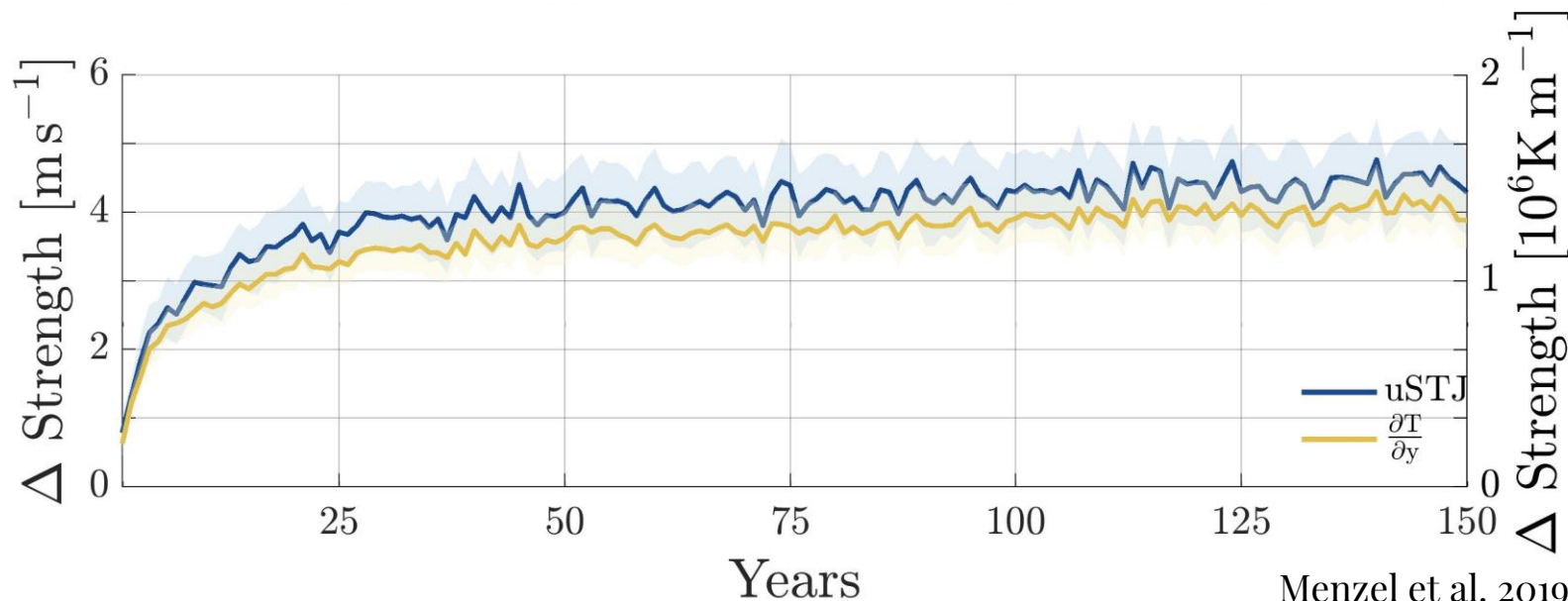
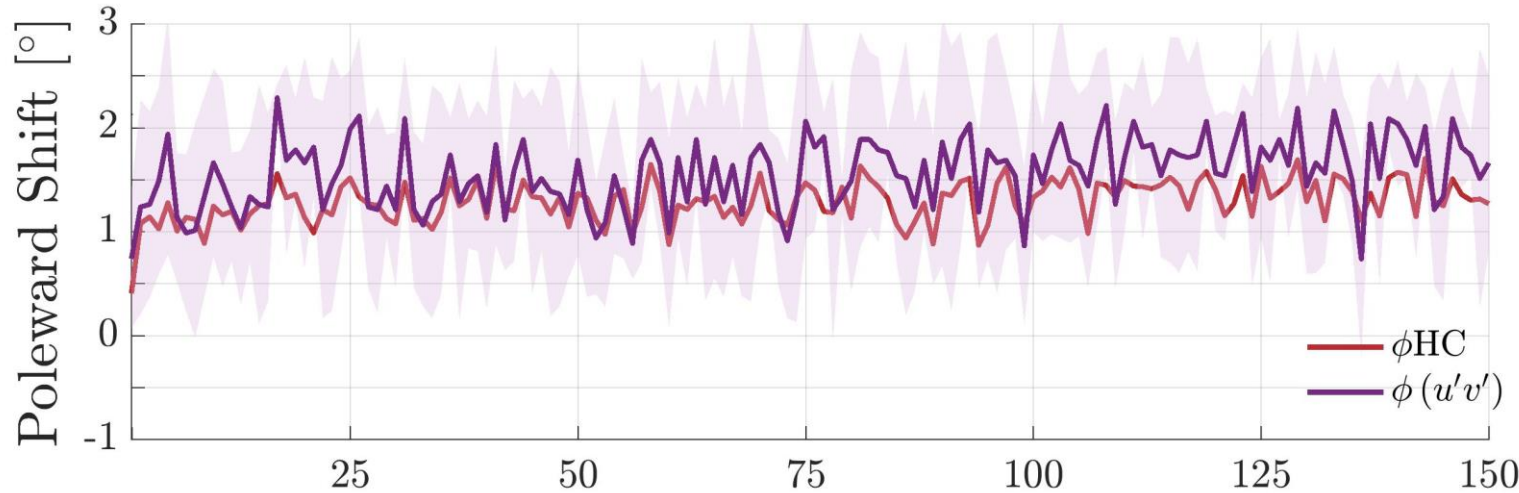
STJ strength:
 – max meridional
 temperature
 gradient $\partial T / \partial y$

Thermal Wind Balance

The subtropic jet's transient response follows
that of the meridional temperature gradients



Physical Balances



	Shift/Change	τ
ϕ_{HC}	poleward	7
$\phi(u'v')$	poleward	5
$uSTJ$	strengthening	40
$\partial T/\partial y$	strengthening	40

HC edge:

- latitude of max eddy momentum flux $\phi(u'v')$

STJ strength:

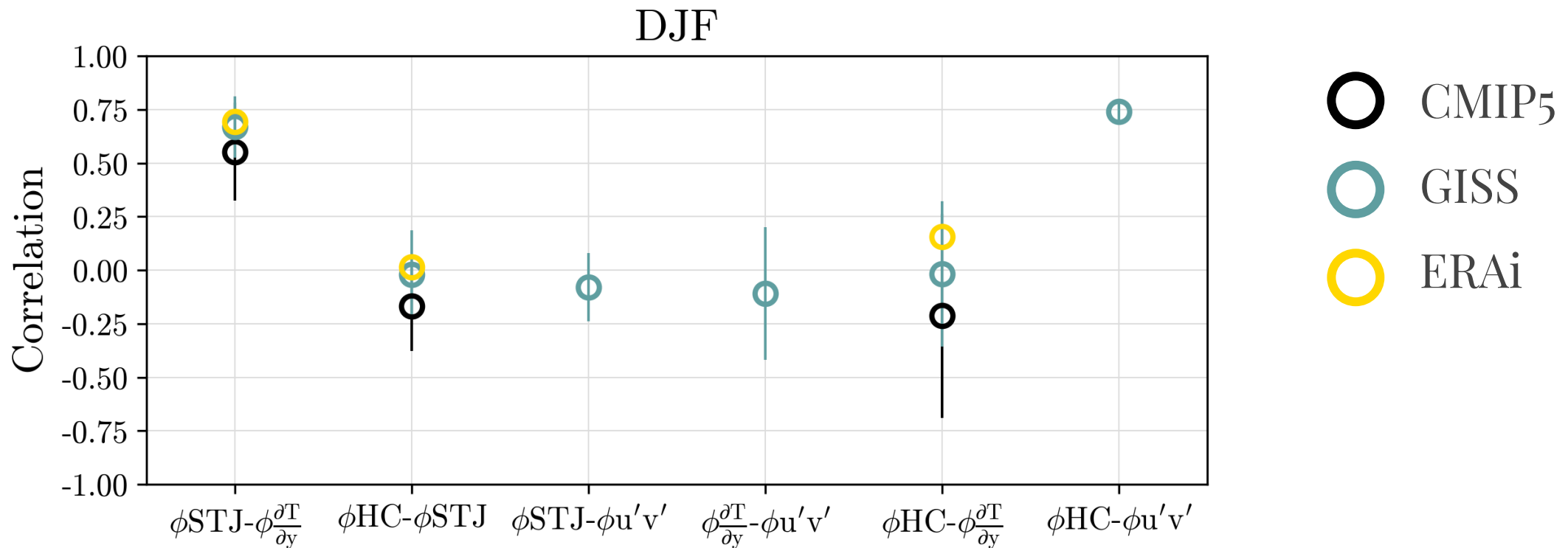
- max meridional temperature gradient $\partial T/\partial y$

Physical Balances

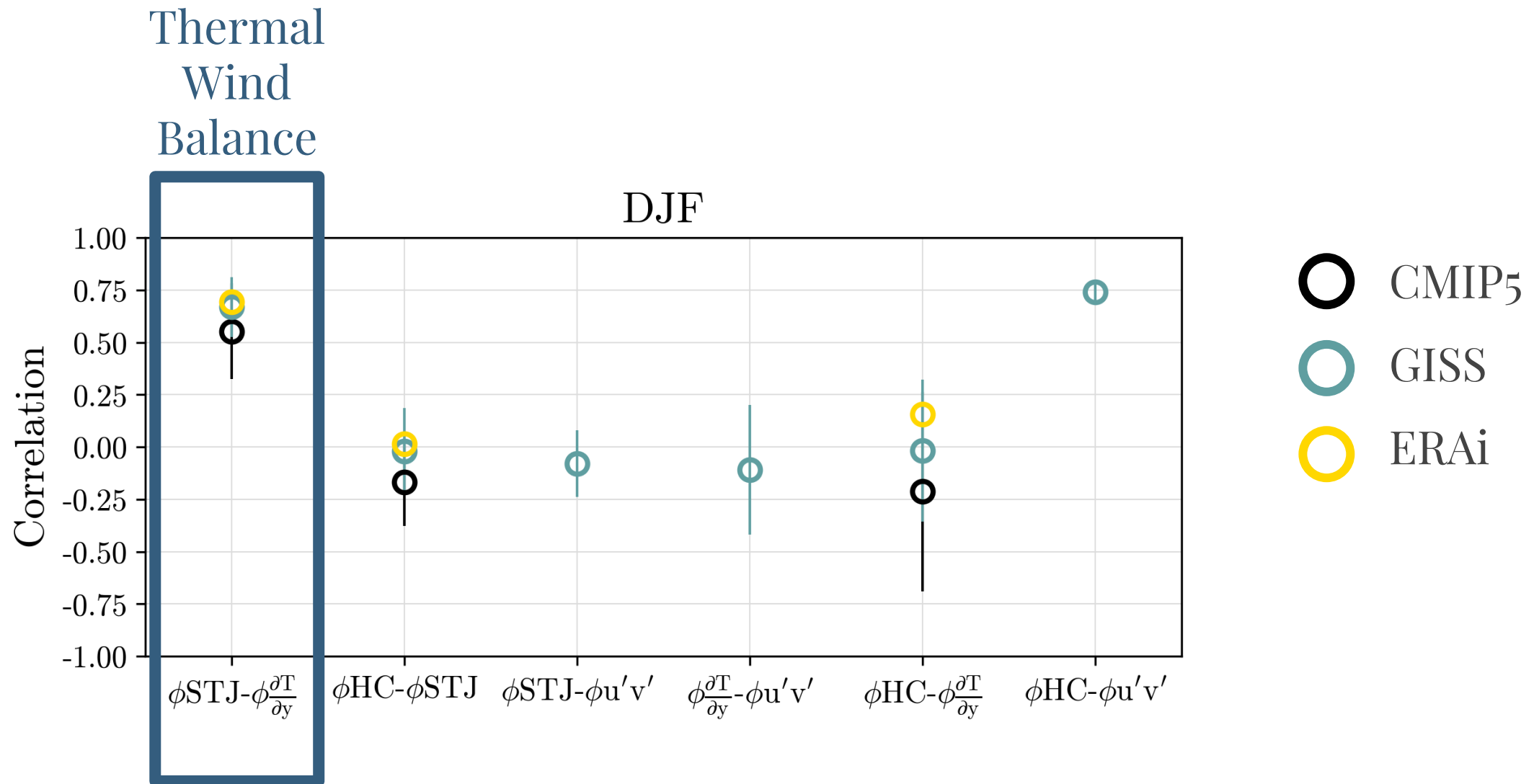
GISS-E2.1: NASA Goddard Institute for Space Studies' Global Climate Model

10 simulations, abrupt NxCO₂

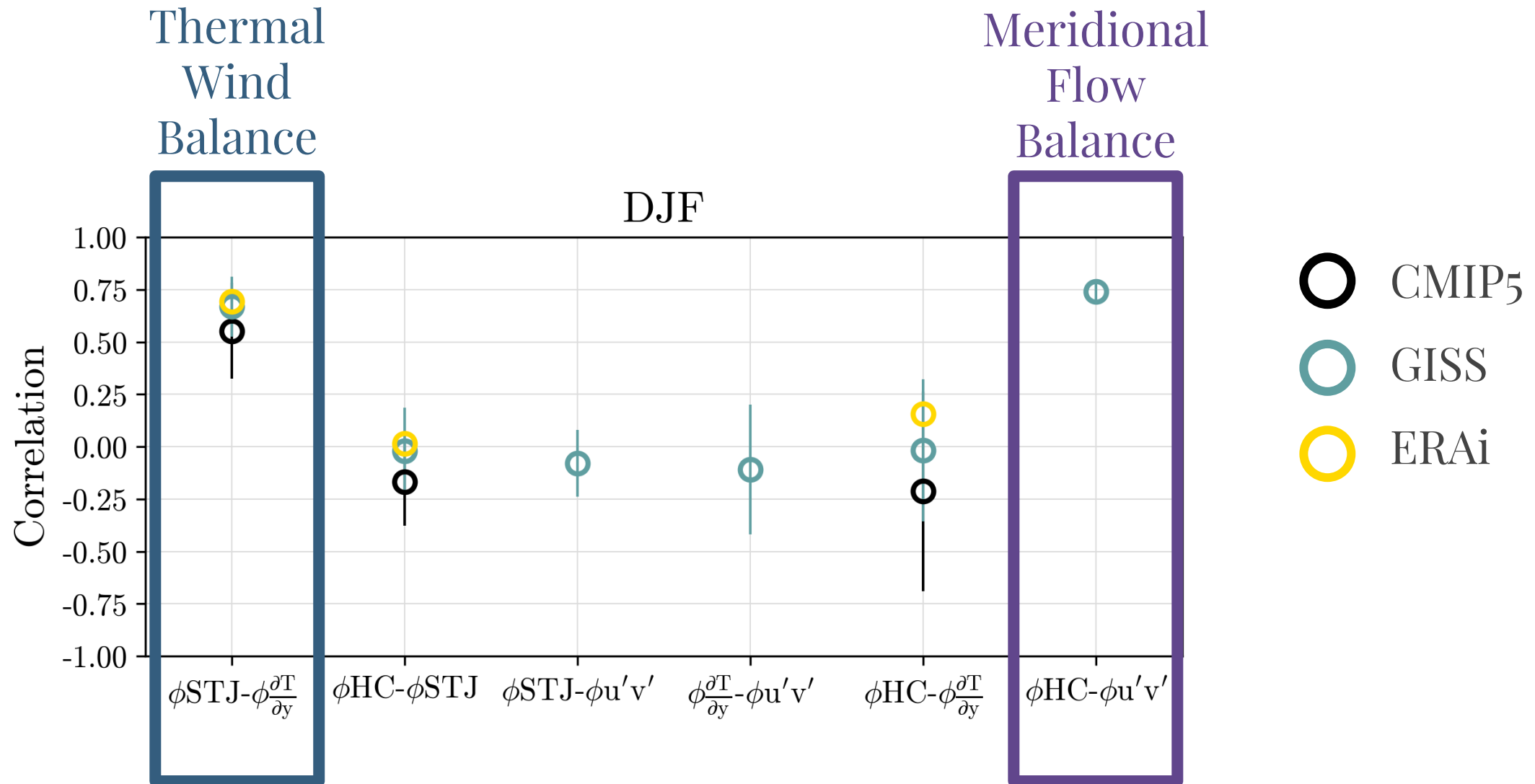
ERAi: ERA-interim reanalysis



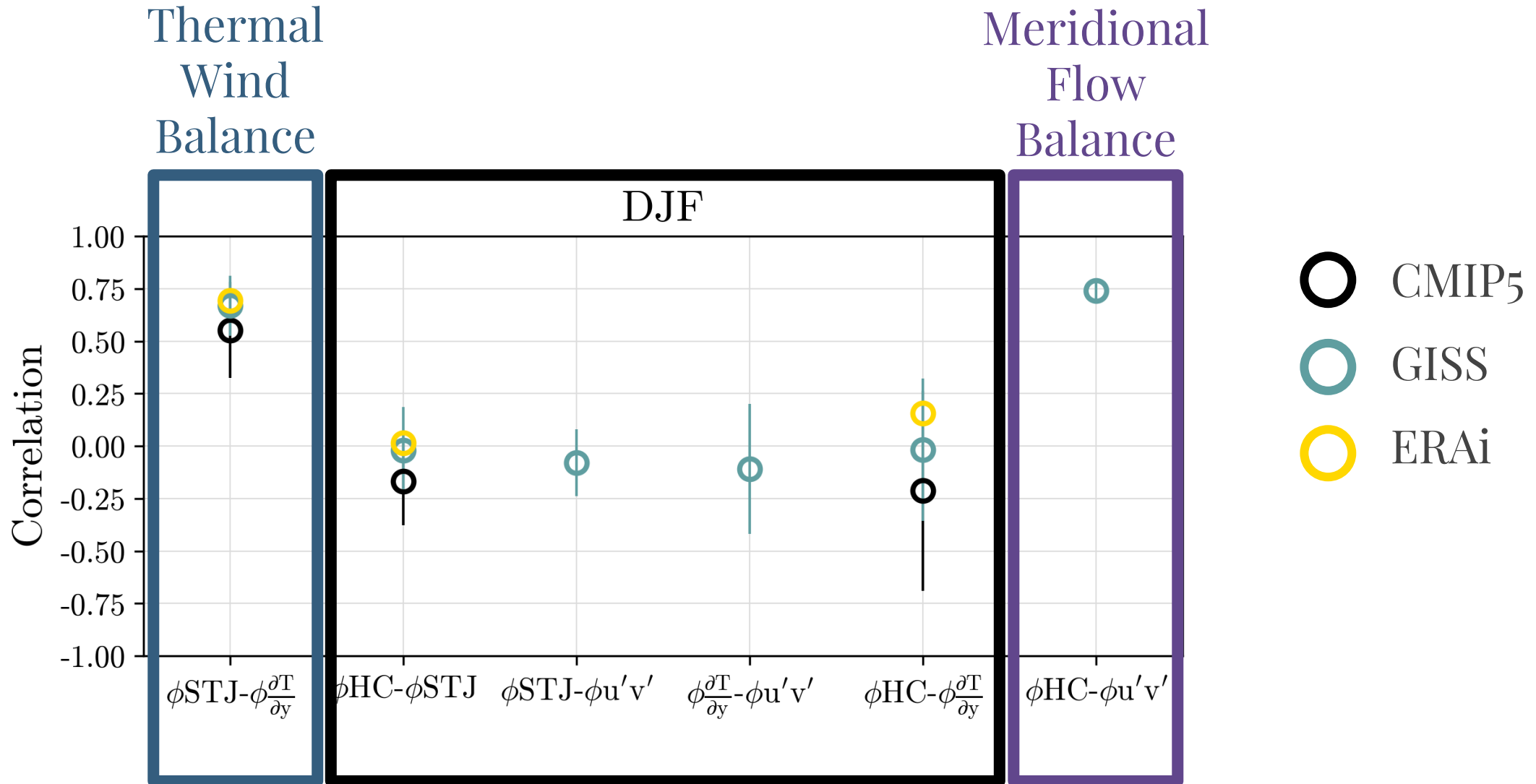
Physical Balances



Physical Balances

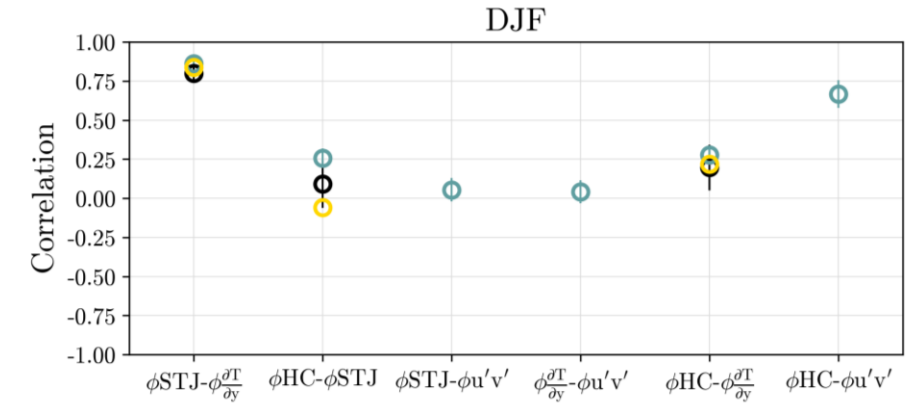
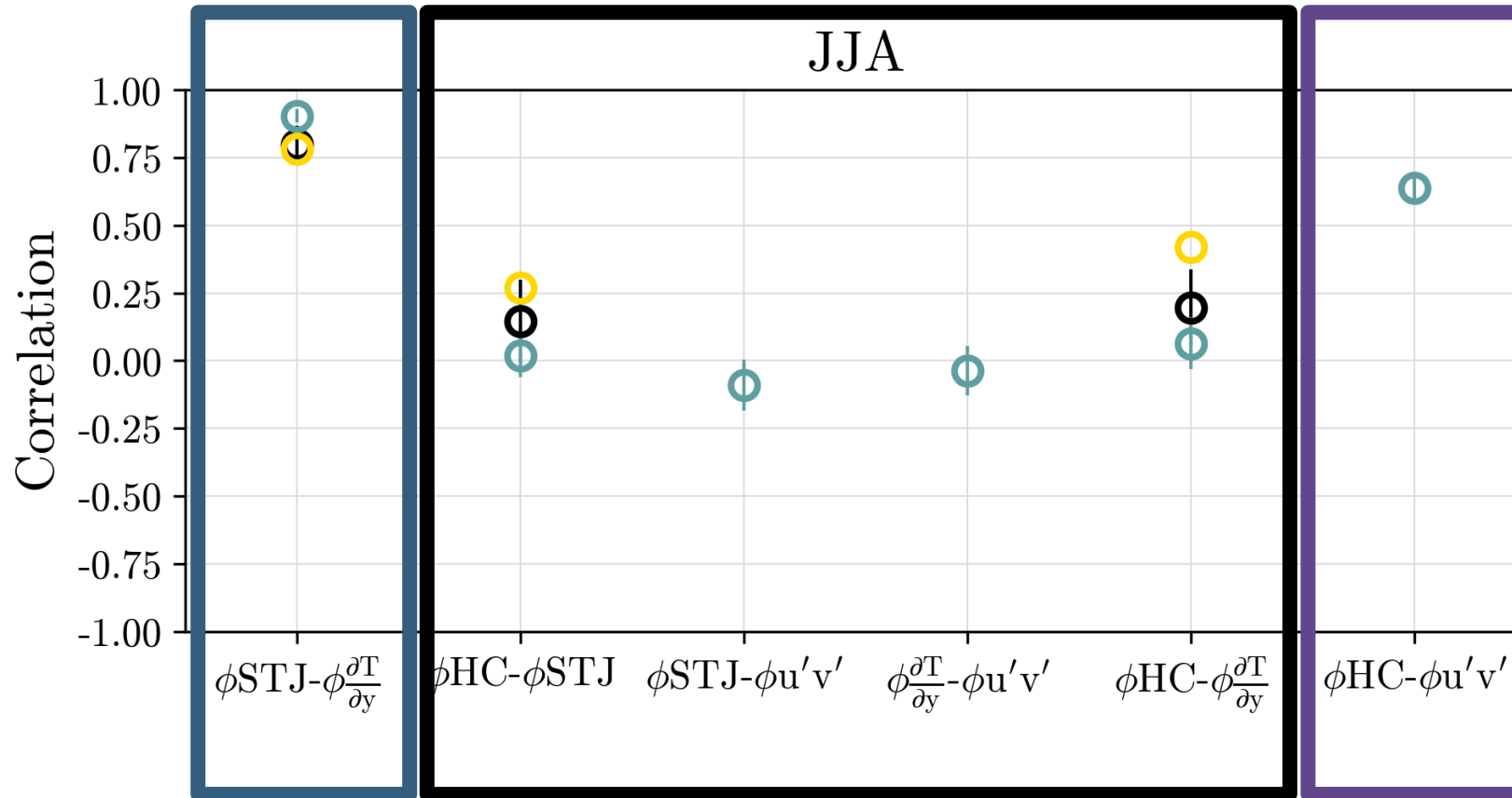


Physical Balances



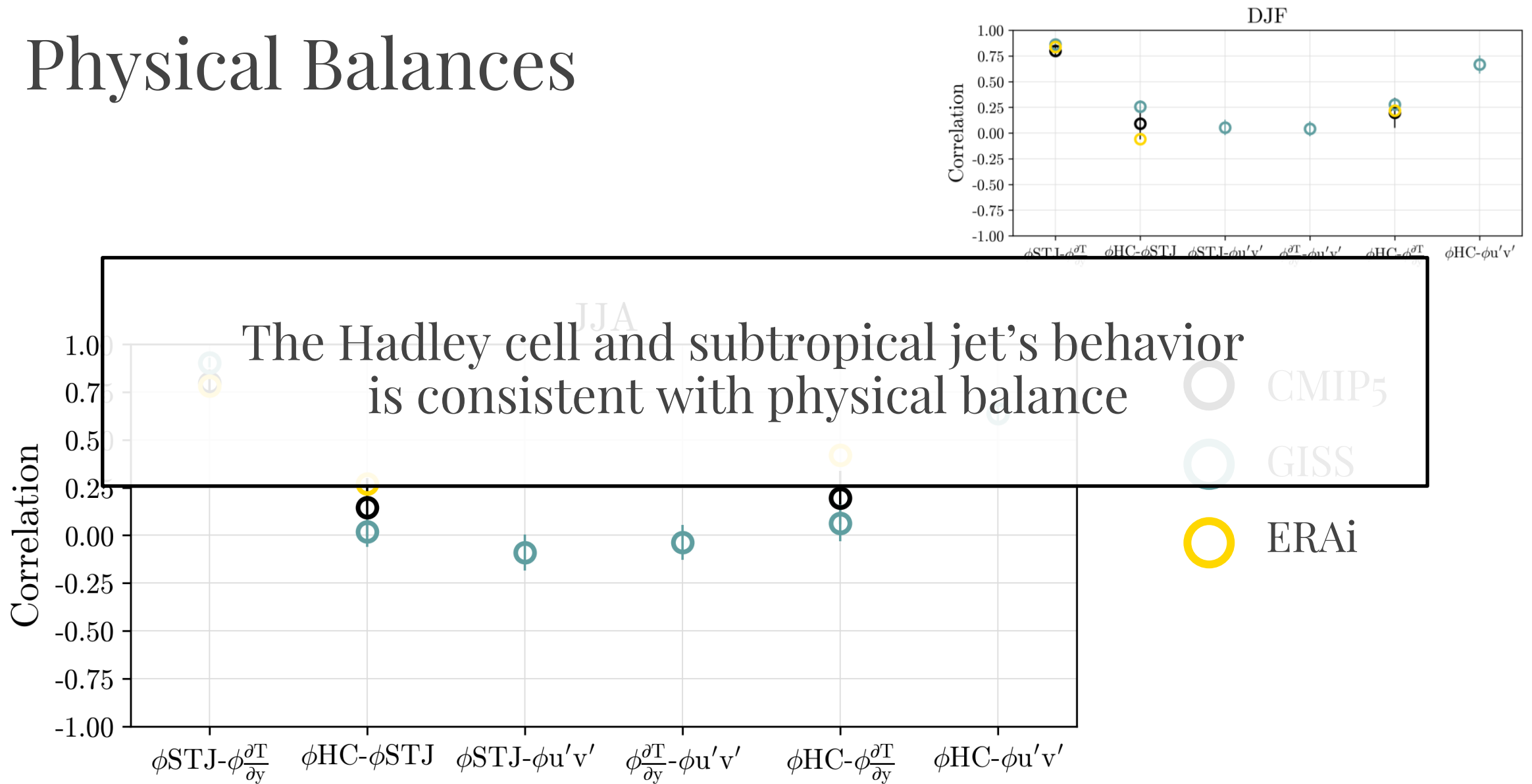
Physical Balances

Thermal
Wind
Balance



Meridional
Flow
Balance

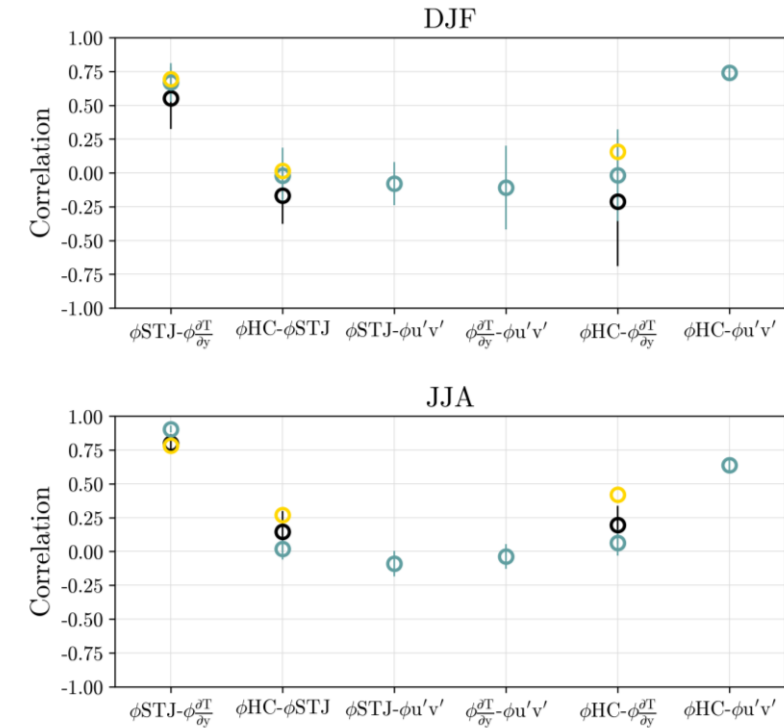
Physical Balances



Physical Balances Analysis

Conclusion:

The Hadley cell (HC)'s behavior is consistent with meridional flow balance and subtropical jet (STJ)'s behavior is consistent with thermal wind balance



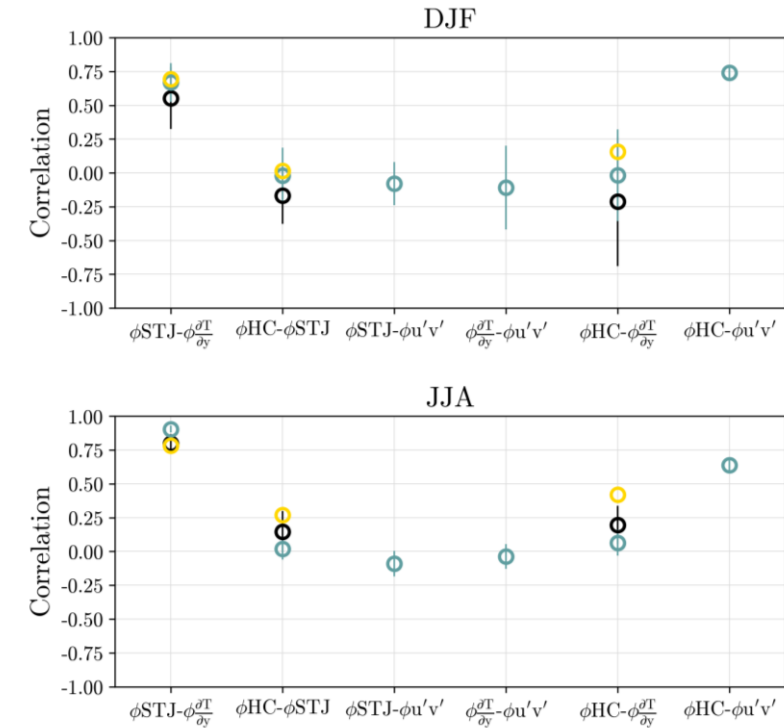
Physical Balances Analysis

Conclusion:

The Hadley cell (HC)'s behavior is consistent with meridional flow balance and subtropical jet (STJ)'s behavior is consistent with thermal wind balance

Lingering Question:

What model processes are necessary to replicate the STJ-HC relationship shown in comprehensive climate models?



Model Hierarchy

Hierarchy of Processes

decreasing complexity →

Large-Scale Atmospheric Circulation	Earth System Models	Atmospheric GCMs	Aquaplanet Simulations	Dry Dynamical Core
Dynamics	Primitive equations	Primitive equations	Primitive equations	Primitive equations
Forcing	Atmospheric composition	Atmospheric composition	Solar heating, atmospheric composition	Equilibrium temperature
Boundary conditions	Dynamical ocean, sea-ice, land models	Prescribed SSTs	Ocean mixed layer, “slab ocean”	informed by Maher et al. 2019
Diabatic processes	Moist processes	Moist processes	Gray radiation, simplified convection scheme	

Model Hierarchy

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Model Hierarchy

Hierarchy of Processes

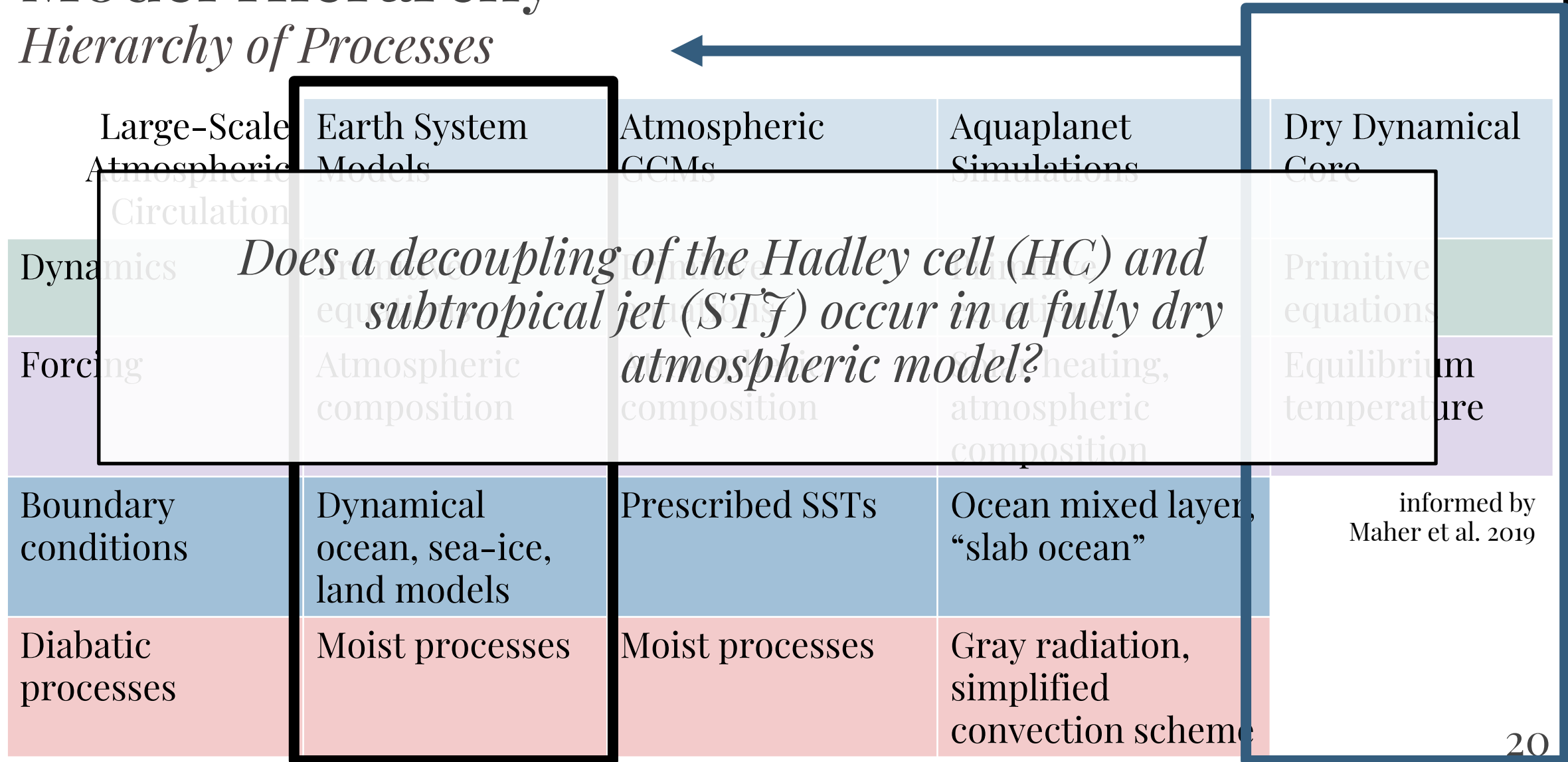
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	Earth System Models	Atmospheric GCMs	Aquaplanet Simulations	Dry Dynamical Core
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Model Hierarchy

Hierarchy of Processes

decreasing complexity →



Dry Dynamical Core Models

Forced by an equilibrium temperature profile

Classical Setup: Held & Suarez (1994)

$$\frac{\partial T}{\partial t} = \frac{T - T_{eq}}{\tau}$$

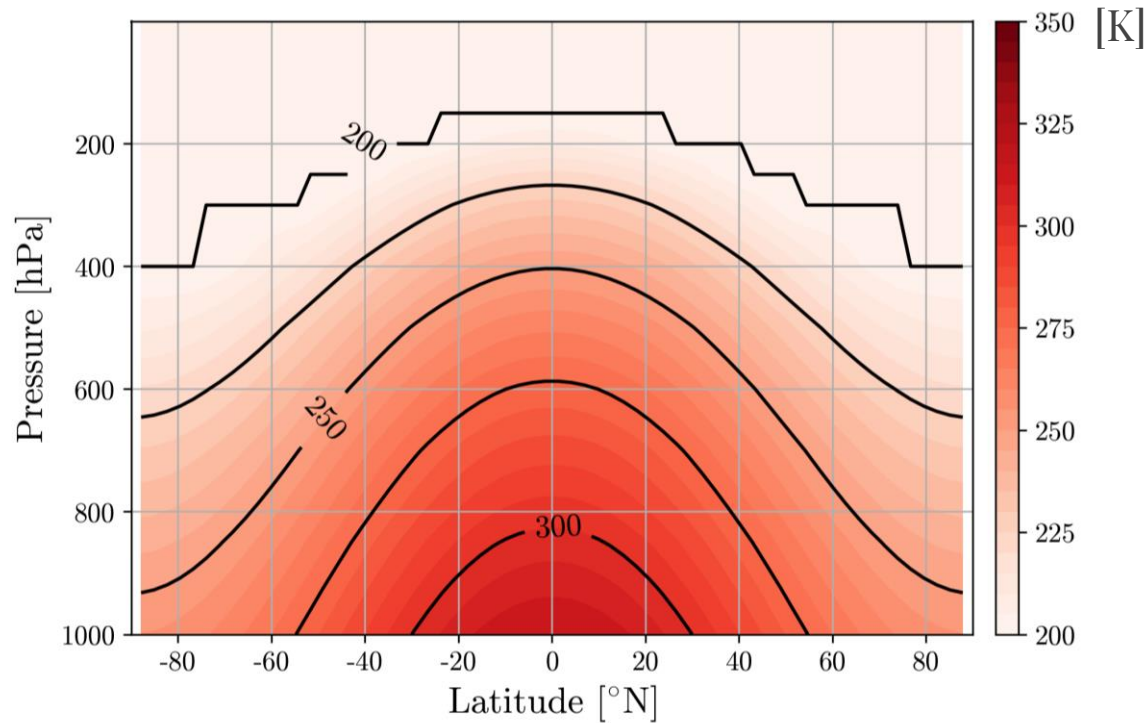
T_{eq} set by an analytical function

$$T_{eq} = \max \left\{ T_{strat}, \left[T_0 - \delta_y (\sin \phi)^2 + T' - \delta_z \log \left(\frac{p}{p_0} \right) (\cos \phi)^2 \right] \left(\frac{p}{p_0} \right)^\kappa \right\}$$

Dry Dynamical Core Models

Held & Suarez (1994)

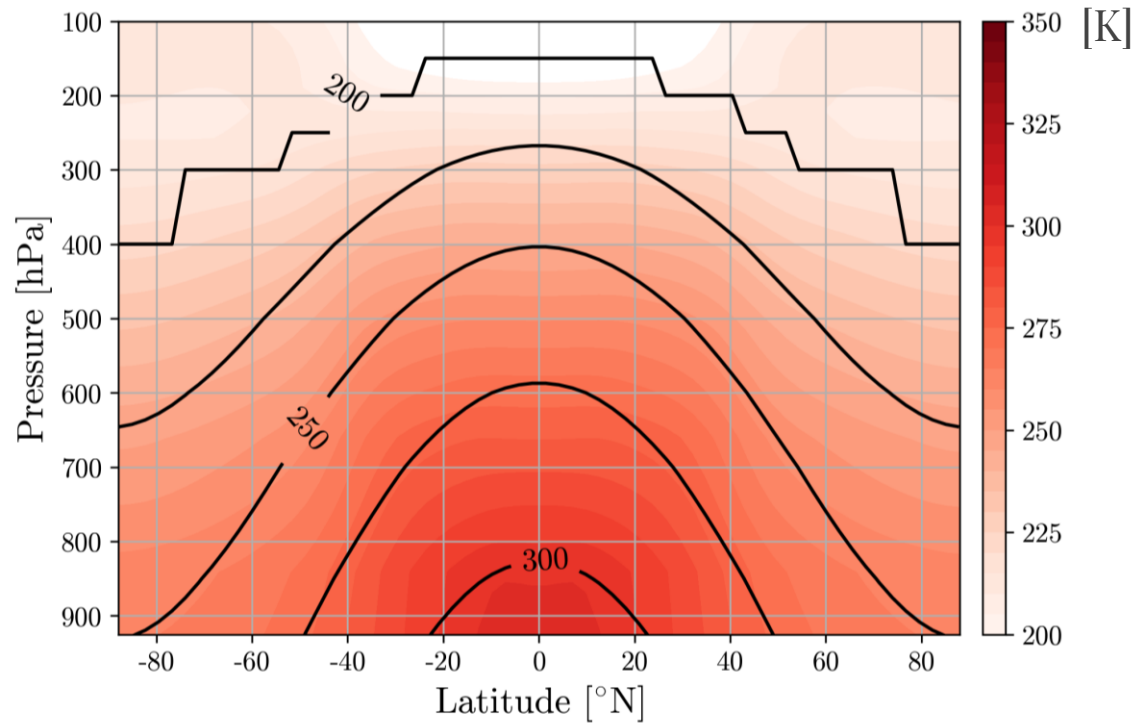
Equilibrium Temperature



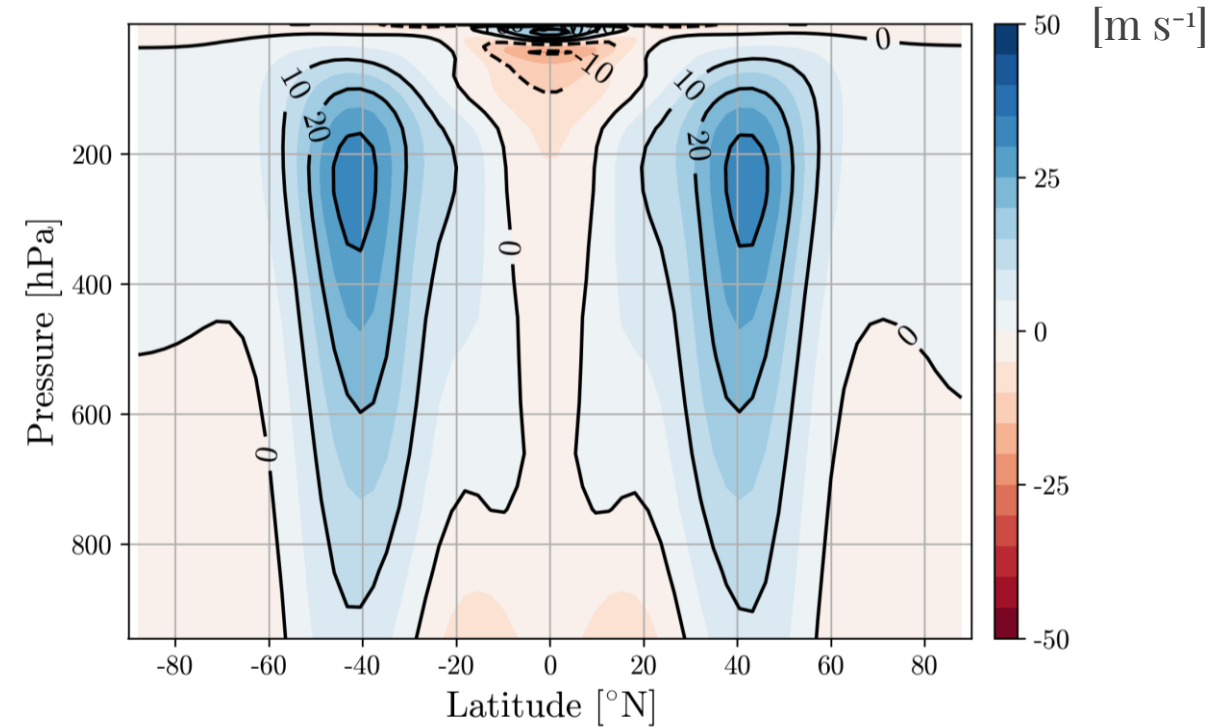
Dry Dynamical Core Models

Held & Suarez (1994)

Simulated Temperature



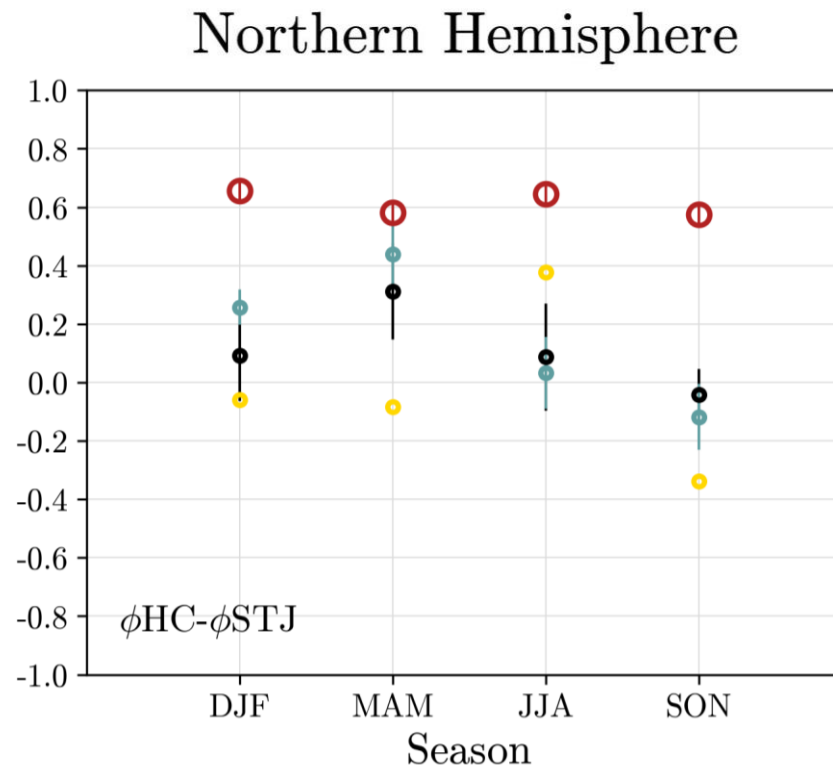
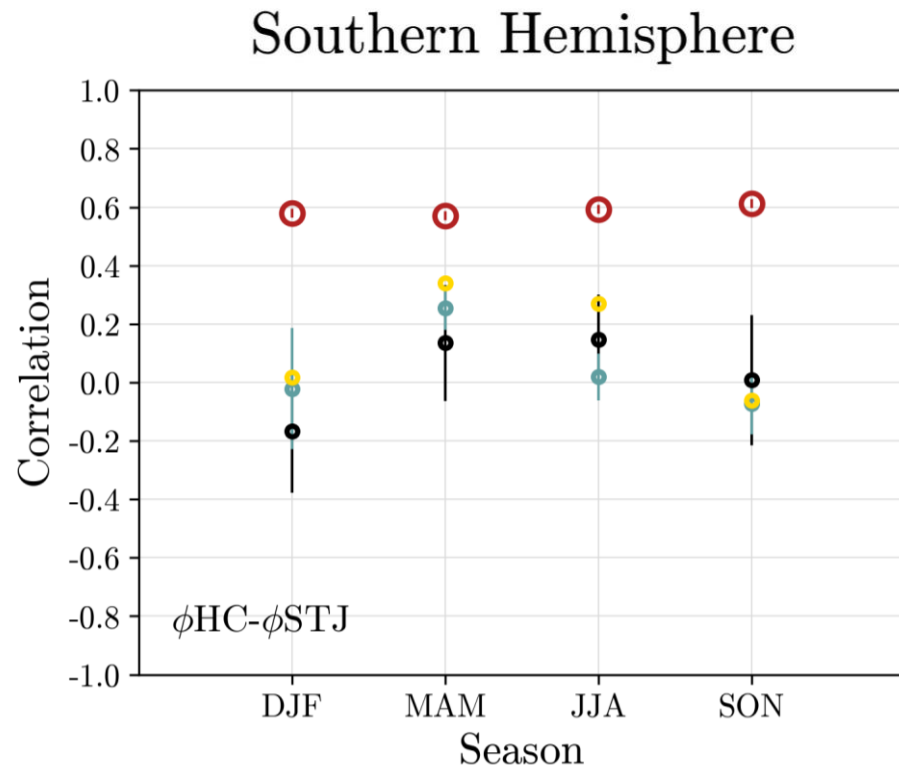
Simulated Wind



Equilibrium Temperature
(contour lines)

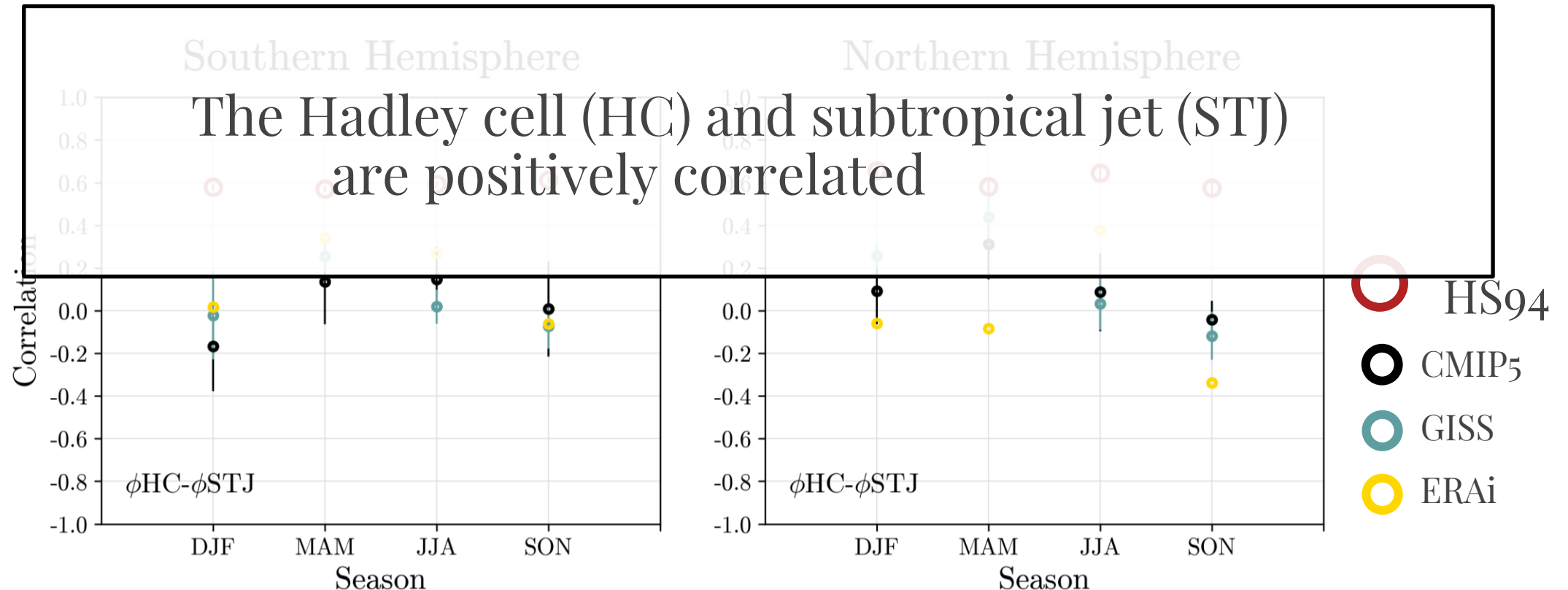
Dry Dynamical Core Models

- Held & Suarez (1994)
Analytical, zonally symmetric T_{eq}



Dry Dynamical Core Models

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Forced by an equilibrium temperature profile

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T_{eq} set by an analytical function

$$T_{eq} = \max \left\{ T_{strat}, \left[T_0 - \delta_y (\sin \phi)^2 + T' - \delta_z \log \left(\frac{p}{p_0} \right) (\cos \phi)^2 \right] \left(\frac{p}{p_0} \right)^\kappa \right\}$$

New Setup: Wu & Reichler (2018)

T_{eq} derived by iteration to improve accuracy

$$T_{eq} = T(\lambda, \phi, p, t)$$

Dry Dynamical Core Models

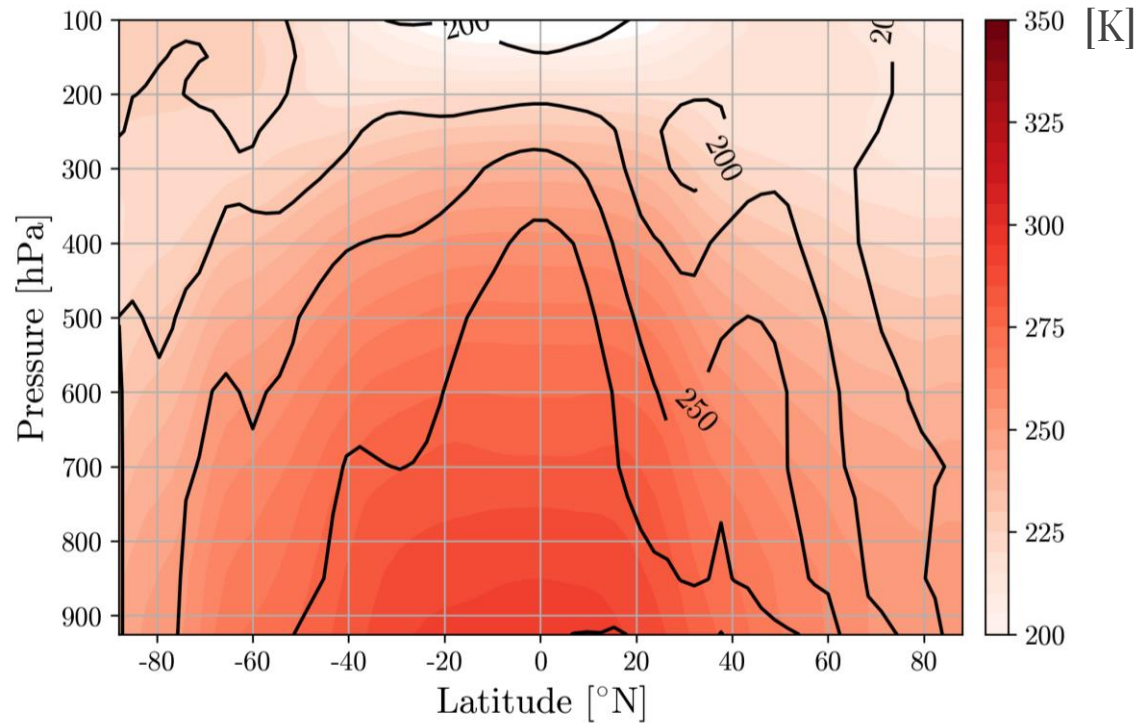
decreasing complexity 

	Wu & Reichler (2018)	Held & Suarez (1994)
T_{eq} Zonal Profile	Zonally varying	Zonally symmetric
Seasonality	Seasonally varying	No seasonality
Topography	Realistic topography	No topography
Stratosphere	Improved (Jucker et al. 2014)	

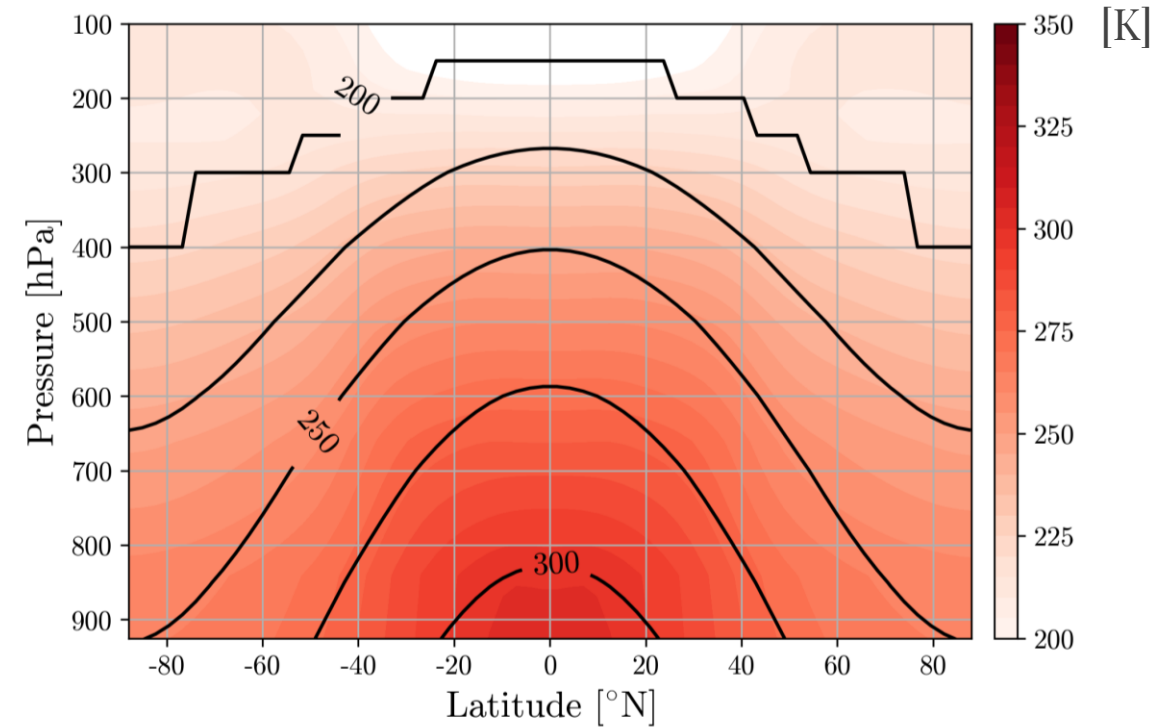
Dry Dynamical Core Models

decreasing complexity →

Wu & Reichler (2018)



Held & Suarez (1994)



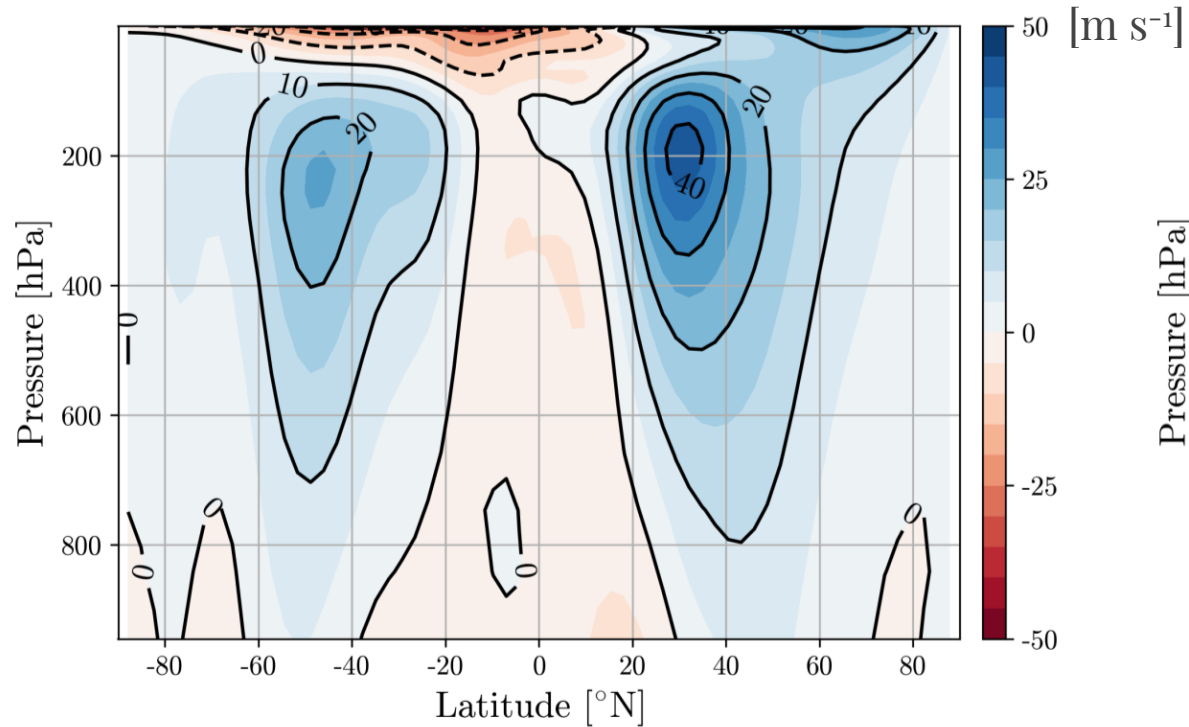
DJF
Climatology

Simulated Temperature,
Equilibrium Temperature (contour lines)

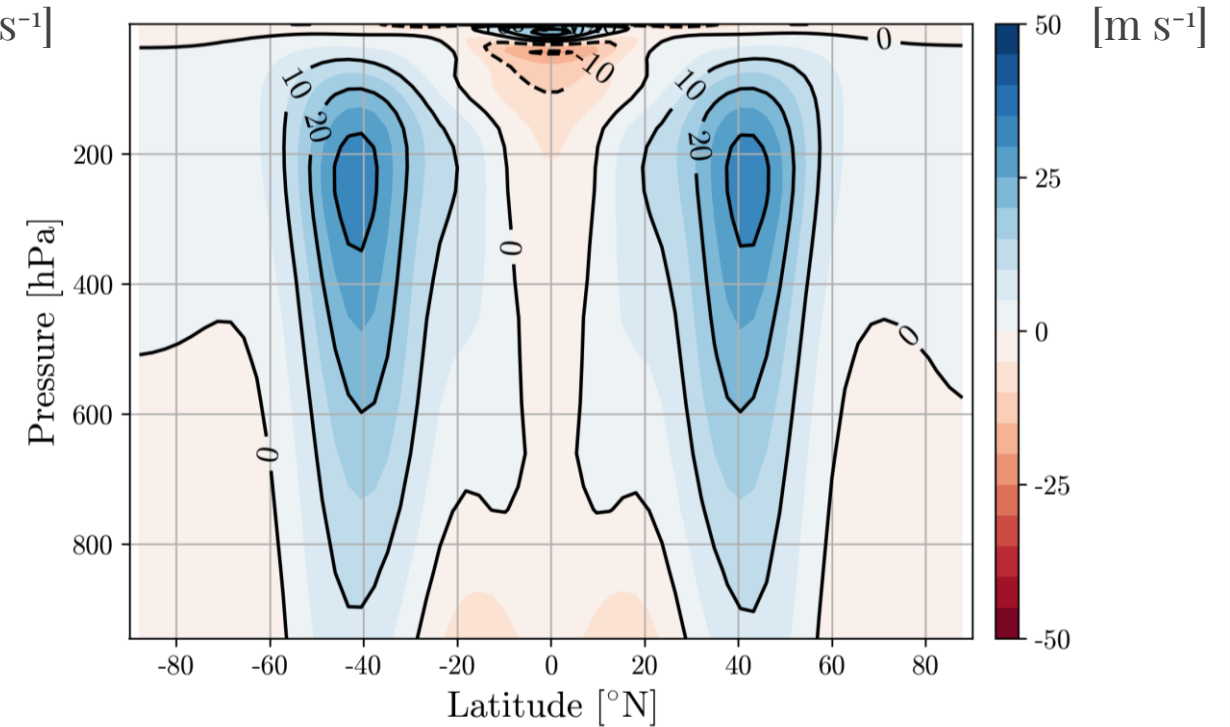
Dry Dynamical Core Models

decreasing complexity →

Wu & Reichler (2018)



Held & Suarez (1994)



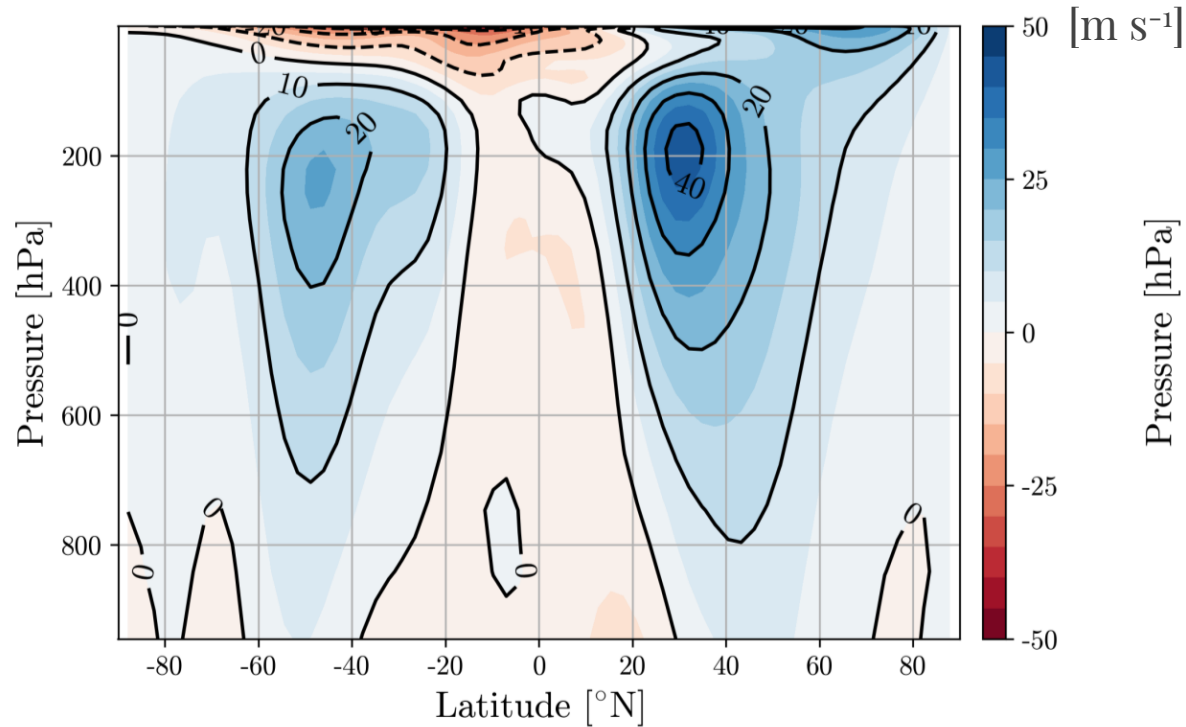
DJF
Climatology

Simulated Wind

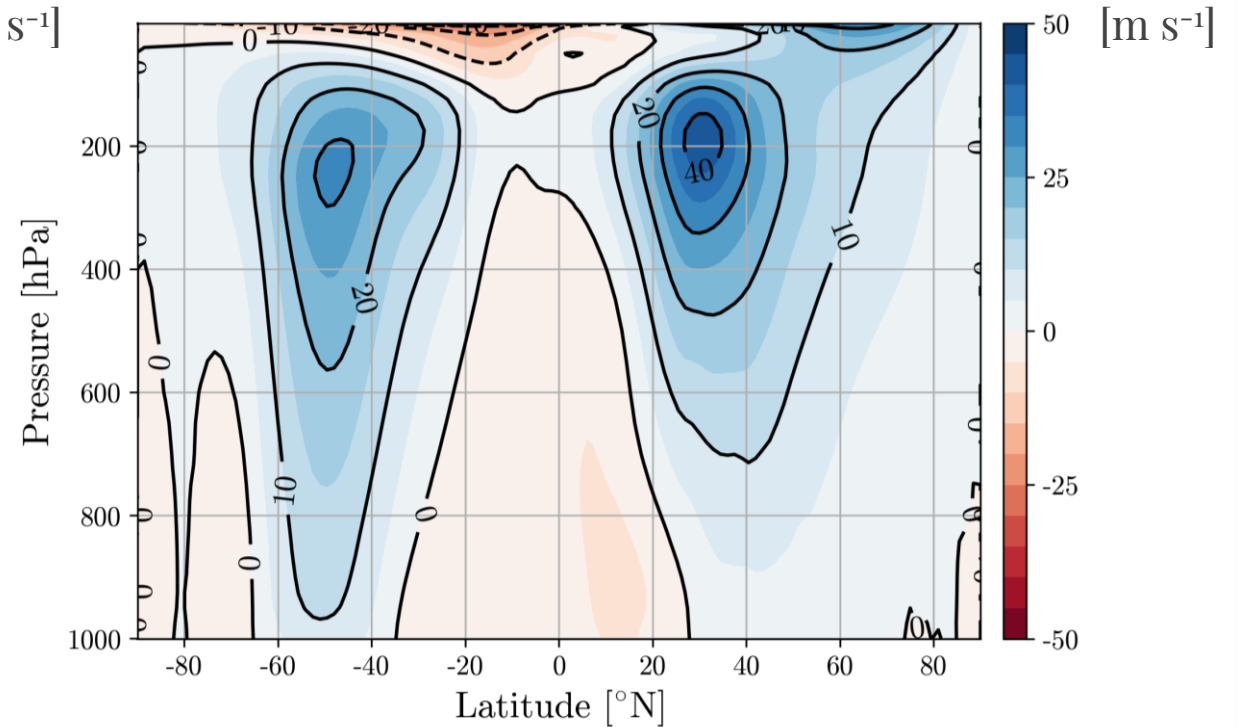
Dry Dynamical Core Models

decreasing complexity →

Wu & Reichler (2018)



ERA-interim Reanalysis

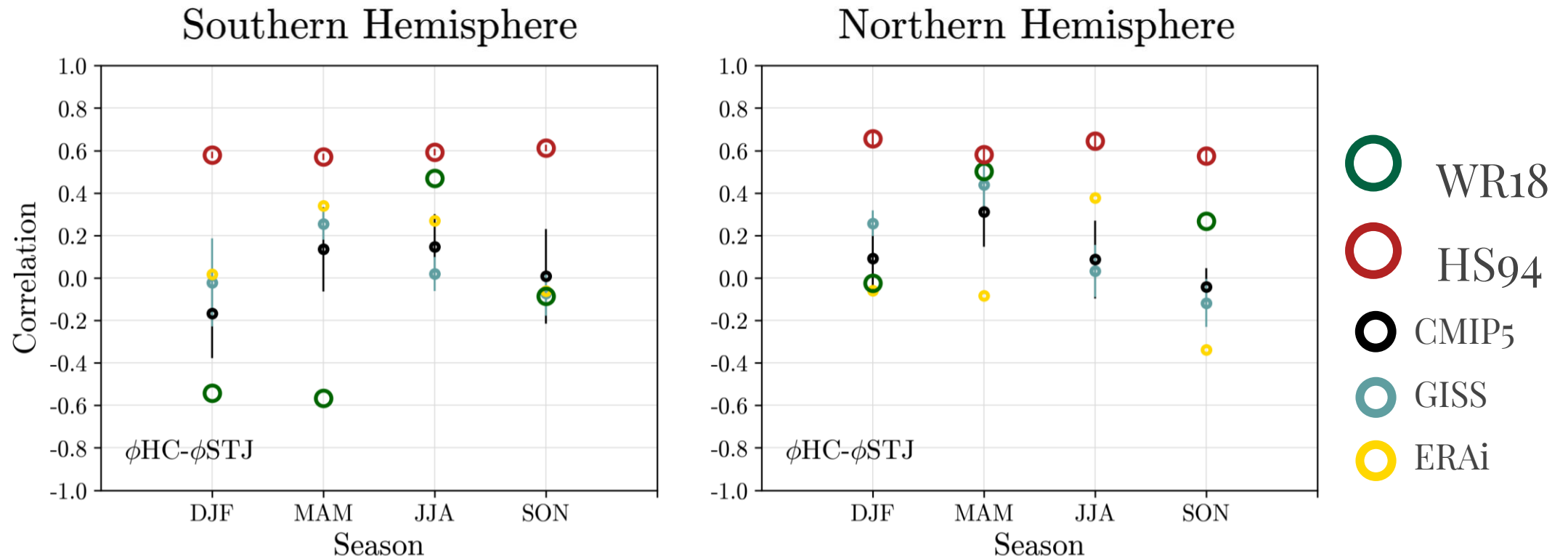


DJF
Climatology

Simulated Wind

Dry Dynamical Core Models

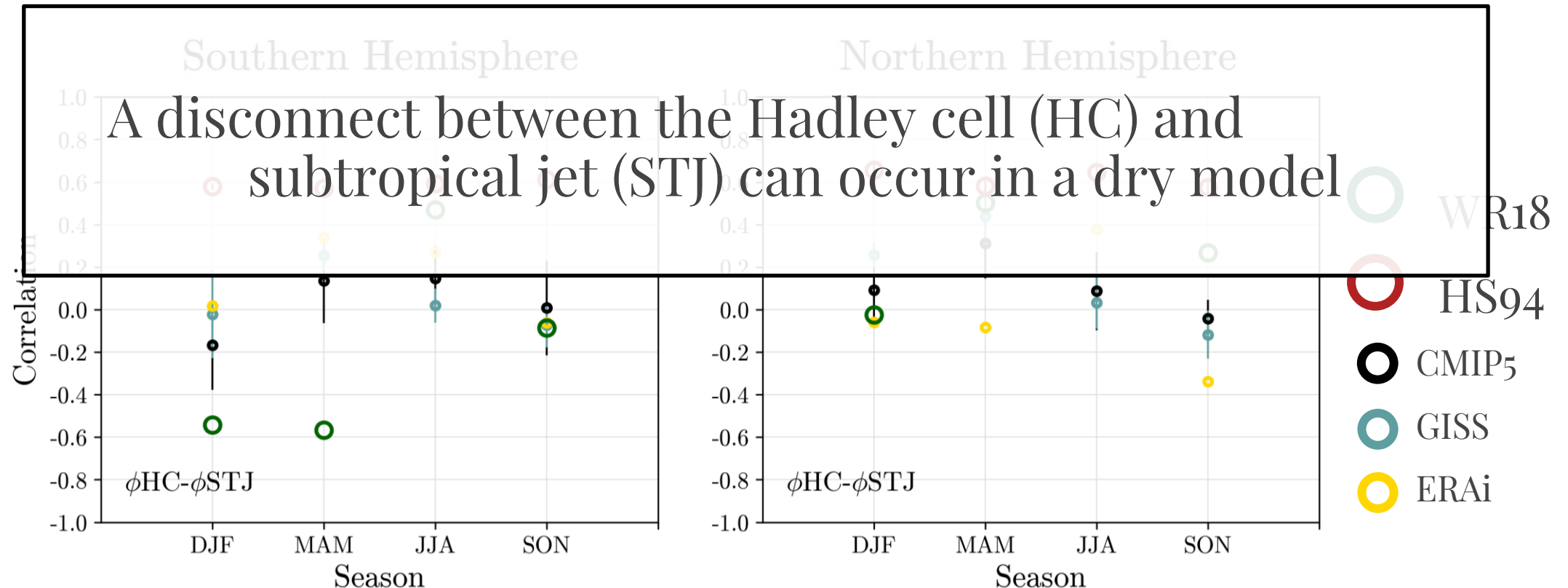
- Wu & Reichler (2018)
Derived by iteration, zonally and seasonally-varying T_{eq}



Dry Dynamical Core Models

○ Wu & Reichler (2018)

Derived by iteration, zonally and seasonally-varying T_{eq}



Dry Dynamical Core Models

A disconnect between the Hadley cell (HC) and subtropical jet (STJ) can occur in a dry model

Follow-up Question:

Is this a result of zonal variability in the “forcing”?

Dry Dynamical Core Models

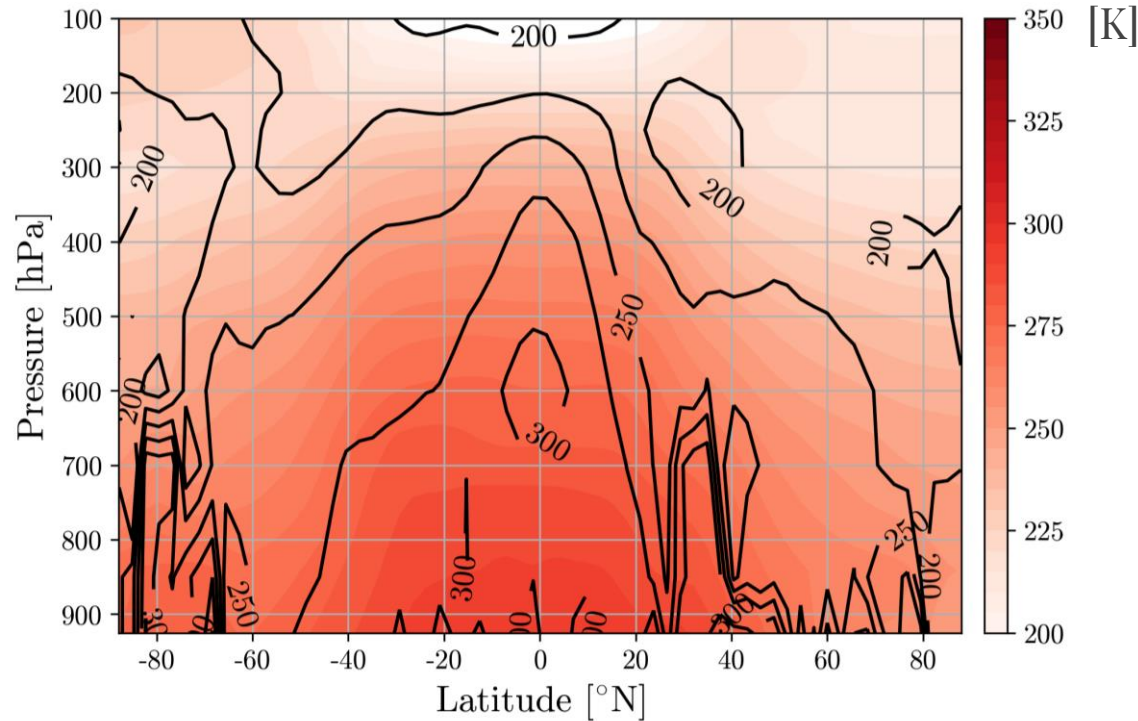
decreasing complexity 

	Wu & Reichler (2018)	Wu & Reichler (2018), Zonal	Held & Suarez (1994)
T_{eq} Zonal Profile	Zonally varying	Zonally symmetric	Zonally symmetric
Seasonality	Seasonally varying	Seasonally varying	No seasonality
Topography	Realistic topography	No topography	No topography
Stratosphere	Improved (Jucker et al. 2014)	Improved (Jucker et al. 2014)	

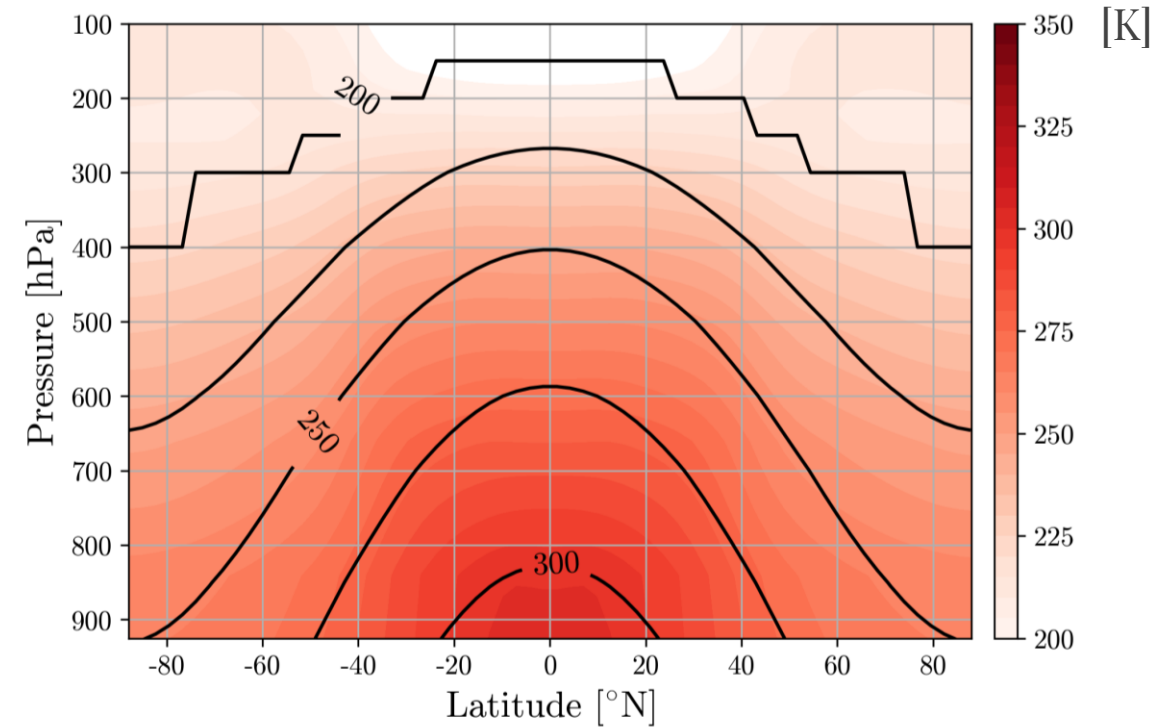
Dry Dynamical Core Models

decreasing complexity →

Zonal Wu & Reichler (2018)



Held & Suarez (1994)



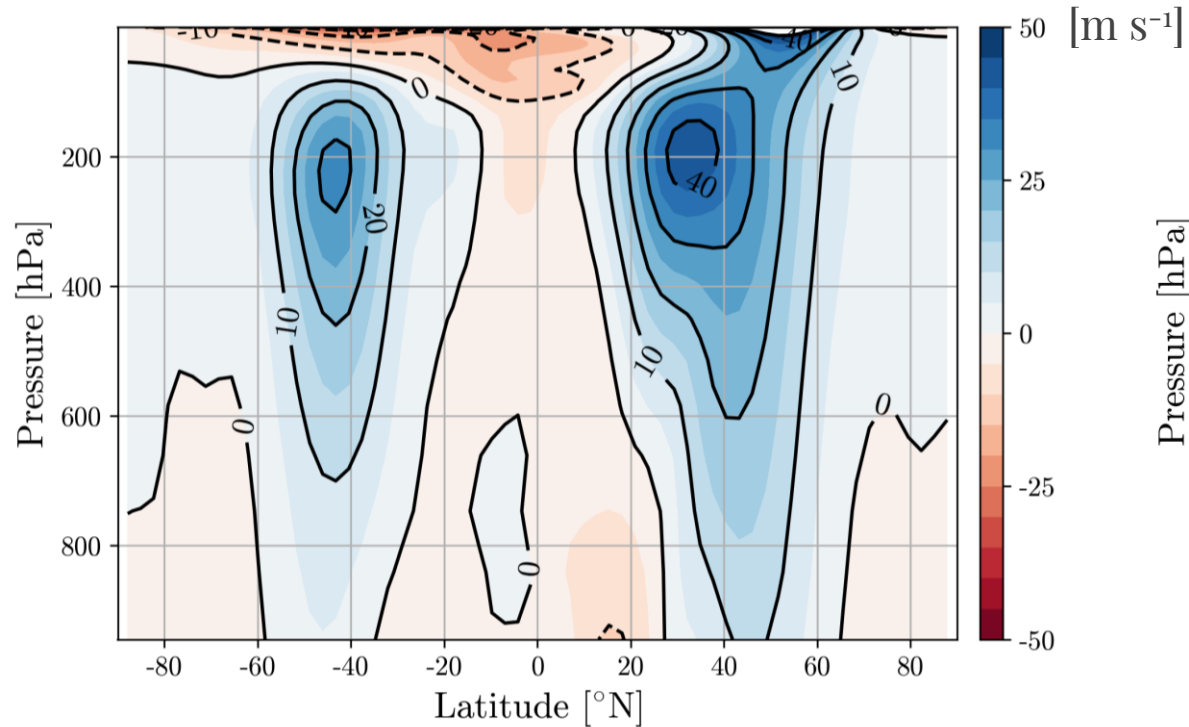
DJF
Climatology

Simulated Temperature,
Equilibrium Temperature (contour lines)

Dry Dynamical Core Models

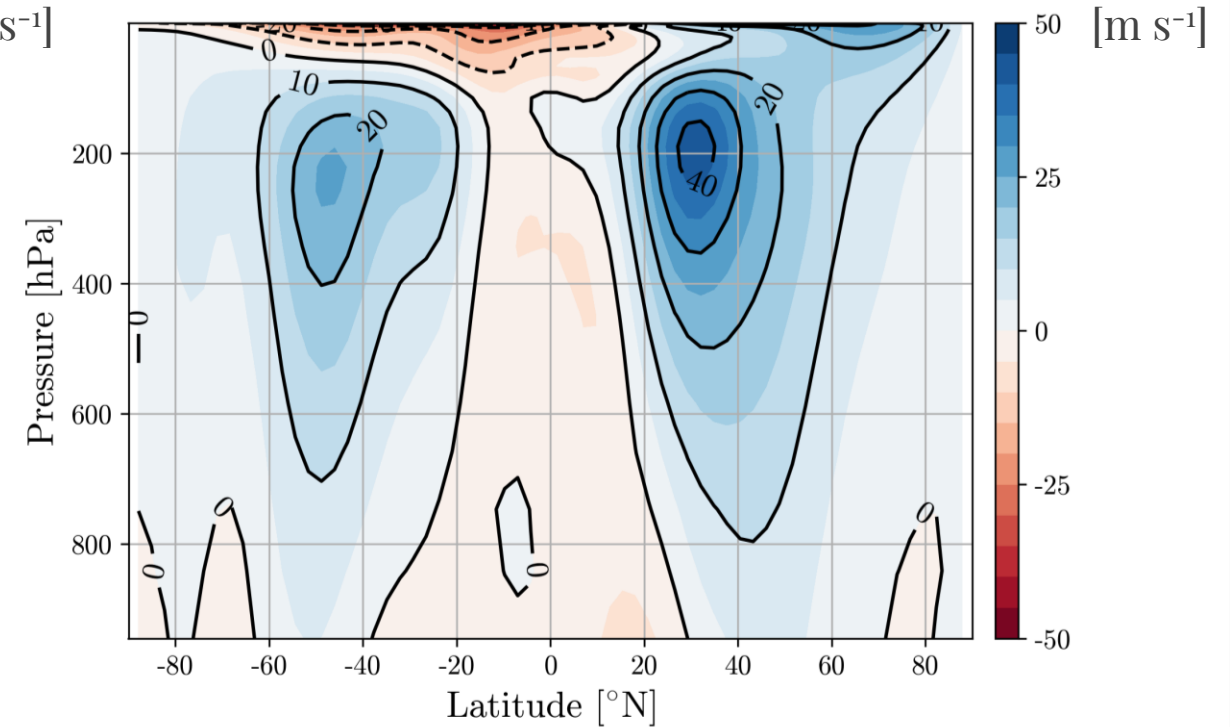
decreasing complexity →

Zonal Wu & Reichler (2018)



DJF
Climatology

3D Wu & Reichler (2018)

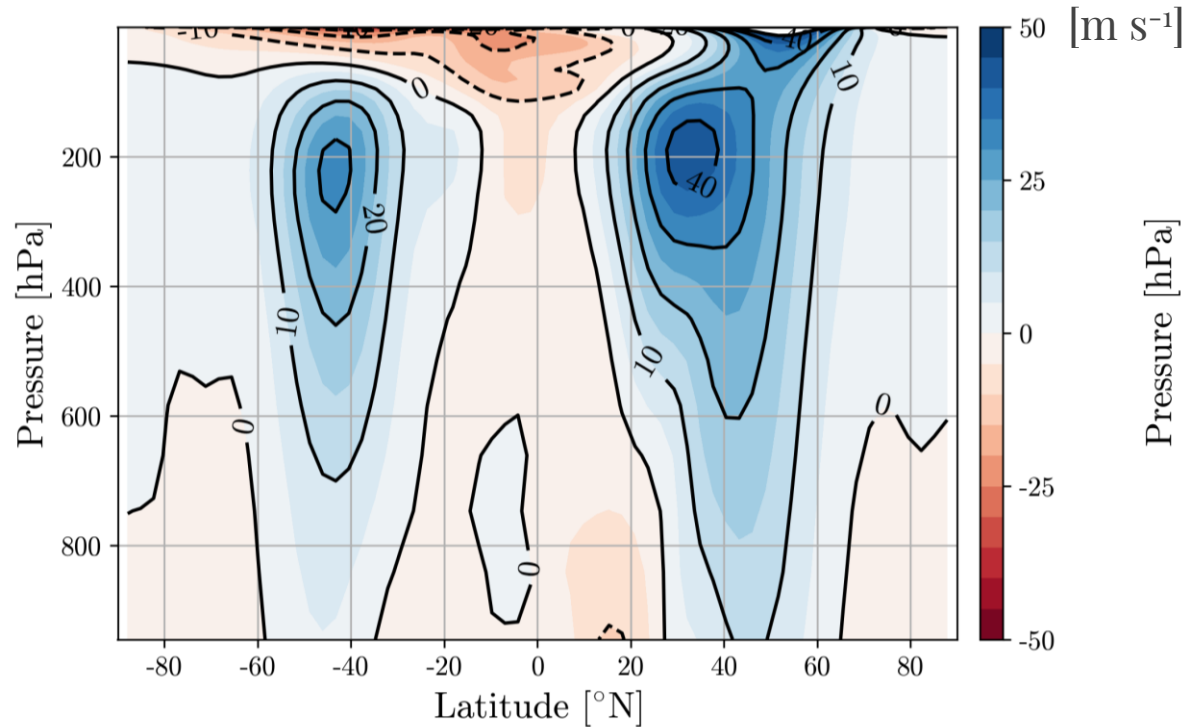


Simulated Wind

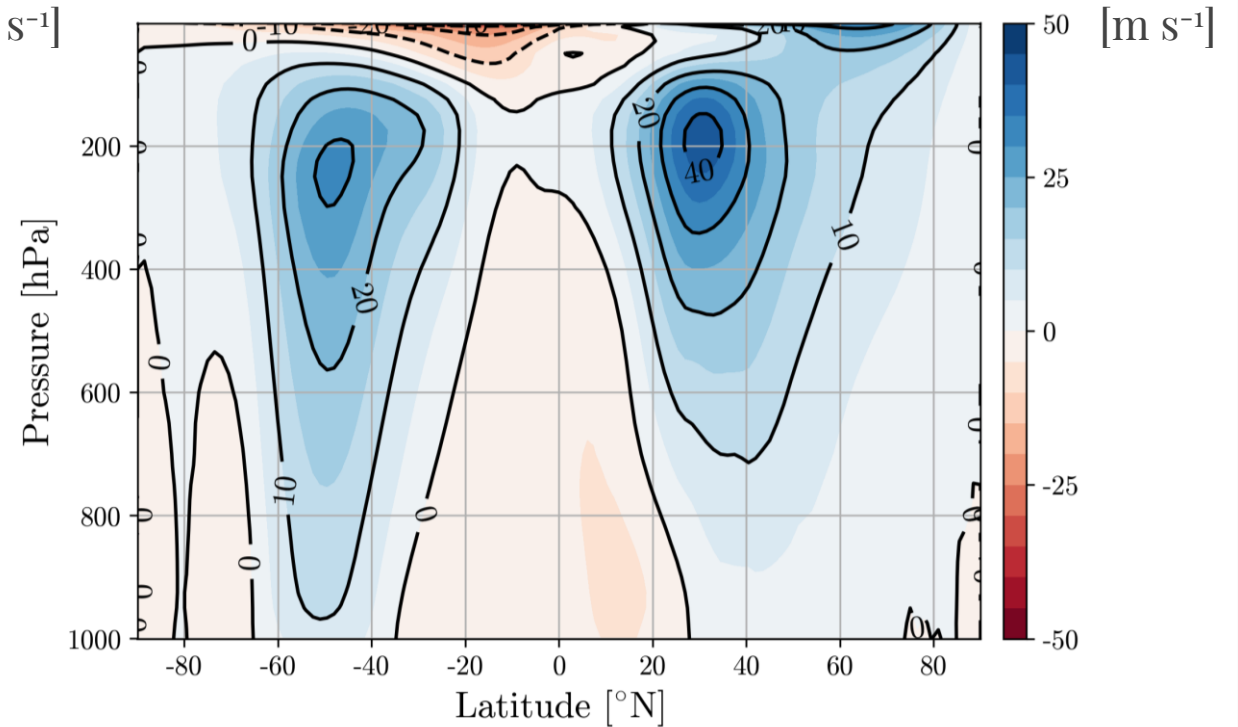
Dry Dynamical Core Models

decreasing complexity →

Zonal Wu & Reichler (2018)



ERA-interim Reanalysis

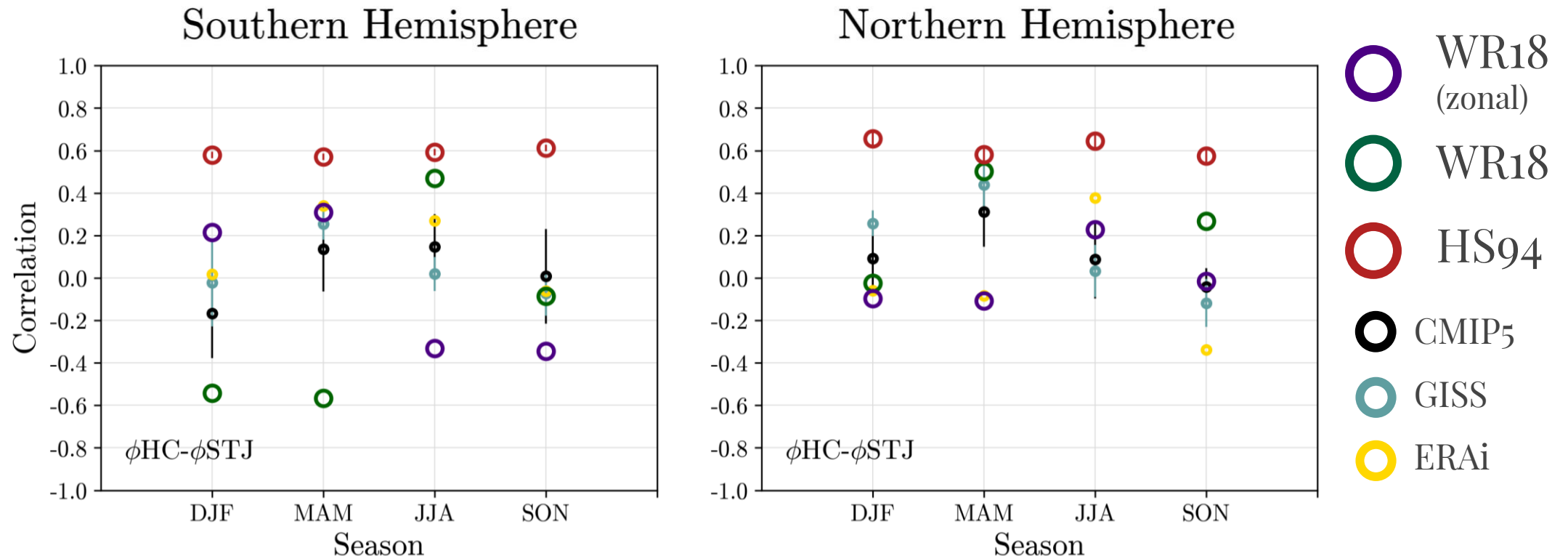


DJF
Climatology

Simulated Wind

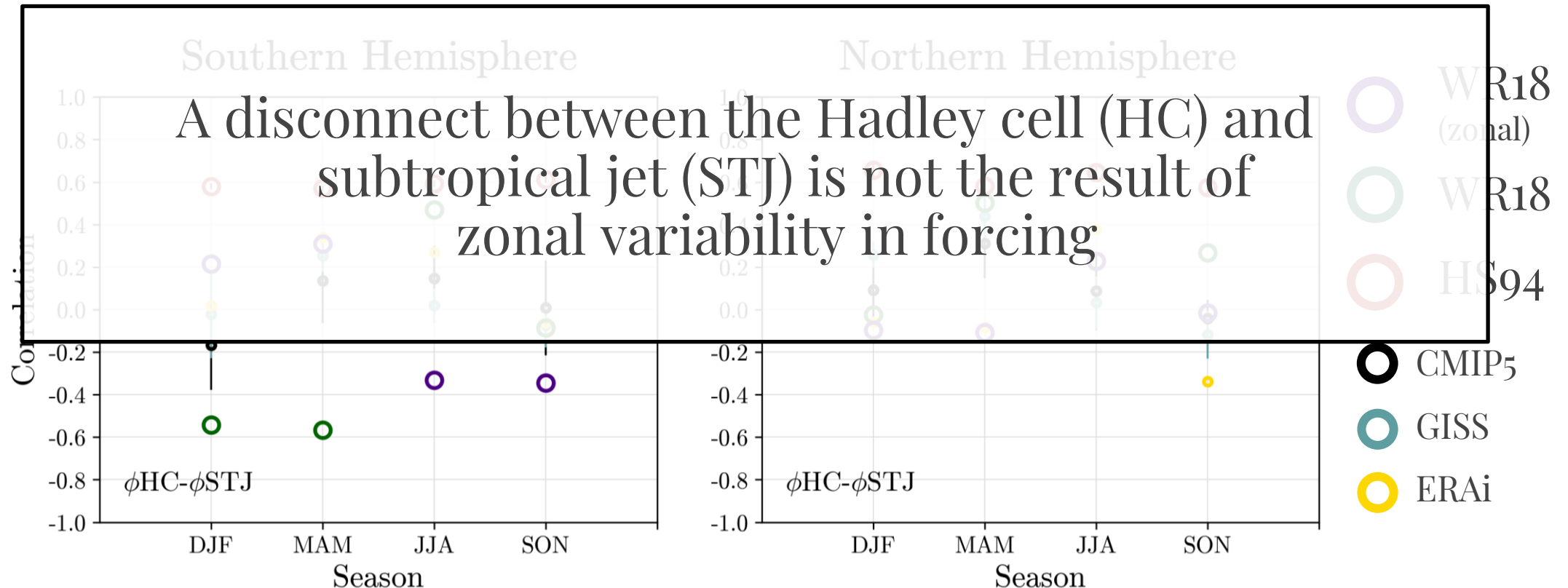
Dry Dynamical Core Models

- Zonal Wu & Reichler (2018)
Derived by iteration, zonally-symmetric, seasonally-varying T_{eq}



Dry Dynamical Core Models

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Derived by iteration, zonally-symmetric, seasonally-varying T_{eq}



Dry Dynamical Core Models

The Hadley cell (HC)
and subtropical jet
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BUT!



Dry Dynamical Core Models

The Hadley cell (HC) and subtropical jet (STJ) are positively correlated in the classical Held & Suarez (1994) setup

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A disconnect between the HC and STJ can occur in a dry model

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It is not the result of zonal variability in the forcing

Dry Dynamical Core Models

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BUT! →

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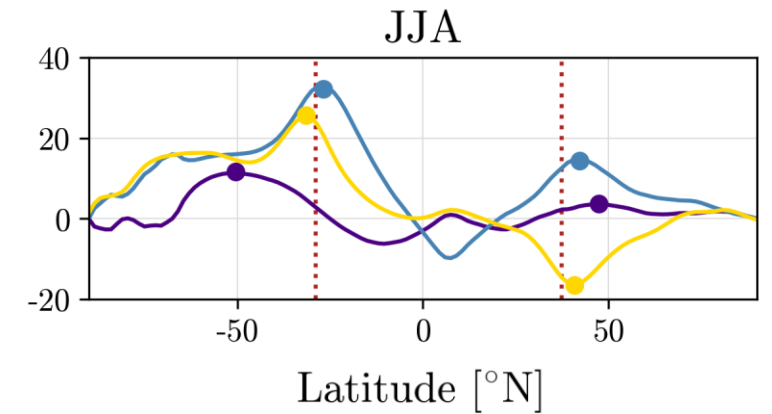
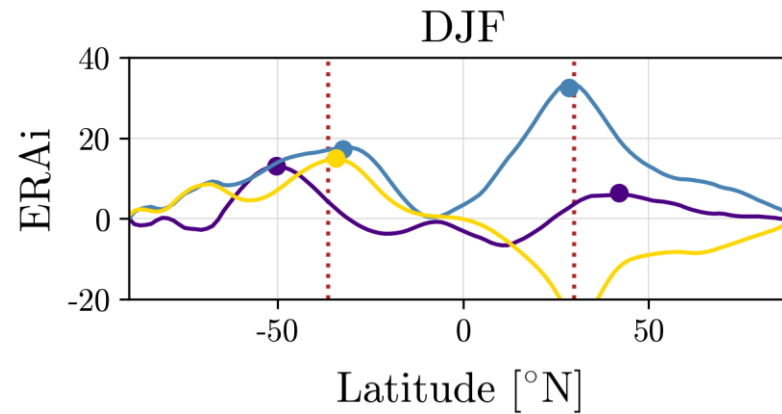
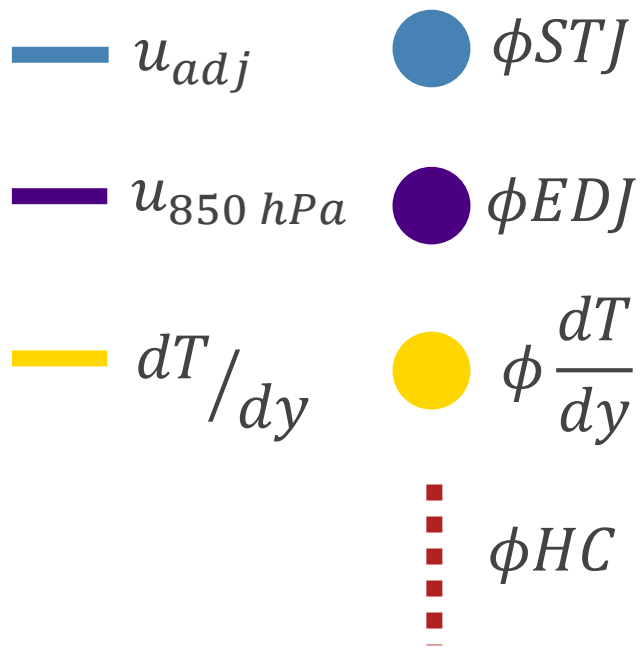
It is not the result of zonal variability in the forcing

Follow-up Question:

What are the differences in the basic state?

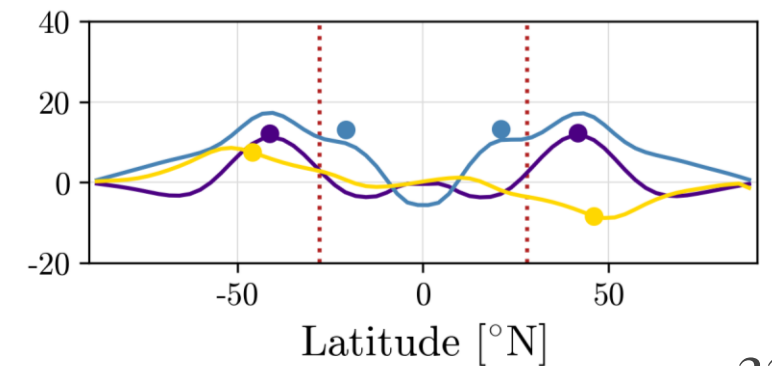
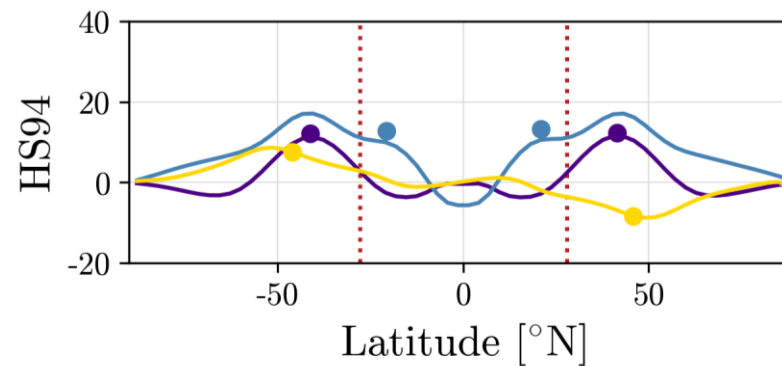
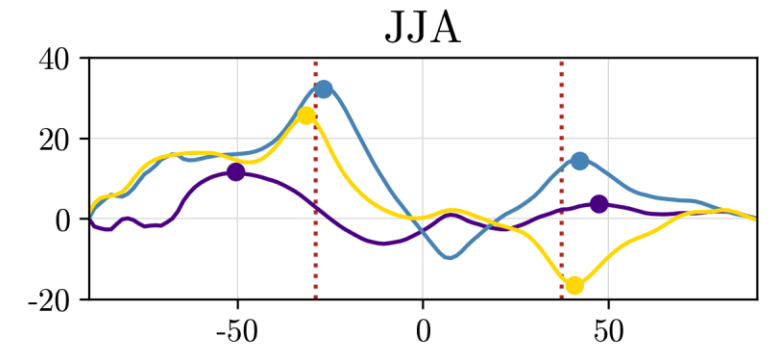
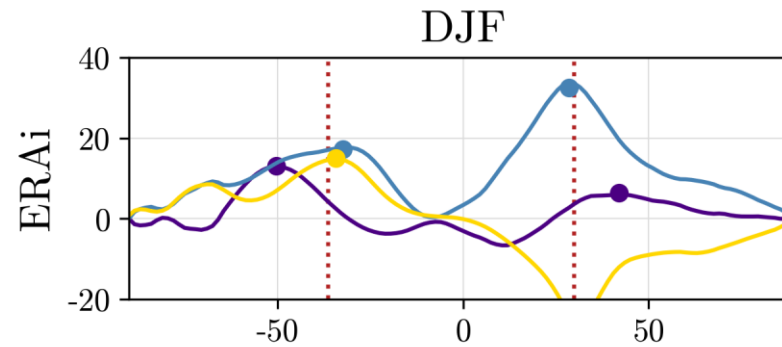
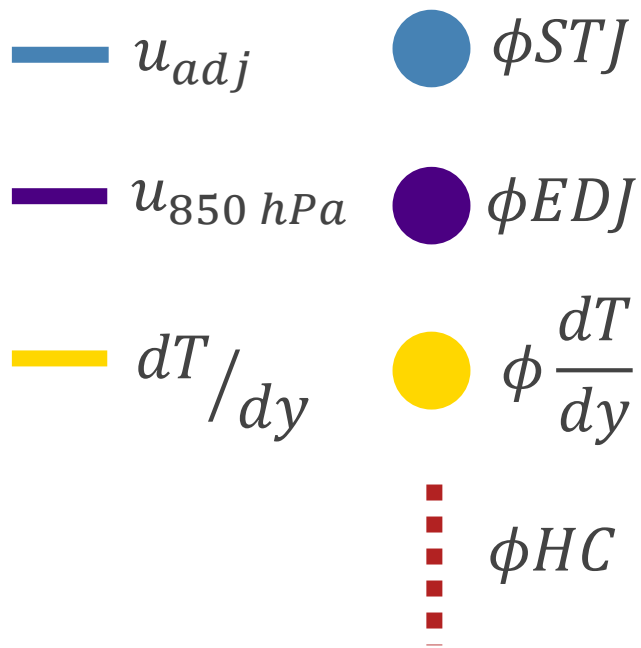
Dry Dynamical Core Models

Basic State Analysis



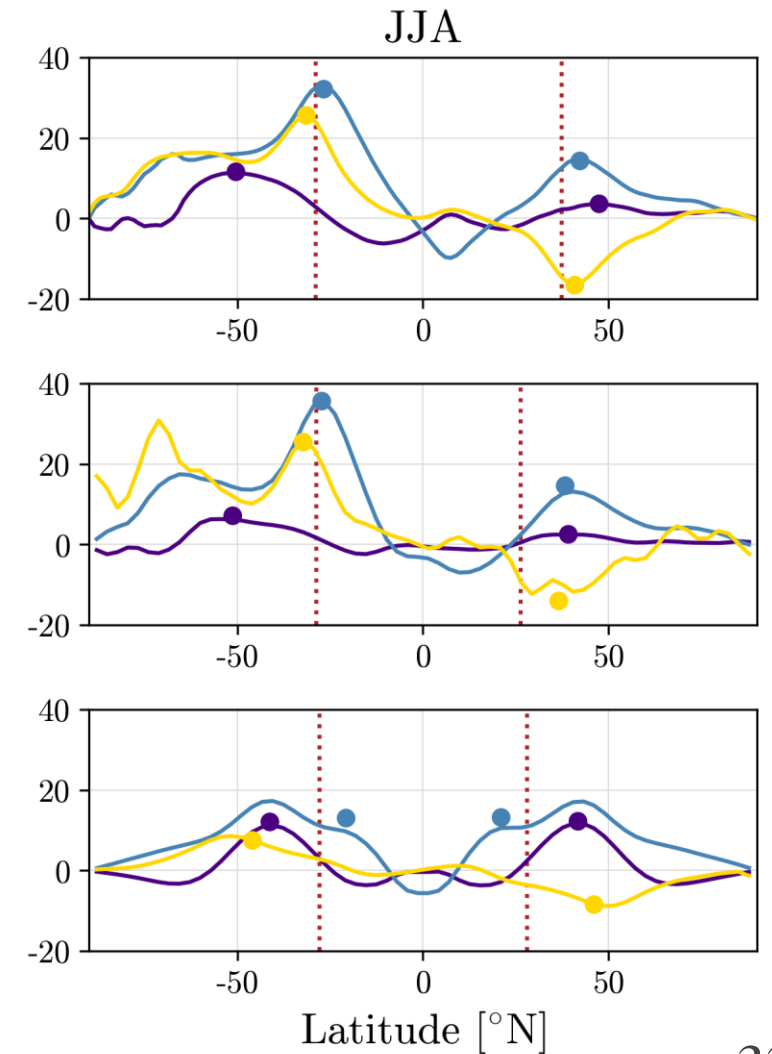
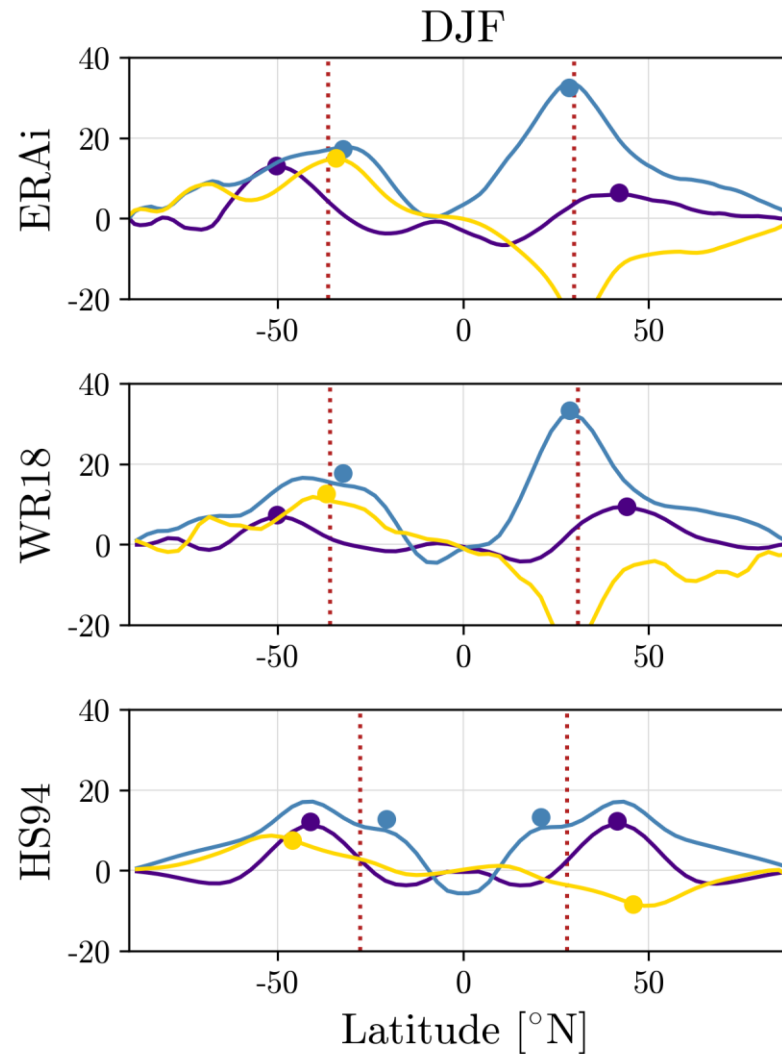
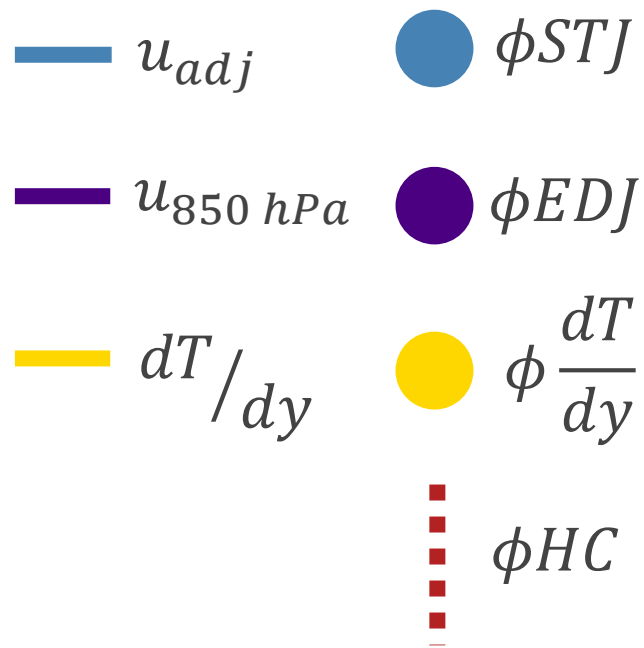
Dry Dynamical Core Models

Basic State Analysis



Dry Dynamical Core Models

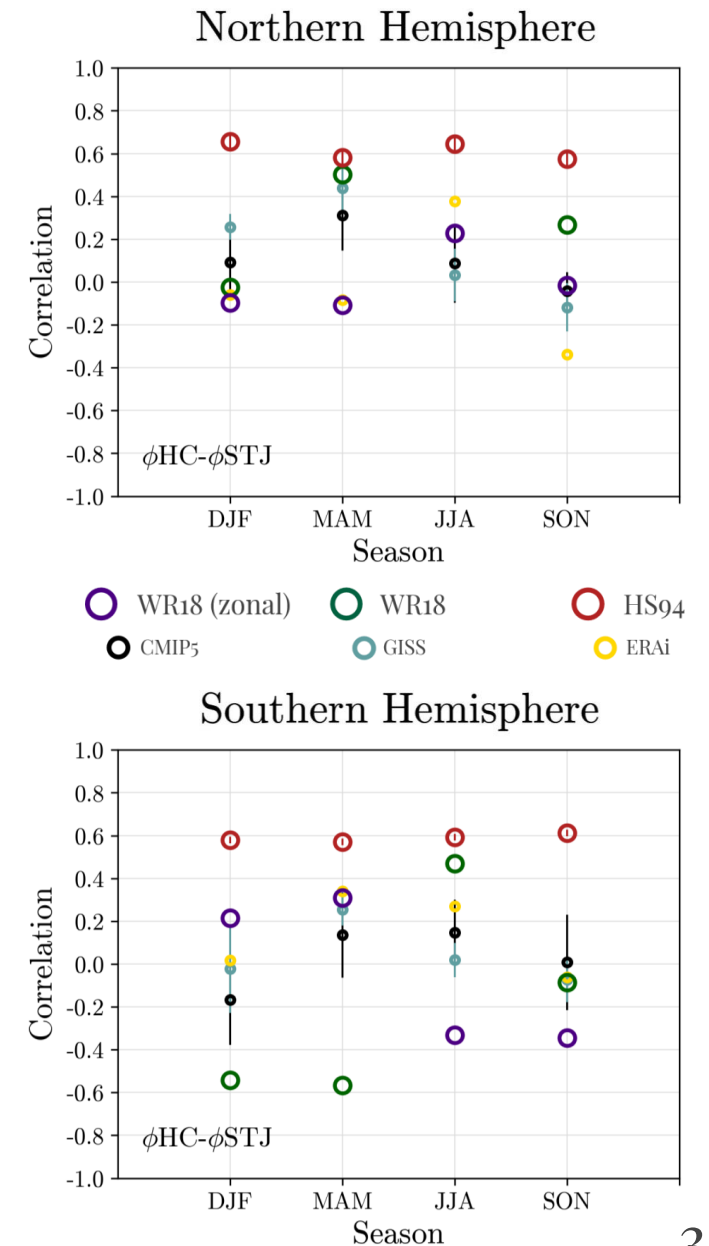
Basic State Analysis



Idealized Modelling

Conclusion:

A disconnect between the STJ and HC can occur in a fully dry atmospheric model when the model simulates a realistic STJ



Conclusions



Conclusions

Key Points:

1. The subtropical jet (STJ) is not coupled to the Hadley Cell (HC)

	Southern Hemisphere					Northern Hemisphere				
	ANN	DJF	MAM	JJA	SON	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.07	-0.1	0.1	0.12	-0.03	0.15	0.02	0.29*	0.2	-0.08
ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)

Conclusions

Key Points:

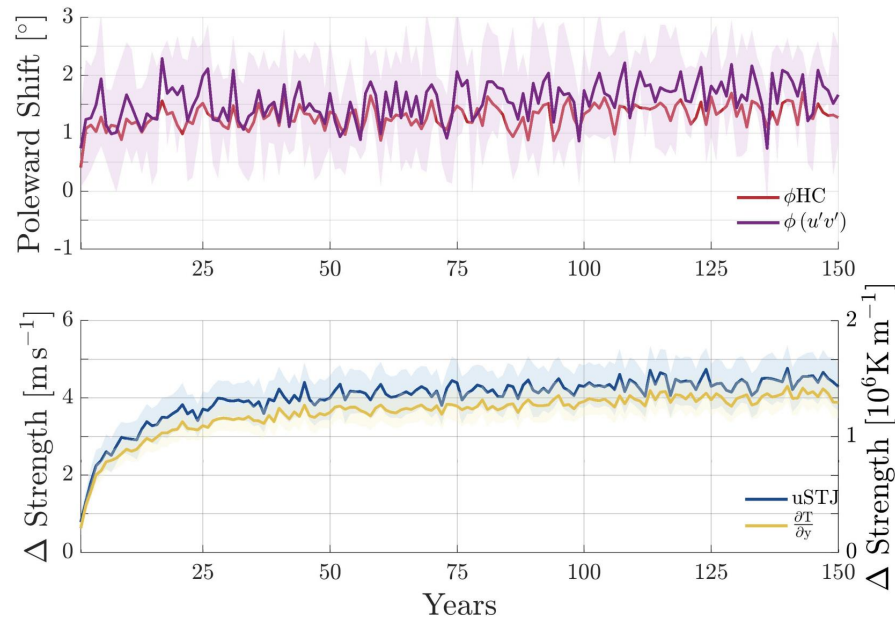
1. The subtropical jet (STJ) is not coupled to the Hadley Cell (HC)
2. The HC and STJ's behavior is consistent with physical balances observed in the atmosphere

Southern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.07	-0.1	0.1	0.12	-0.03
ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)

Northern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.15	0.02	0.29*	0.2	-0.08
ϕ_{STJ}	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)



Conclusions

Key Points:

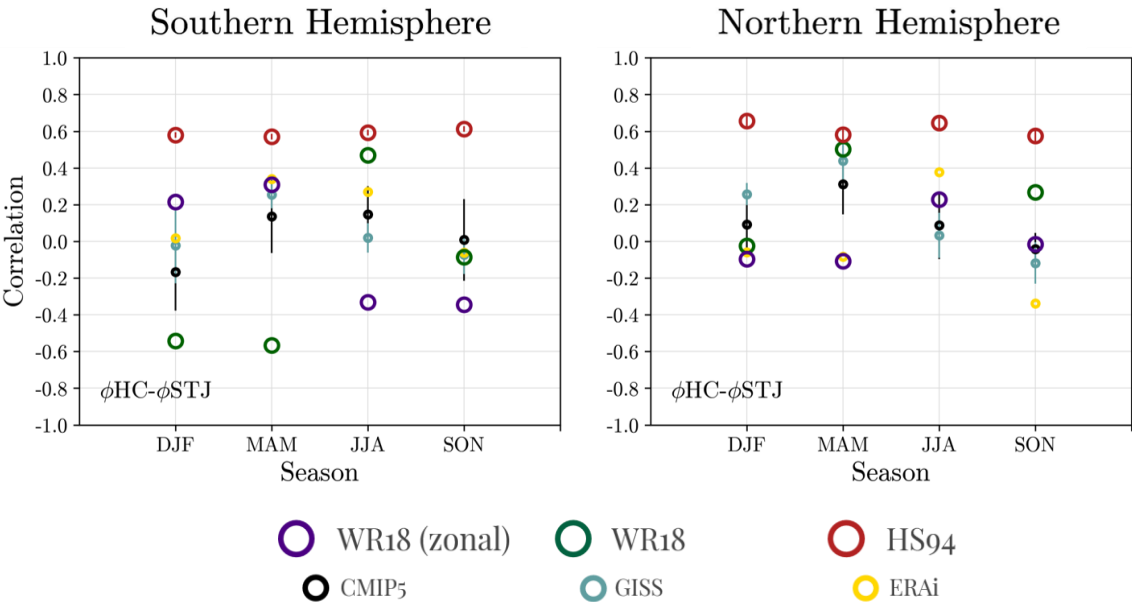
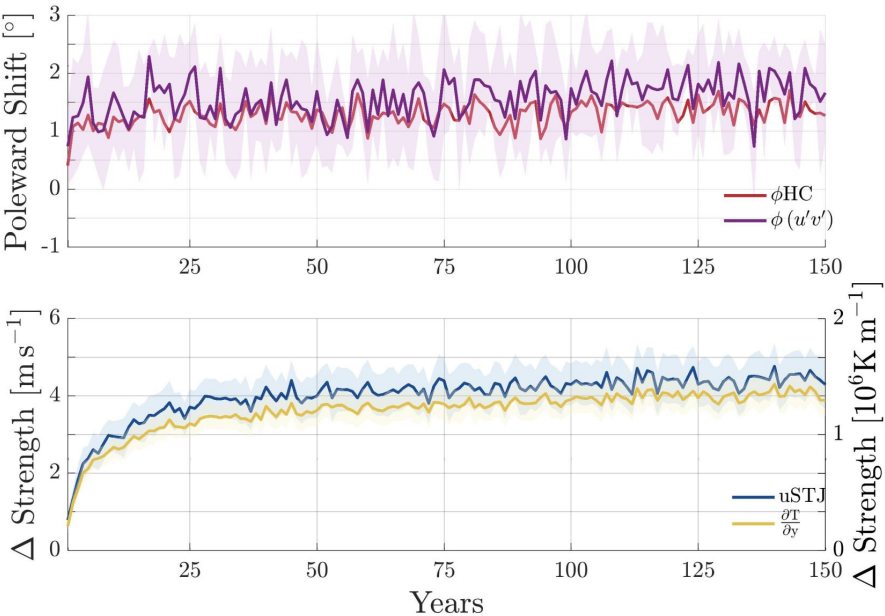
1. The subtropical jet (STJ) is not coupled to the Hadley Cell (HC)
2. The HC and STJ's behavior is consistent with physical balances observed in the atmosphere
3. A disconnect between the STJ and HC can occur in a fully dry atmospheric model when the model simulates a realistic STJ

Southern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.07	-0.1	0.1	0.12	-0.03
ϕ_{STJ}	(0.23)	(0.3)	(0.22)	(0.15)	(0.22)

Northern Hemisphere

	ANN	DJF	MAM	JJA	SON
ϕ_{HC}	0.15	0.02	0.29*	0.2	-0.08
ϕ_{STJ}	(0.18)	(0.12)	(0.16)	(0.17)	(0.09)



Conclusions

Key Points:

1. The subtropical jet (STJ) is not coupled to the Hadley Cell (HC)
2. The HC and STJ's behavior is consistent with physical balances observed in the atmosphere
3. A disconnect between the STJ and HC can occur in a fully dry atmospheric model when the model simulates a realistic STJ

Hypothesis:

A coupling between the Hadley cell and subtropical jet occurs when eddy influence on the subtropical thermal wind balance is non-negligible

Continuing Work

Hypothesis:

A coupling between the
Hadley cell (HC) and subtropical jet (STJ) occurs
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Balances in Idealized Models

What is the relationship between the midlatitude eddies and subtropical thermal wind balance in the different simulations?

Continuing Work

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Balances in Idealized Models

What is the relationship between the midlatitude eddies and subtropical thermal wind balance in the different simulations?

Non-metric Analysis

What are the limitations of metric analysis?

Will analysis of atmospheric fields provide greater insight into the system dynamics?

Continuing Work

Hypothesis:

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Questions?

Balances in Idealized Models

What is the relationship between the midlatitude eddies and subtropical thermal wind balance in the different simulations?

Non-metric Analysis

What are the limitations of metric analysis?

Will analysis of atmospheric fields provide greater insight into the system dynamics?

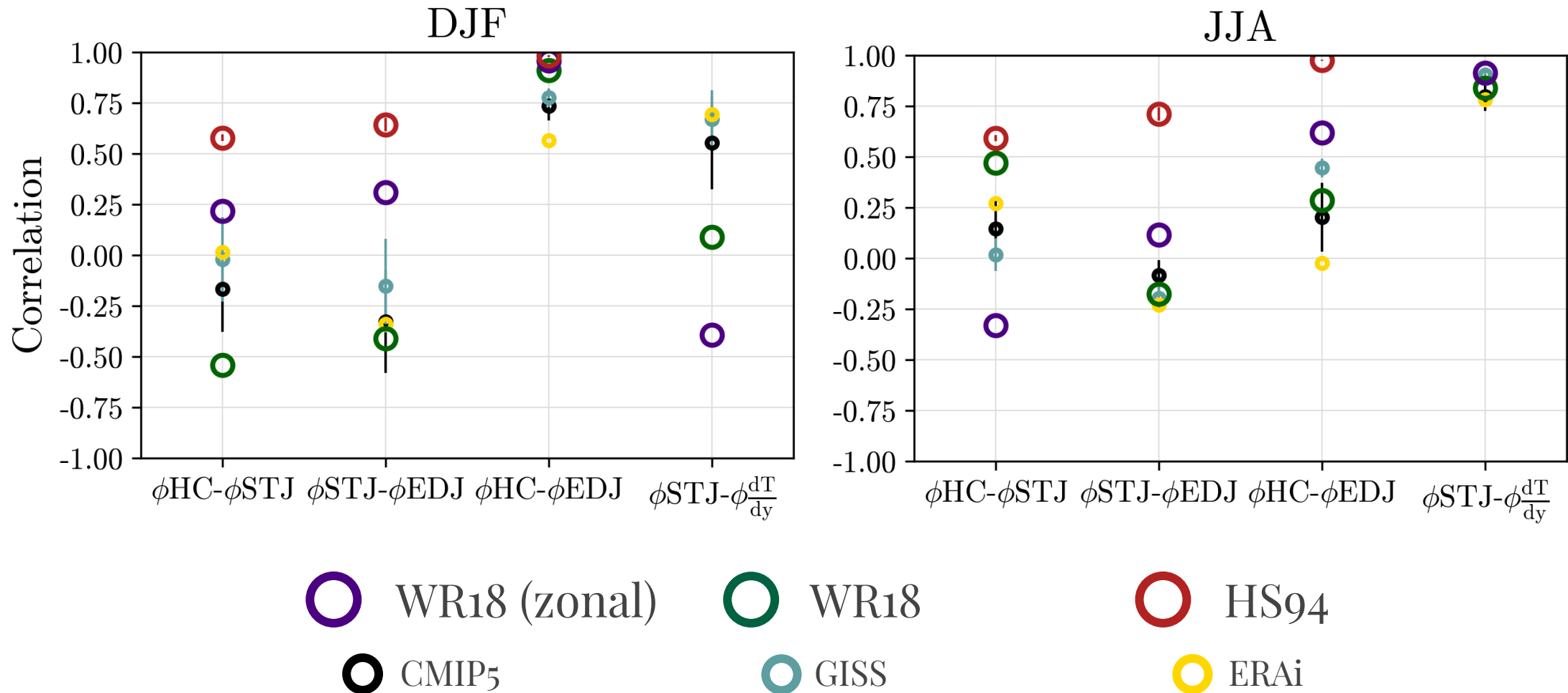
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Extra Slides

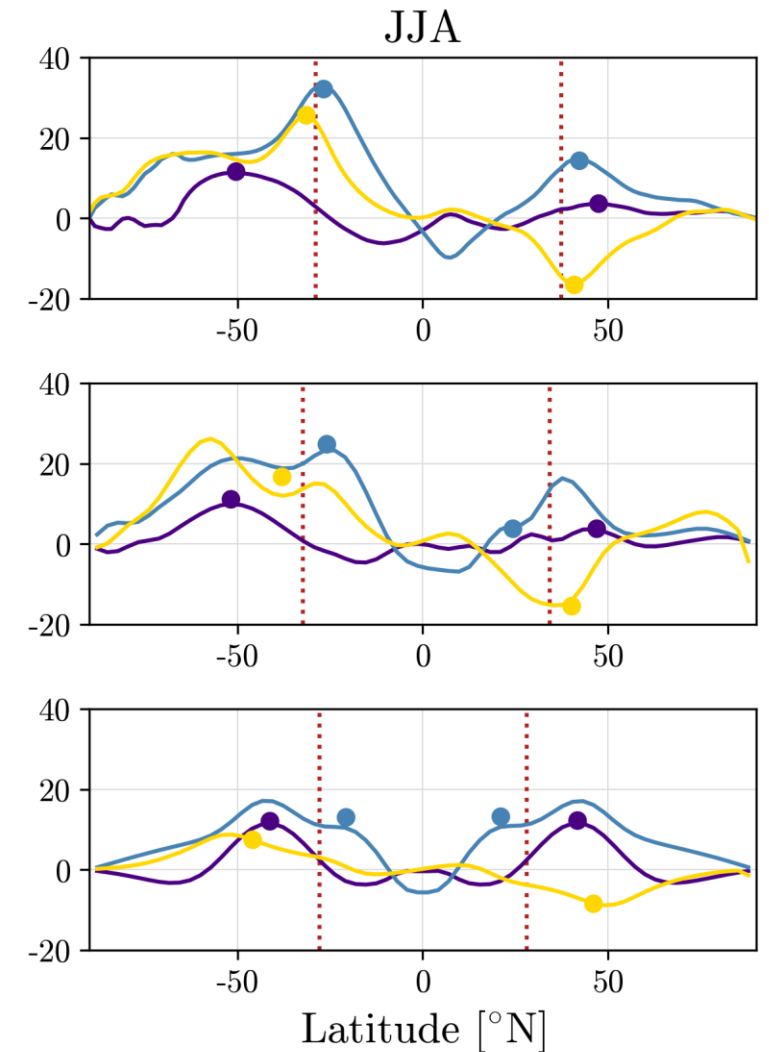
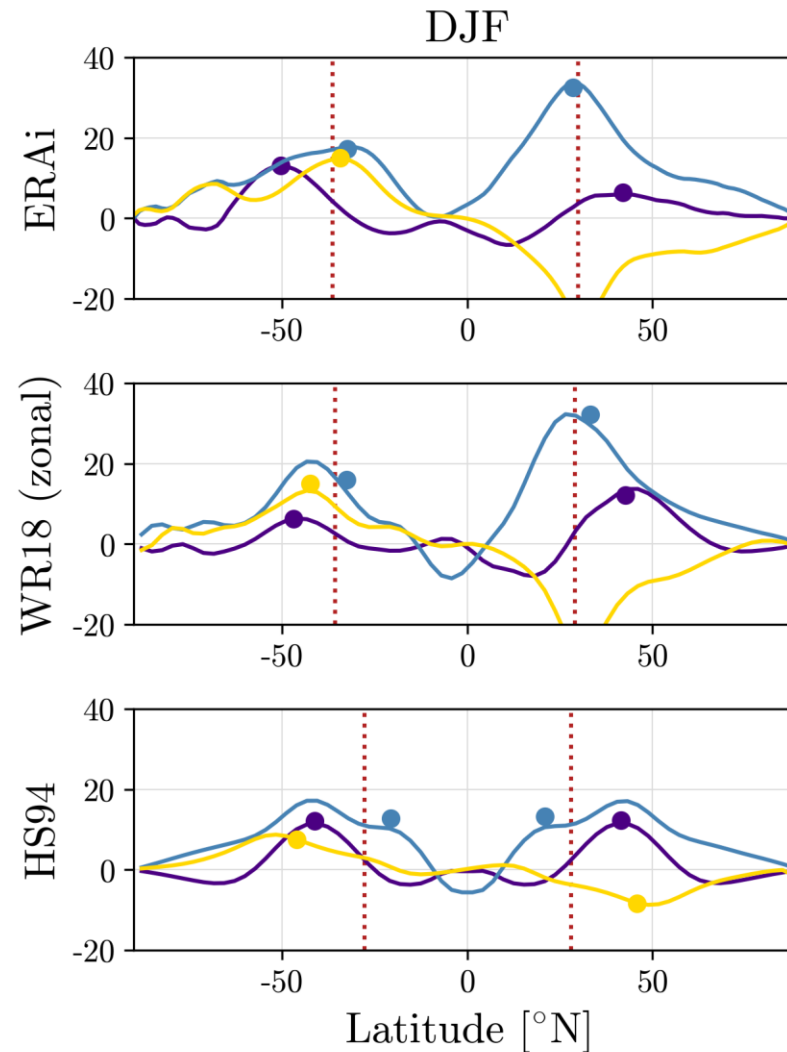
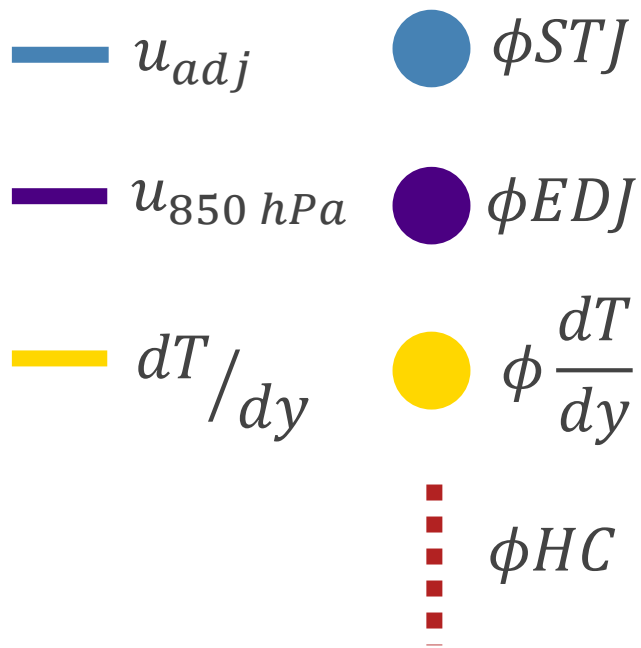


Dry Dynamical Core Models



Dry Dynamical Core Models

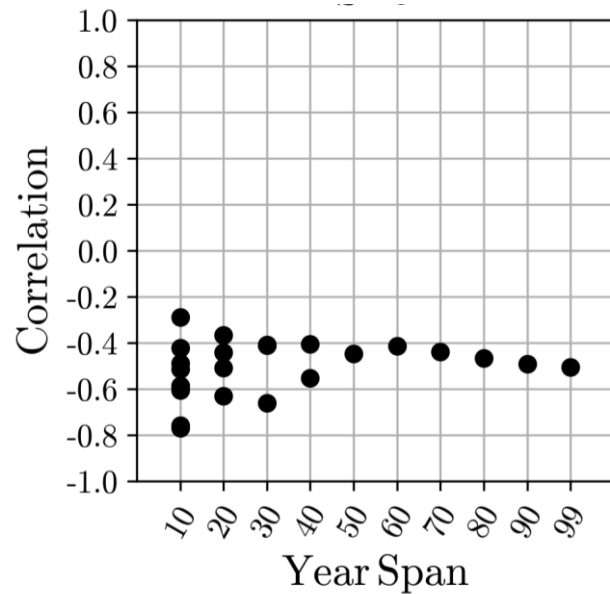
Basic State Analysis



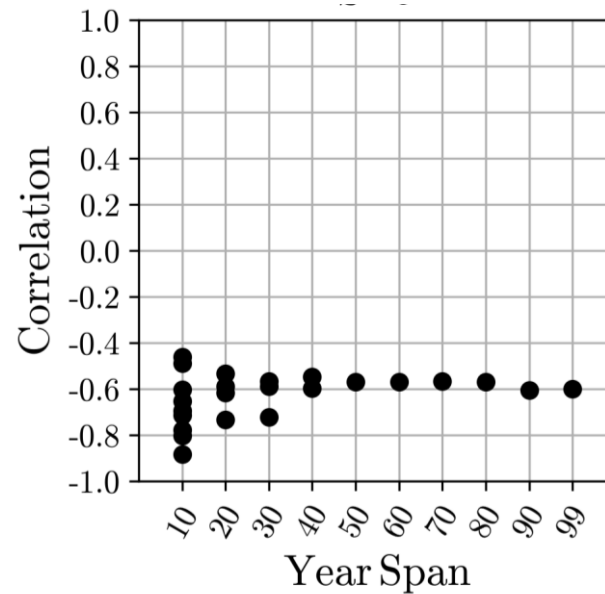
Dry Dynamical Core Models

Correlation Convergence

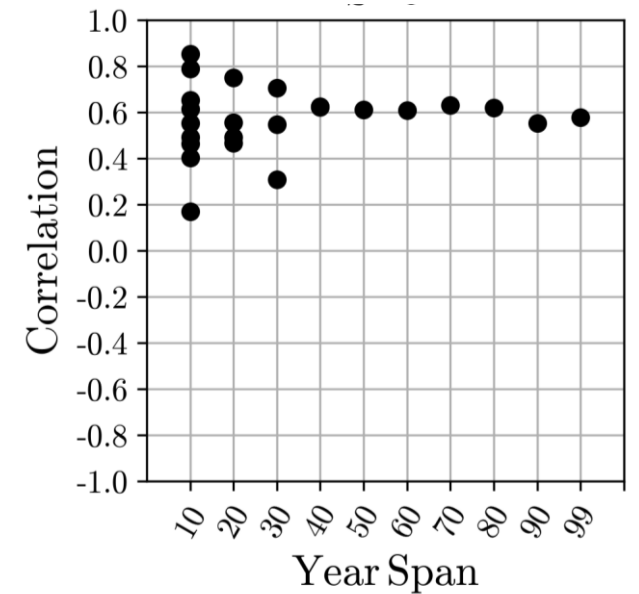
Wu & Reichler (2018)



Zonal Wu & Reichler (2018)



Held & Suarez (1994)



GISS-E2.1 Model

NASA Goddard Institute for Space Studies' Global Climate Model

abruptNxCO2

Abrupt multiplying
of CO₂, held fixed

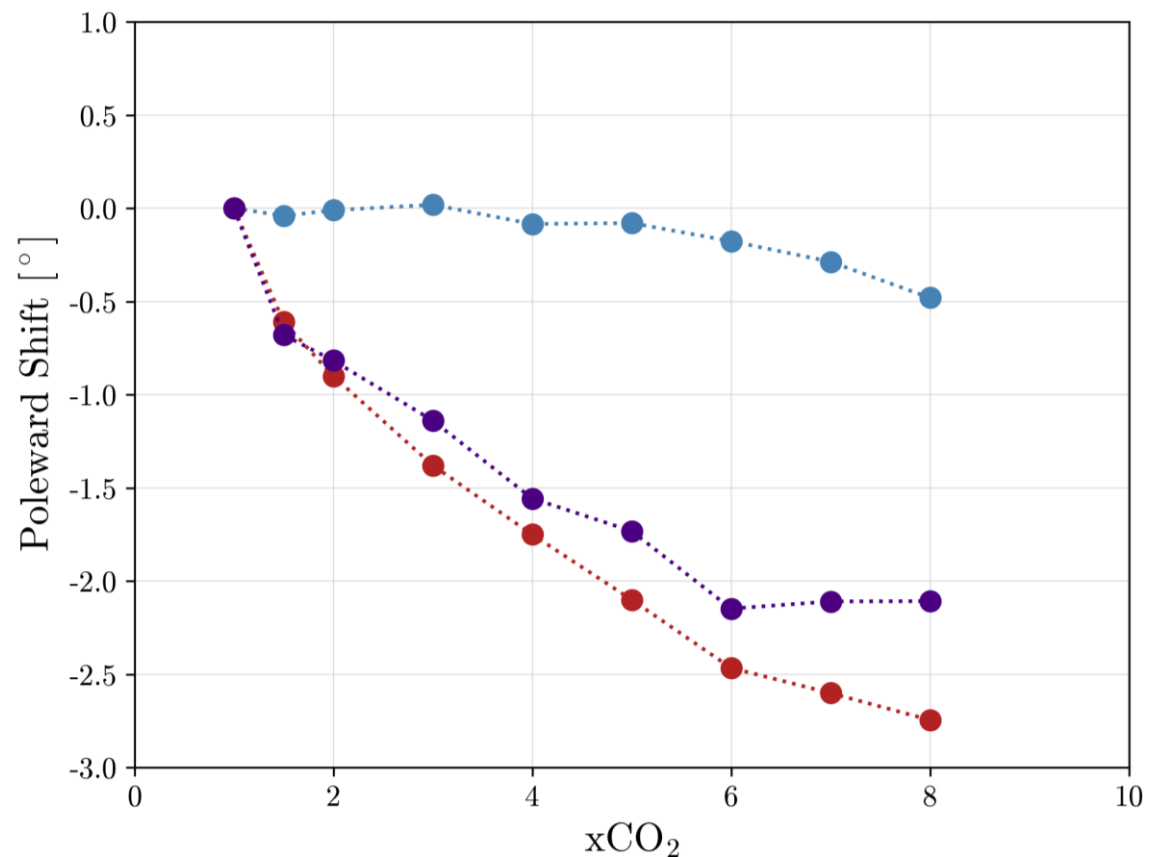
1xCO ₂	5xCO ₂
1.5xCO ₂	6xCO ₂
2xCO ₂	7xCO ₂
3xCO ₂	8xCO ₂
4xCO ₂	

provided by
Clara Orbe and Ivan Mitevski

● ϕ_{STJ}

● ϕ_{EDJ}

● ϕ_{HC}



CMIP5: Cooling

Hadley cell:

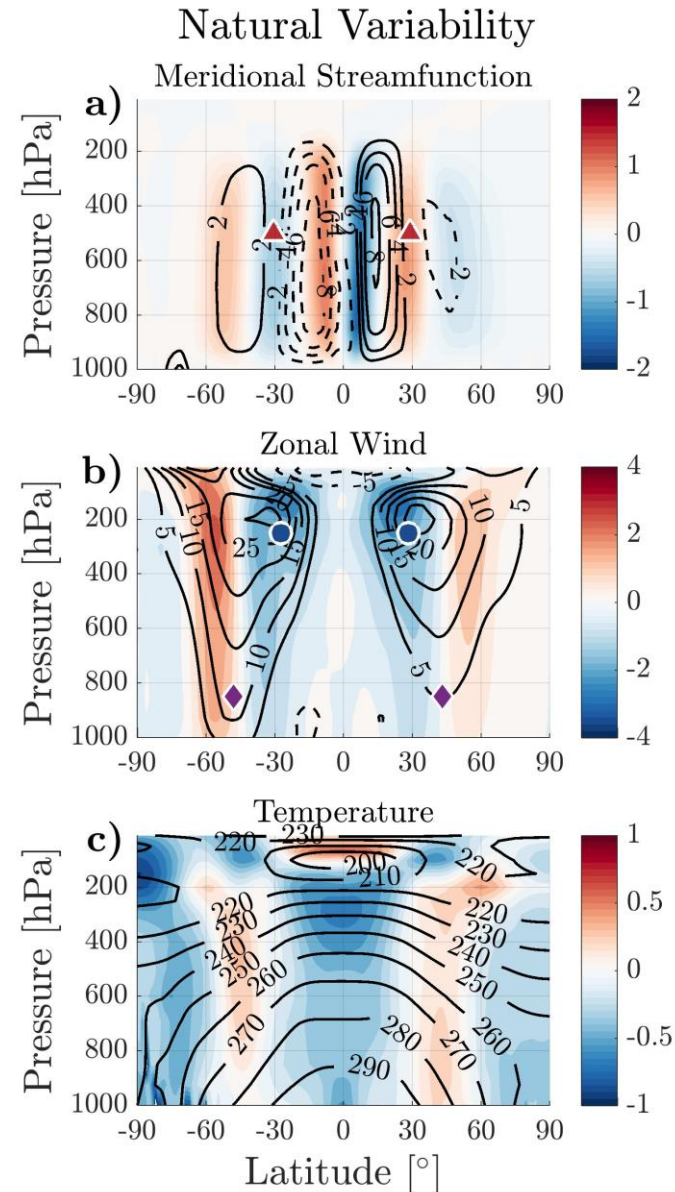
- Expands
- Weakens

Eddy-driven jet:

- Shifts poleward
- Strengthens

Subtropical jet:

- Shifts poleward
- Weakens



Menzel et al. 2019

CMIP5: Warming

Hadley cell:

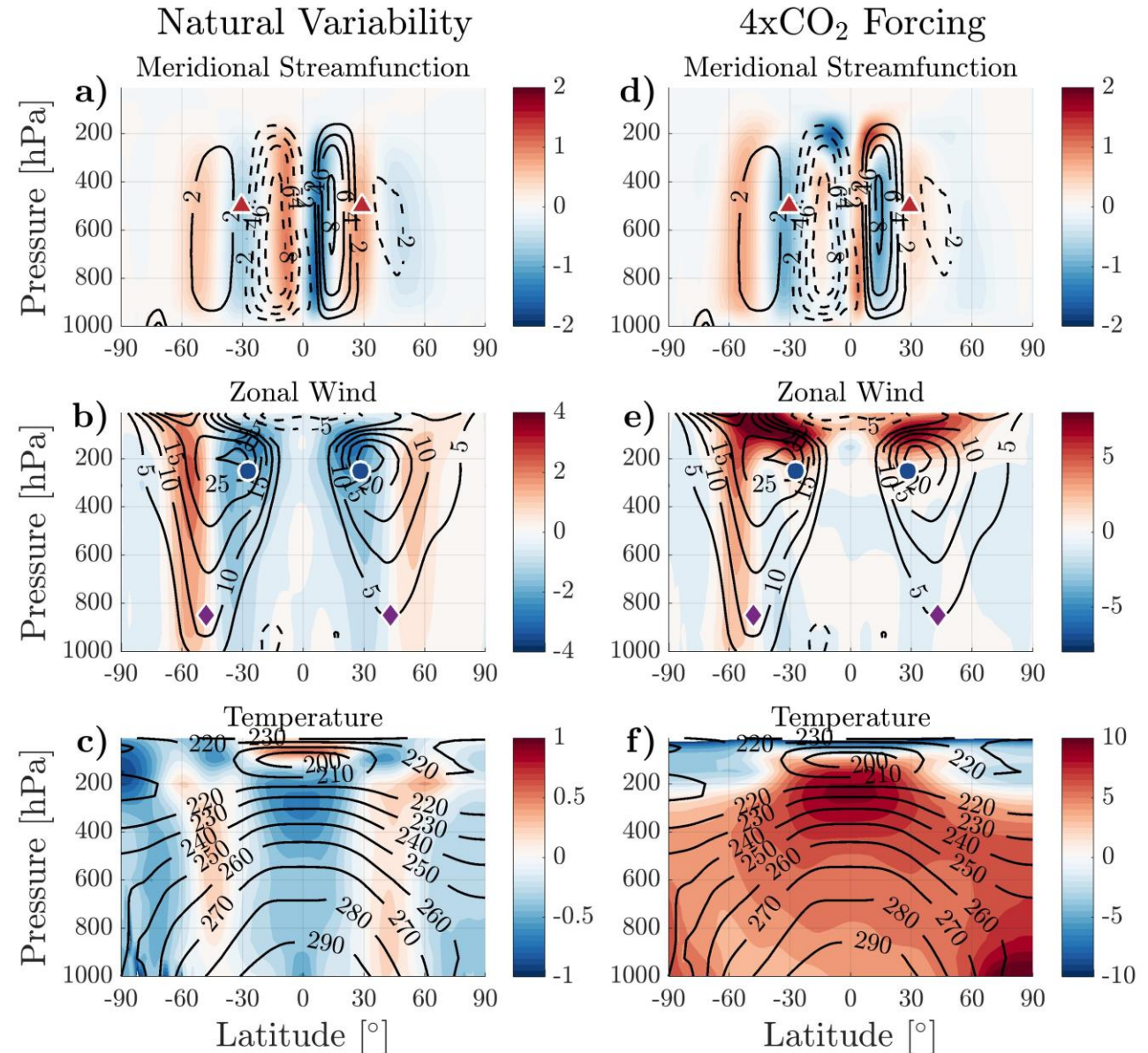
- Expands
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Eddy-driven jet:

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Subtropical jet:

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- Strengthens

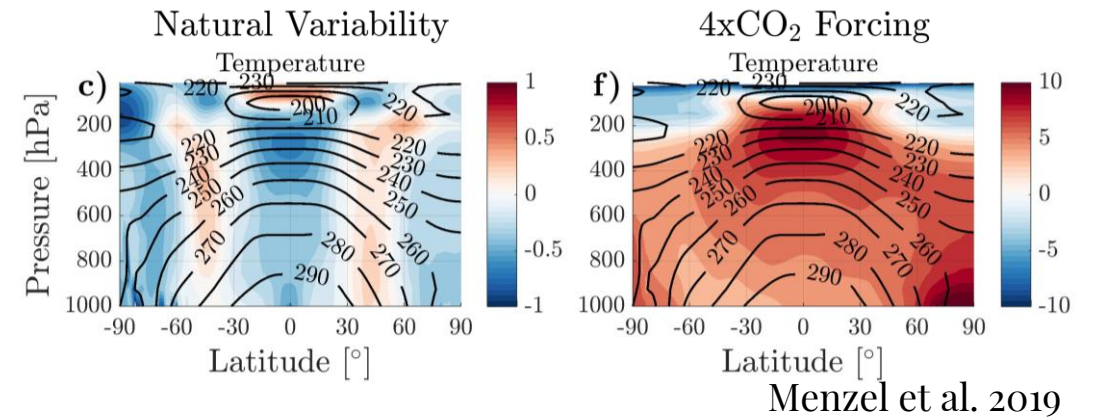
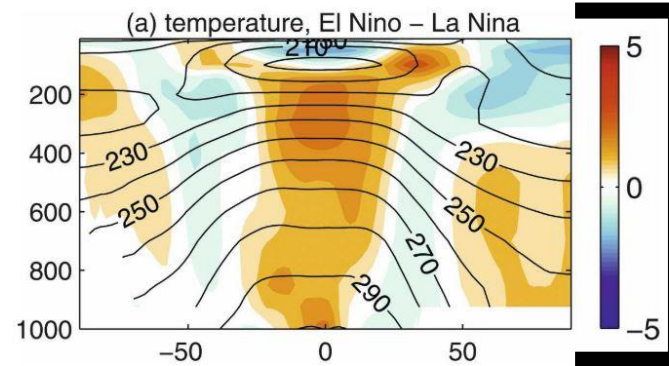


Warming Width

- Lu et al. 2008
- Sun et al. 2014
- Tandon et al. 2014

Narrow tropical
warming (ENSO)

Lu et al. 2008



more narrow
HC

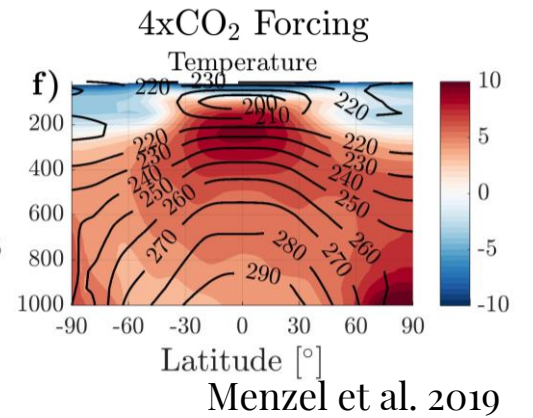
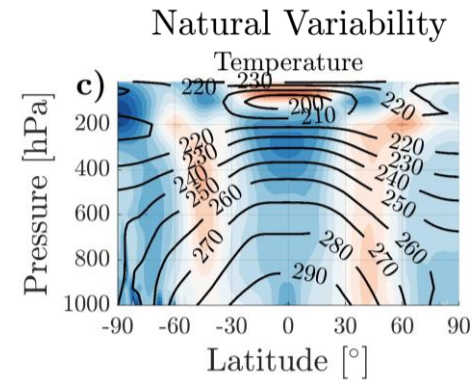
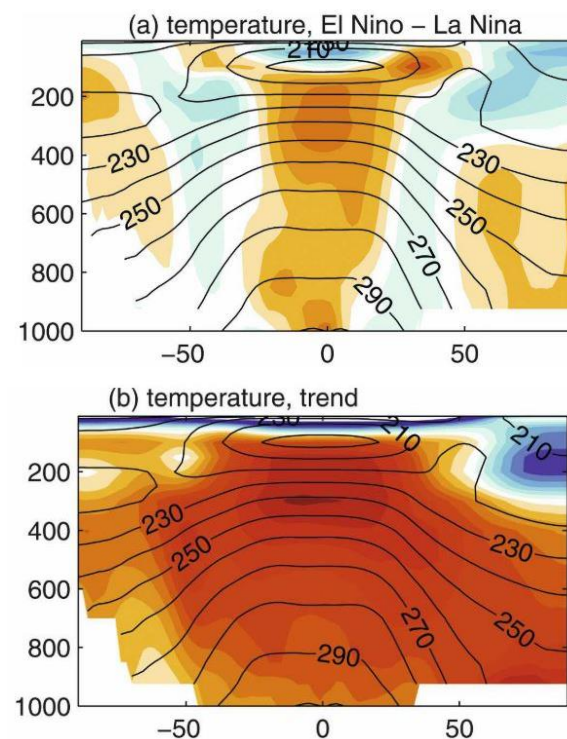
Warming Width

- Lu et al. 2008
- Sun et al. 2014
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Narrow tropical
warming (ENSO)

Broad warming
(global forcing)

Lu et al. 2008



more narrow
HC

wider HC

Warming Width

- Lu et al. 2008
- Sun et al. 2014
- Tandon et al. 2014

Narrow tropical
warming (ENSO)

Broad warming
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Lu et al. 2008

